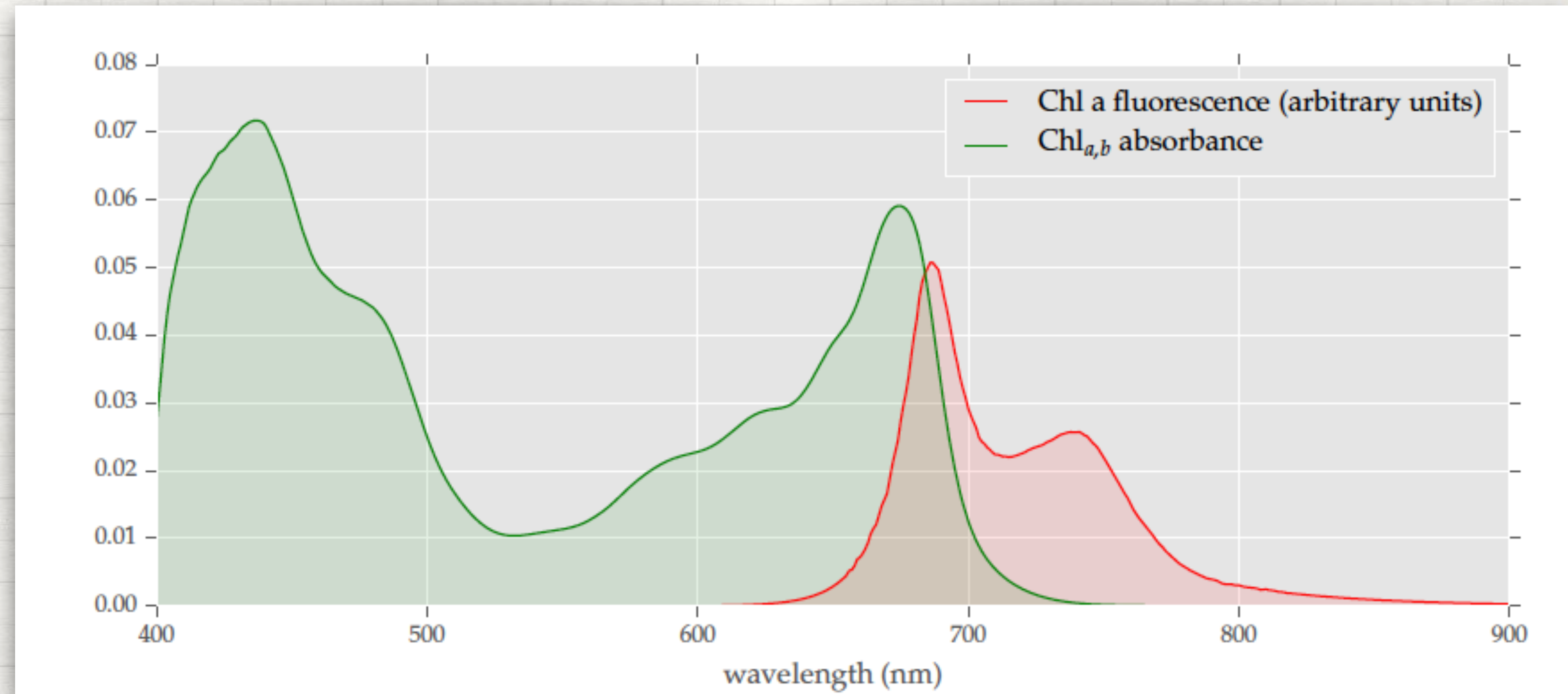
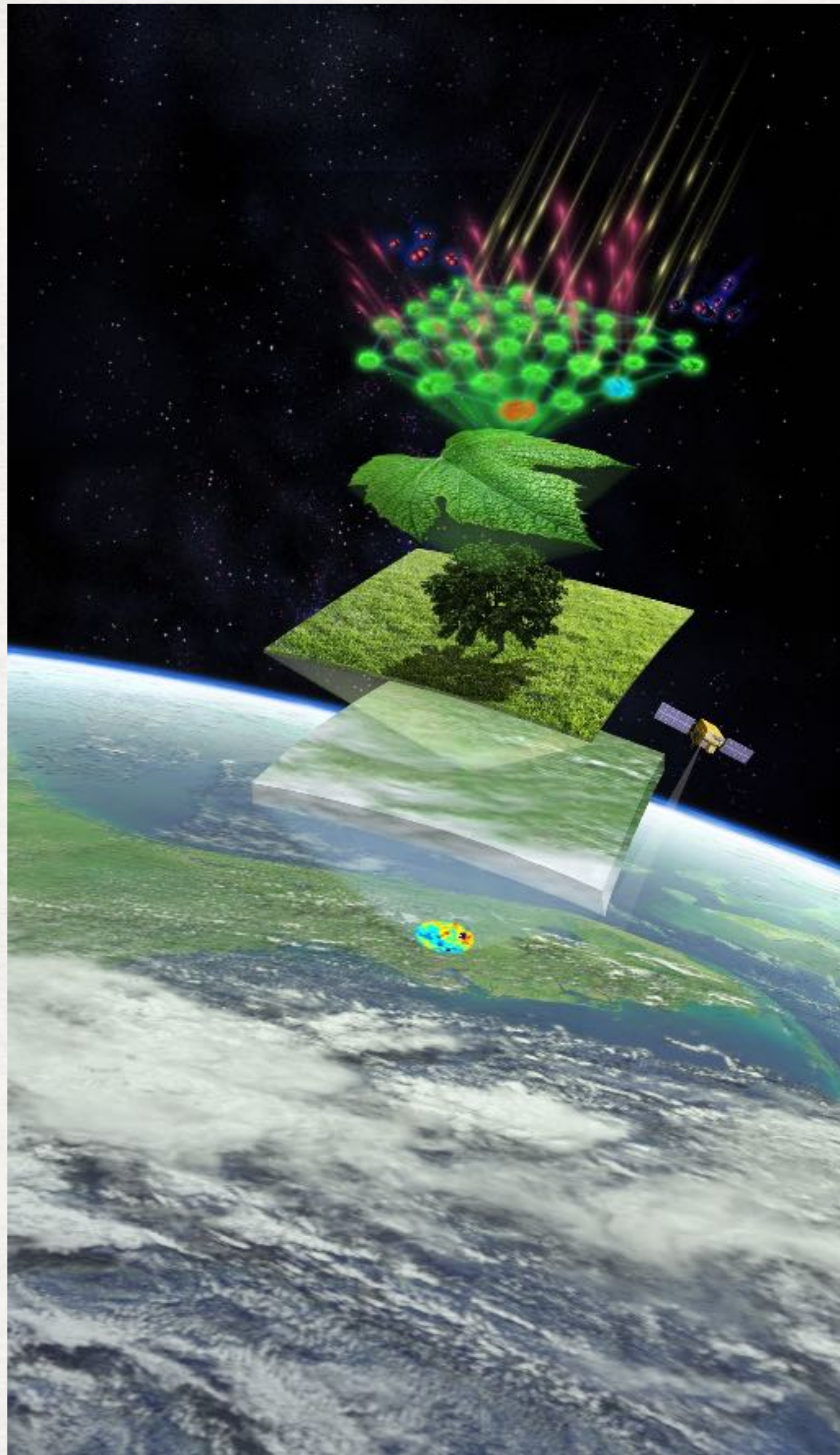


Canopy Structure, SIF and Photosynthetic Capacity

Joe Berry, Carnegie Inst. for Science



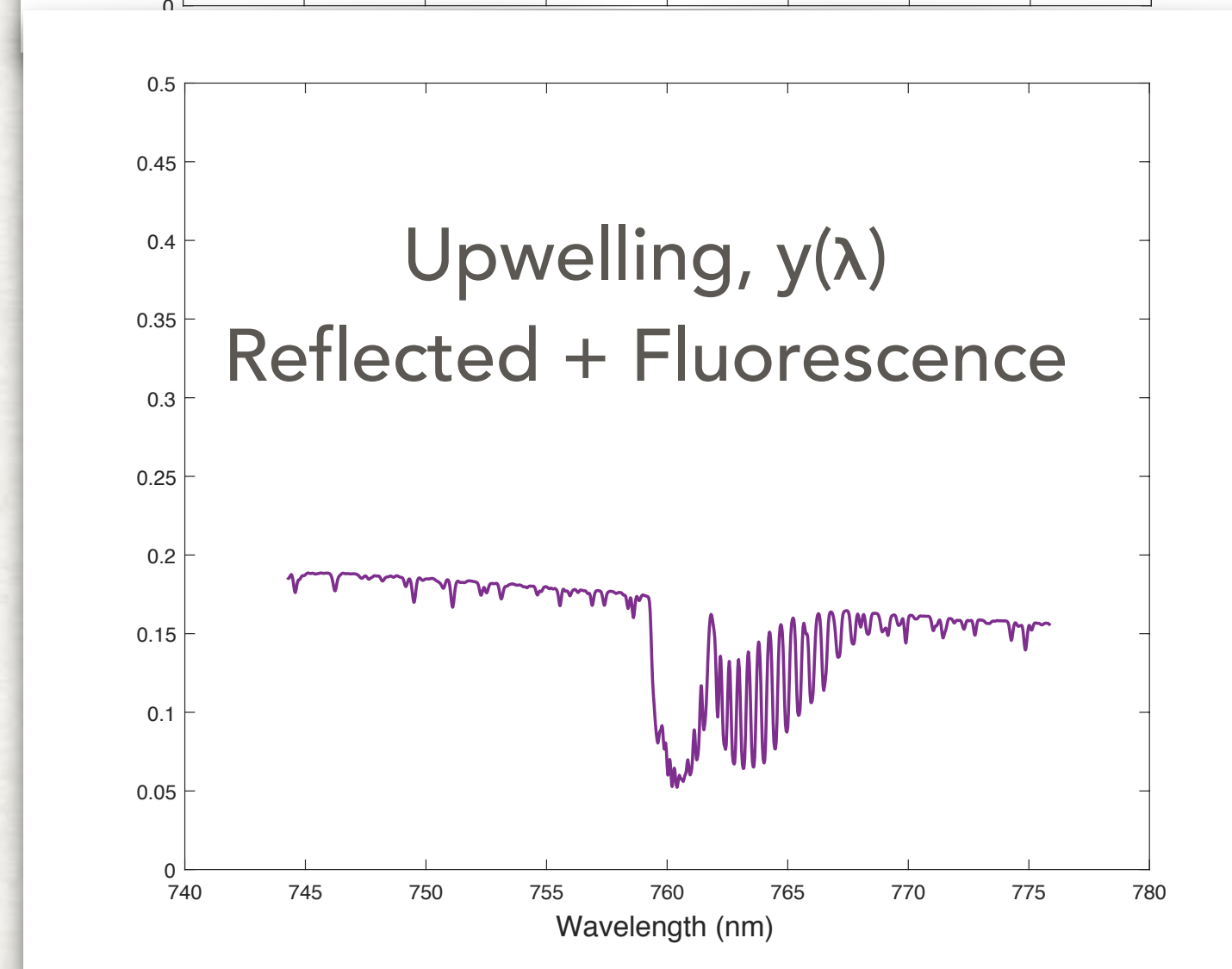
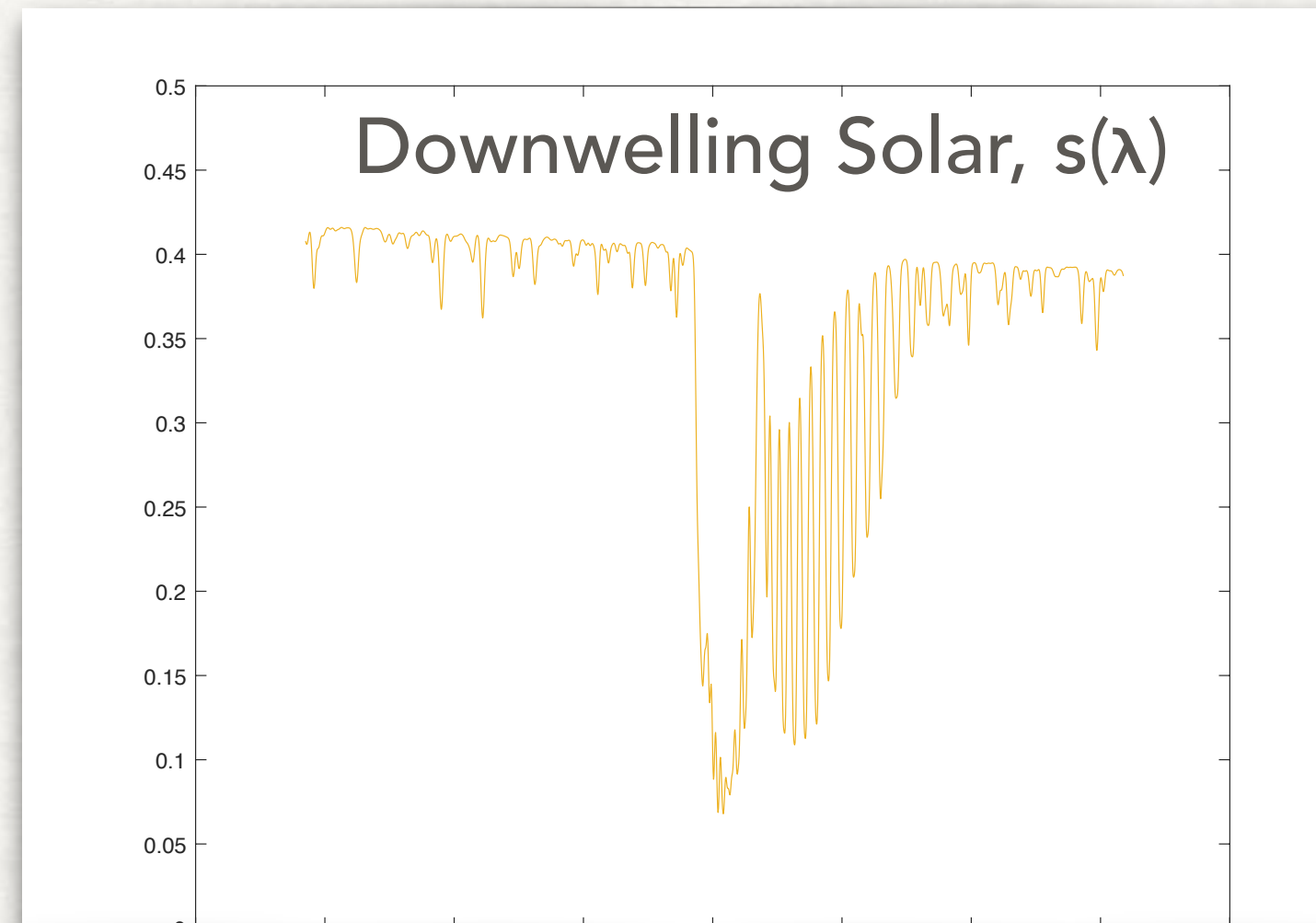
Separation of Reflected and Fluorescence Light

Forward Model

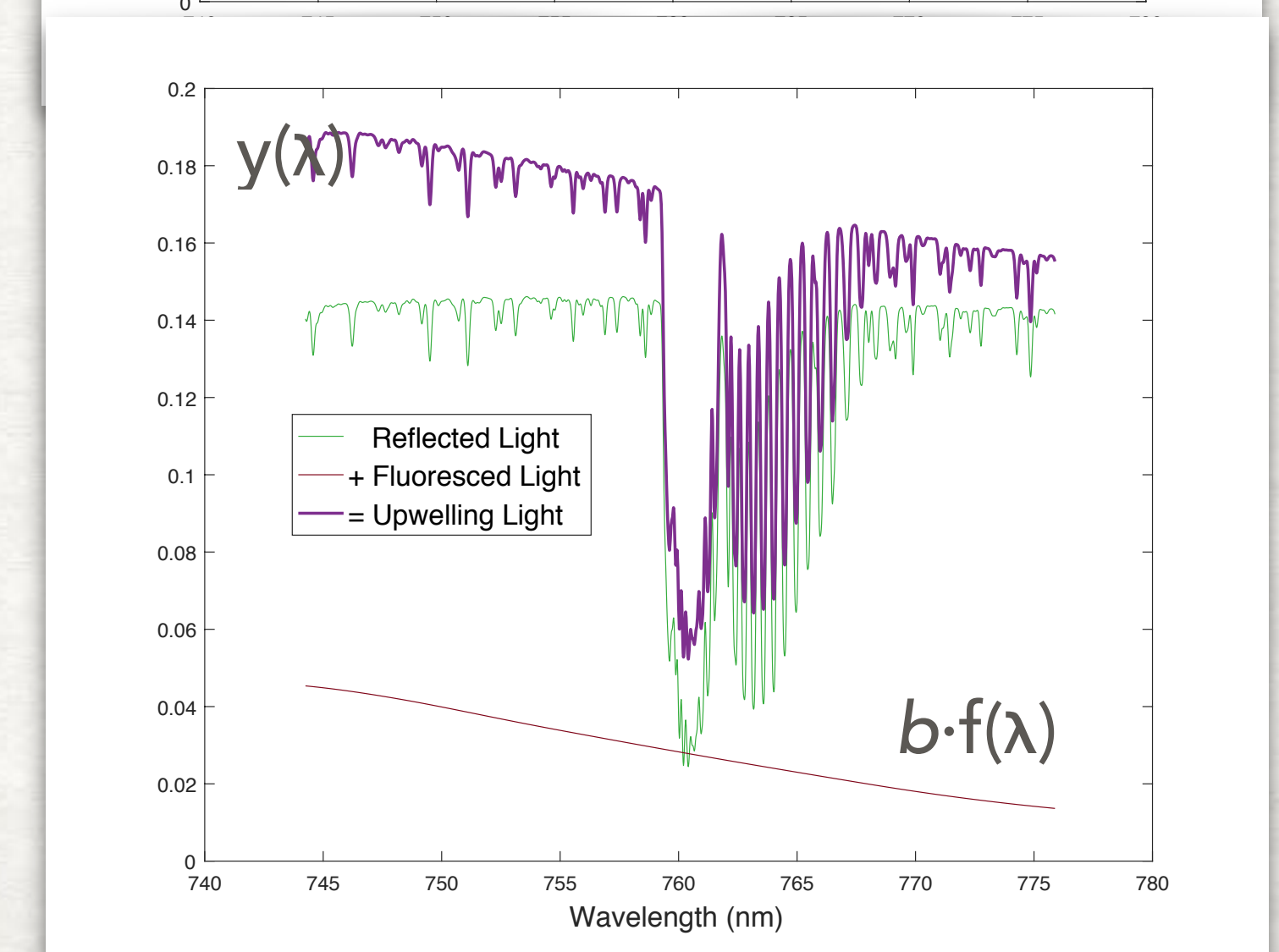
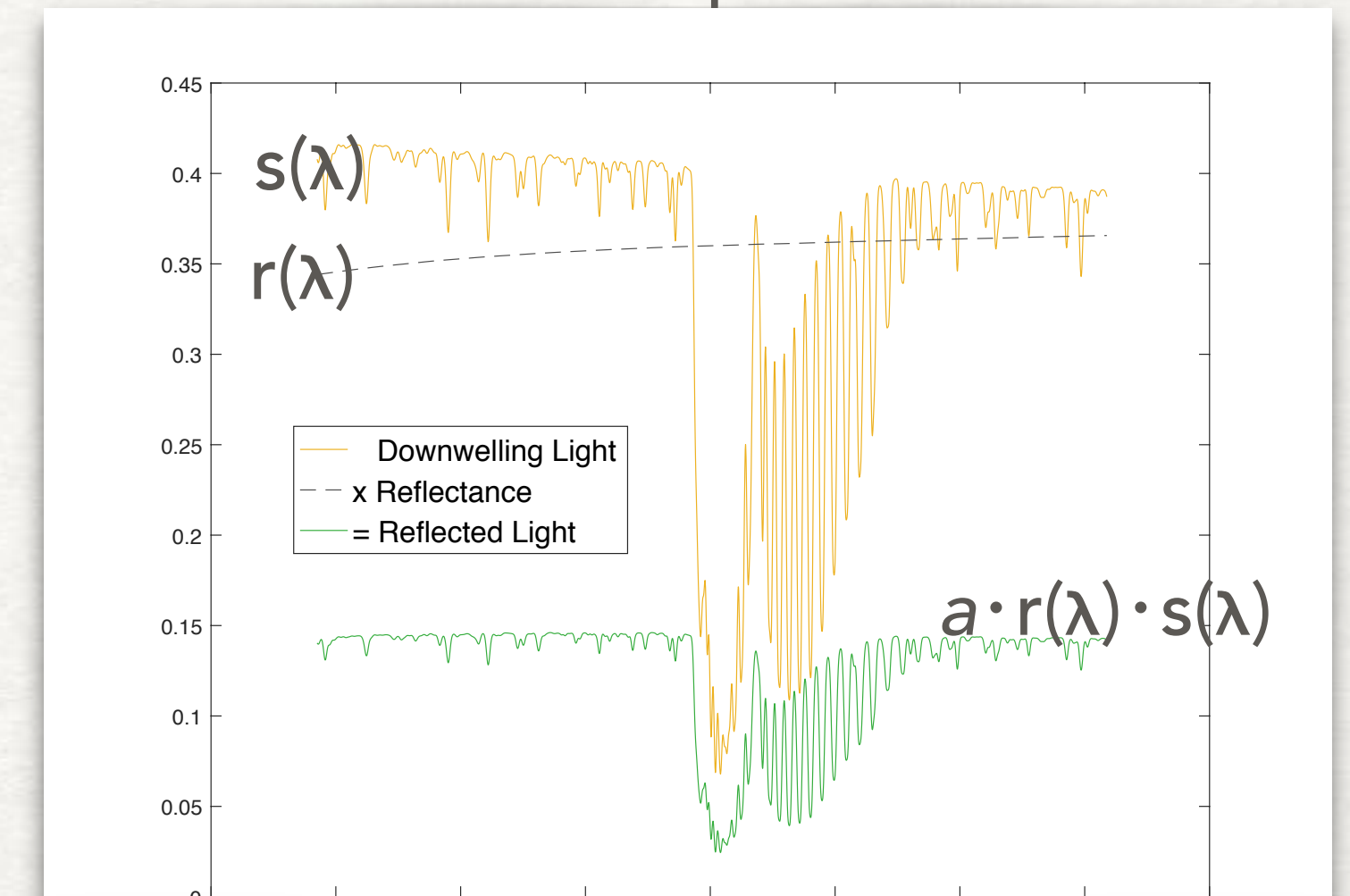
(a) Reflected light
 \downarrow

$$y(\lambda) = a \cdot r(\lambda) \cdot s(\lambda) + b \cdot f(\lambda)$$
 \uparrow
 (b) Fluorescence light

Measured



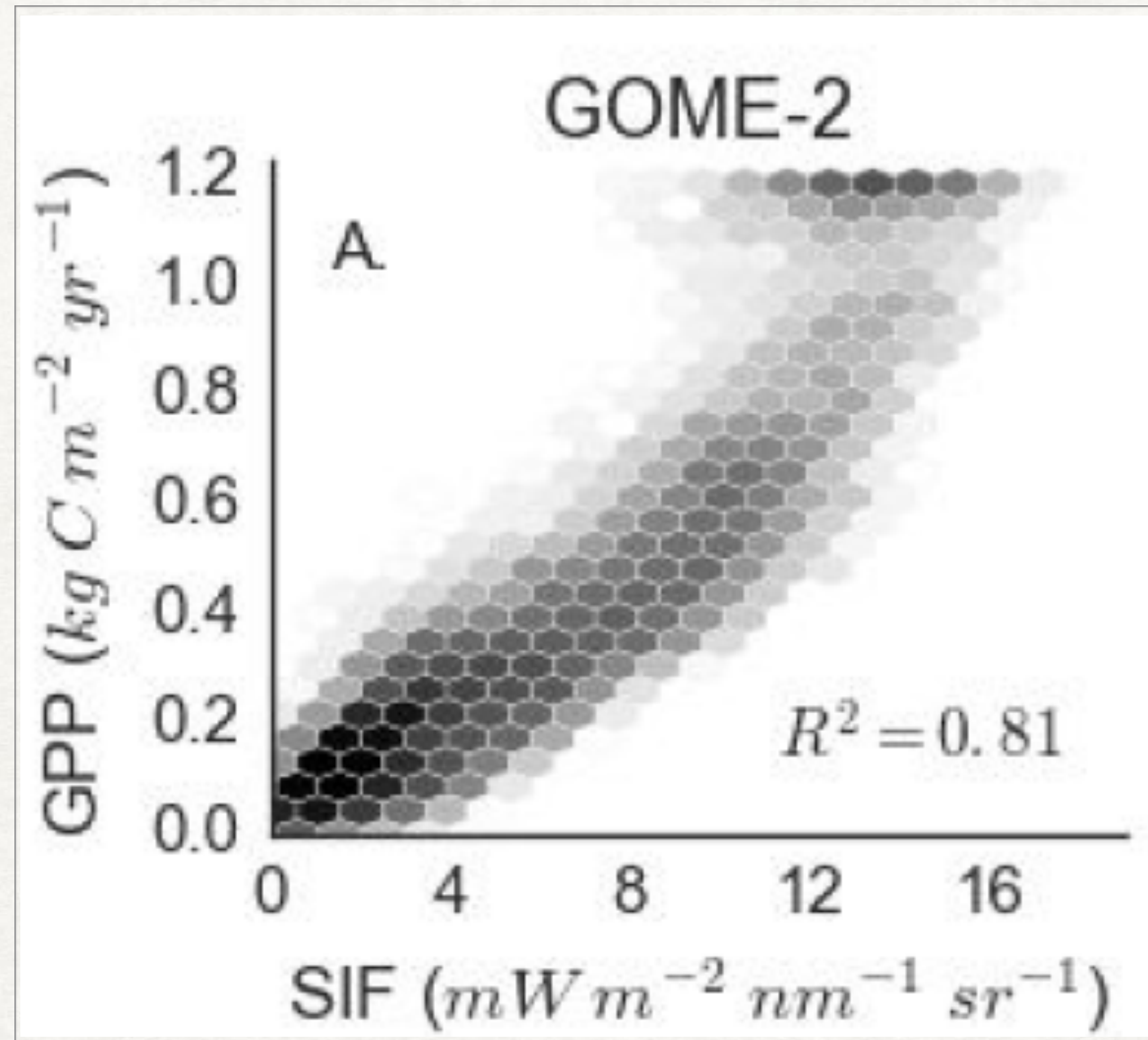
Components





SIF can be used to Estimate GPP

Monthly MPI-GPP at 0.5° vs SIF (GOME-2)



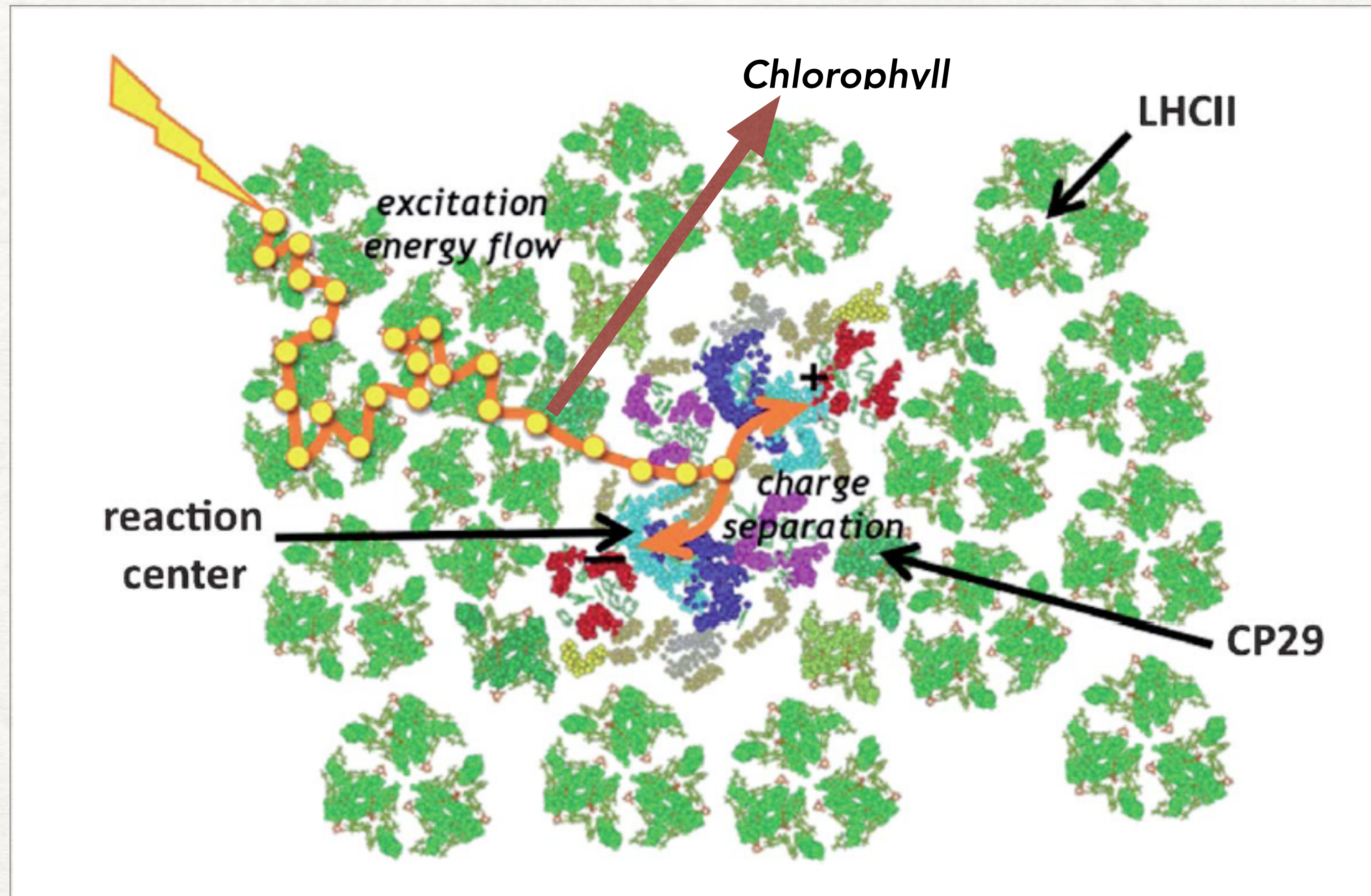
SIF:

- A new method based on resolving fine spectral details permits us to resolve the light emitted by chlorophyll.
- It seems to be a powerful proxy for photosynthesis.
- We want to know the basis of this:

Is it fluorescence yield

Or, is it biophysics?

SIF Results from the Decay of an Excited Chlorophyll Molecule



The yield of fluorescence (τ_F), is related to the lifetime of the excited state (τ_P) and the intrinsic lifetime of chlorophyll (τ_{Chl}).

$$\Phi_F = \frac{\tau_P}{\tau_{Chl}} \approx 0.01$$

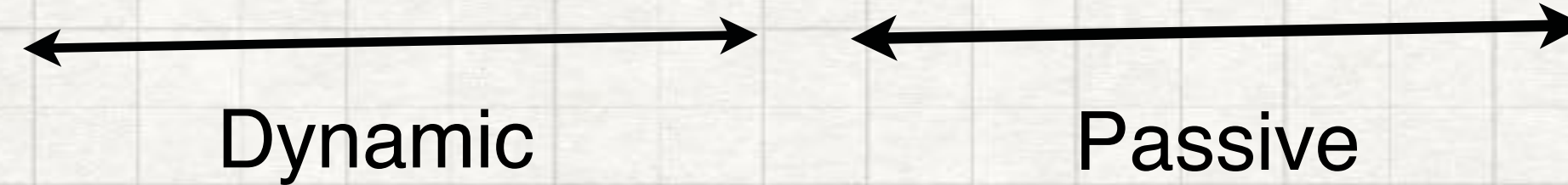
The Fundamental Conservation Equation

$$\Phi_P + \Phi_N + \Phi_D + \Phi_F = 1$$

Loss + PSI

Yield of fluorescence ~ 0.01 to 0.05

Directly proportional to life time, τ of the excited state



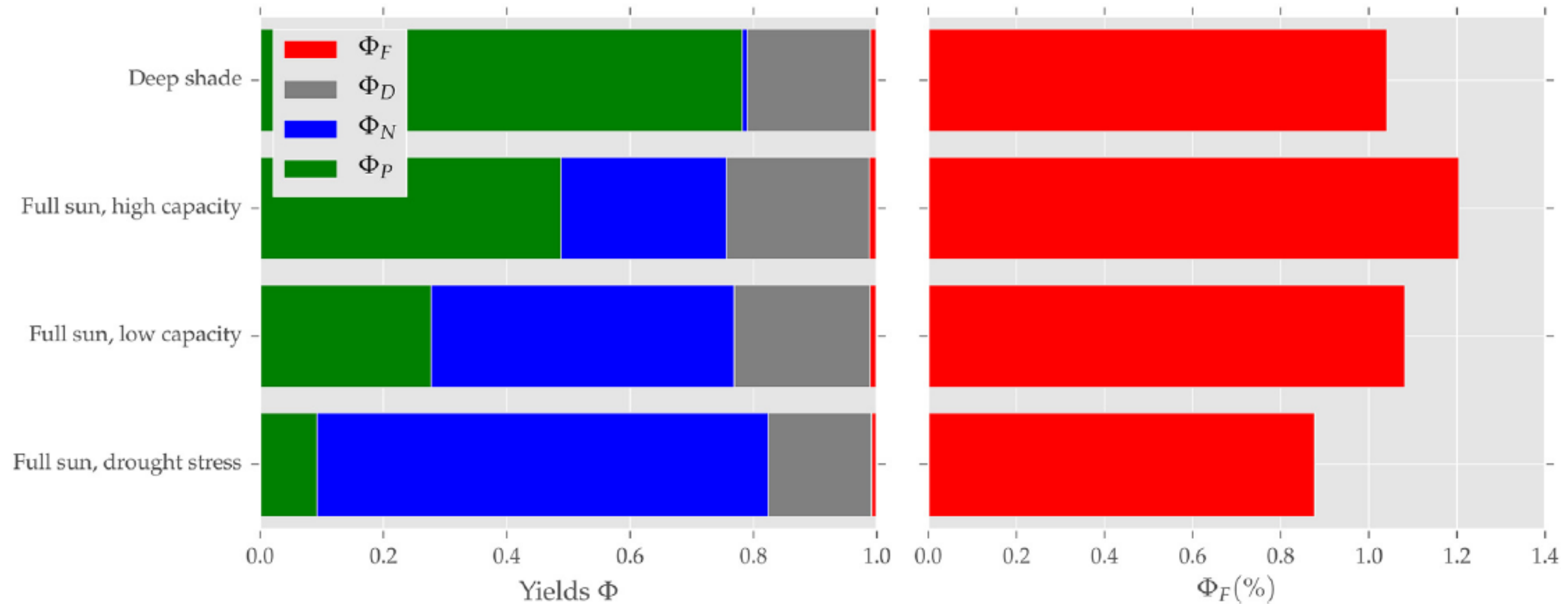
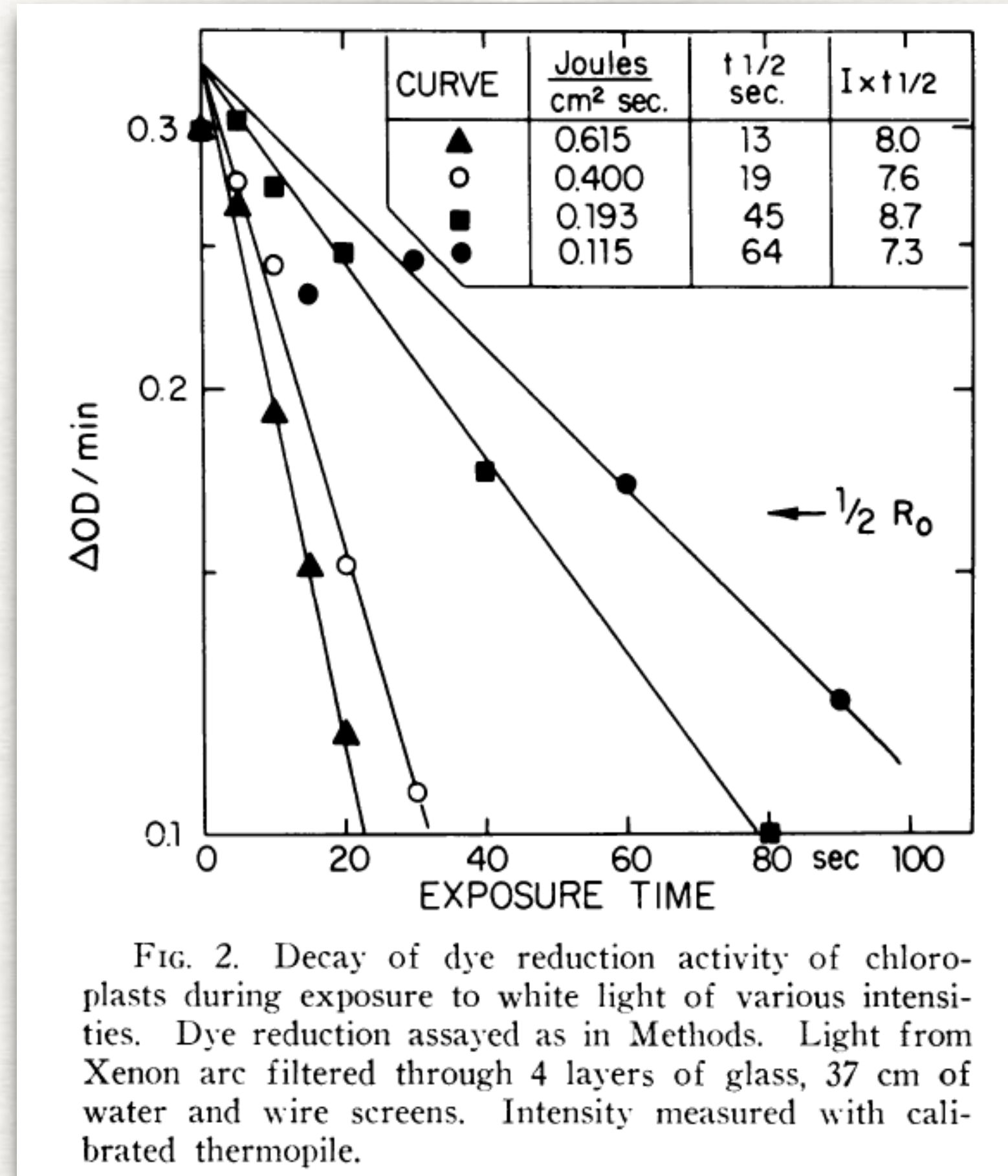


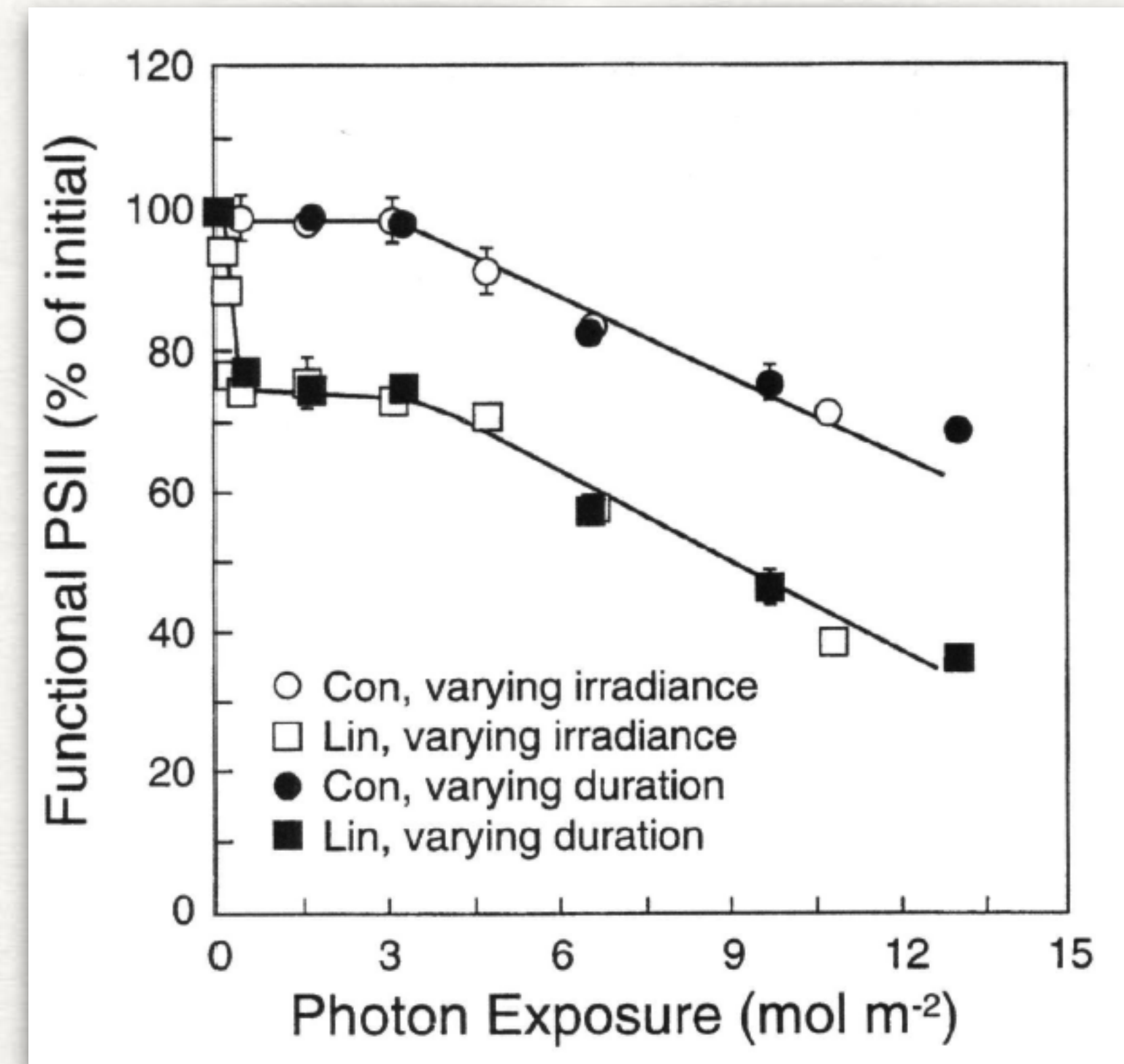
Fig. 2 Quantum yields (Φ) for the four pathways used by leaves to process photons depend on whether the leaf is shaded or exposed to full sun, whether the leaf has a high or low photosynthetic capacity, and whether drought stress is present. In the histogram, *green* denotes photosynthesis (PSII yield); *blue*, nonphotochemical quenching; *gray*, nonradiative decay; and *red*, fluorescence. The results shown here were obtained with a laboratory instrument called a PAM fluorometer. As the *right panel* shows, the fluorescence quantum yield changes less than the PSII yield but also noticeably with varying conditions, especially once nonphotochemical quenching becomes dominant.

Photoinhibition is a first order photochemical process



Jones, L. W., & Kok, B. (1966). Photoinhibition of Chloroplast Reactions. I. Kinetics and Action Spectra. *Plant Physiology*, 41(6), 1037–1043. <http://doi.org/10.1104/pp.41.6.1037>

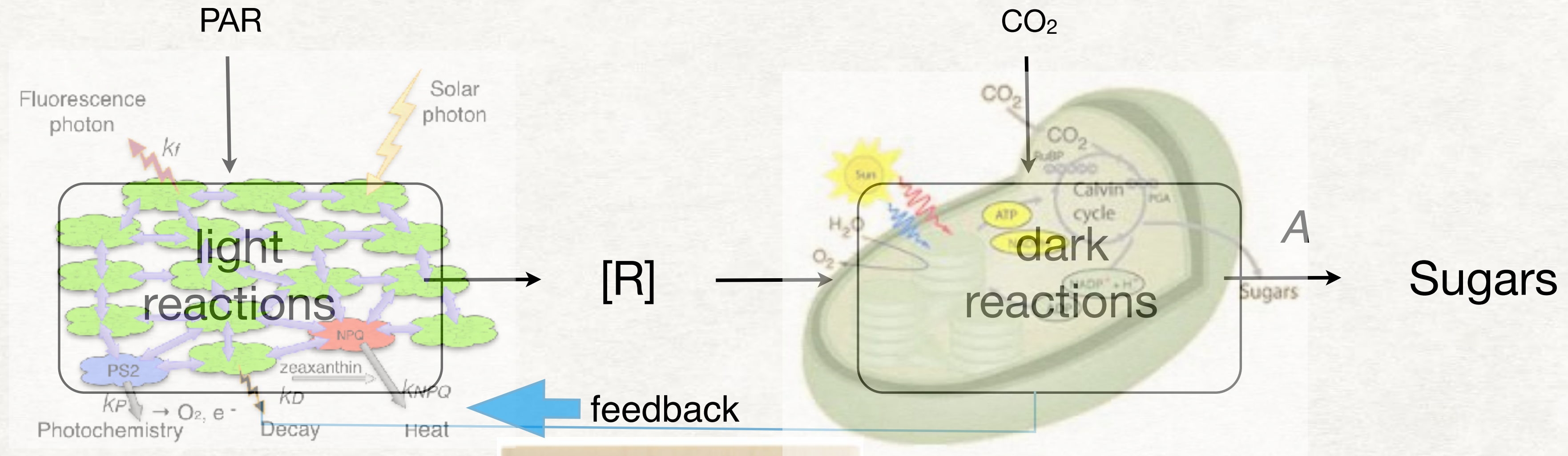
Anderson, J. M., Park, Y. I., & Chow, W. S. (1997). Photoinactivation and photoprotection of photosystem II in nature. *Physiologia Plantarum*, 100(2), 214–223. <http://doi.org/10.1111/j.1399-3054.1997.tb04777.x>



$$\Phi_i = \frac{K_i}{K_P + K_N + K_D + K_F + K_i}$$

$$K_i \approx 1.5 \times 10^{-6}$$

Interactions between the light and dark reactions:



$$A \approx \min \begin{cases} \text{light reactions } (W_L) \\ \text{dark reactions } (W_C) \end{cases}$$

PAM Fluorometry - the key for understanding SIF

$$F \propto \Phi_F,$$

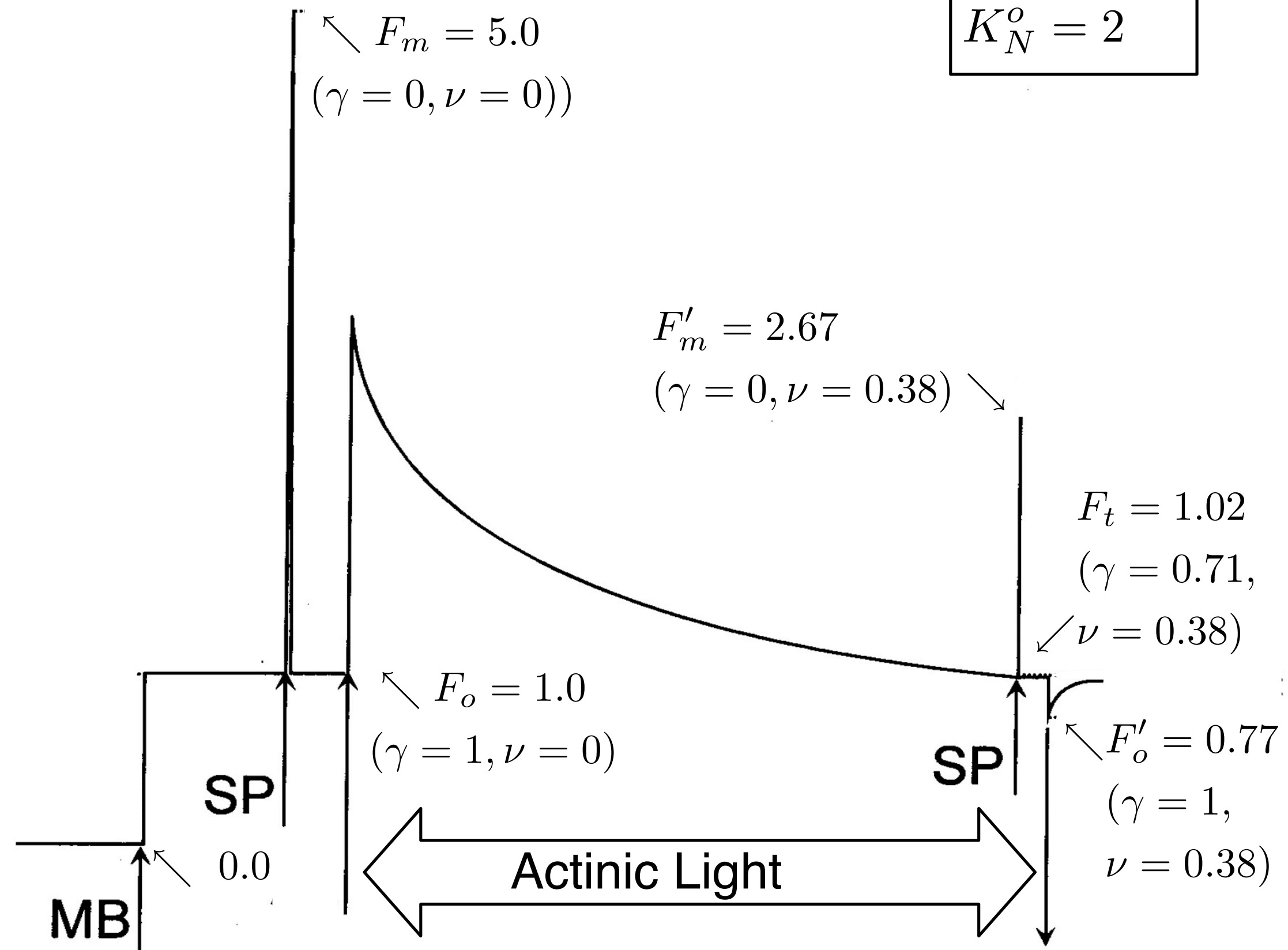
$$\Phi_F = \frac{K_F}{K_F + K_D + \gamma K_P^o + \nu K_N^o}$$

$$K_F = 0.05$$

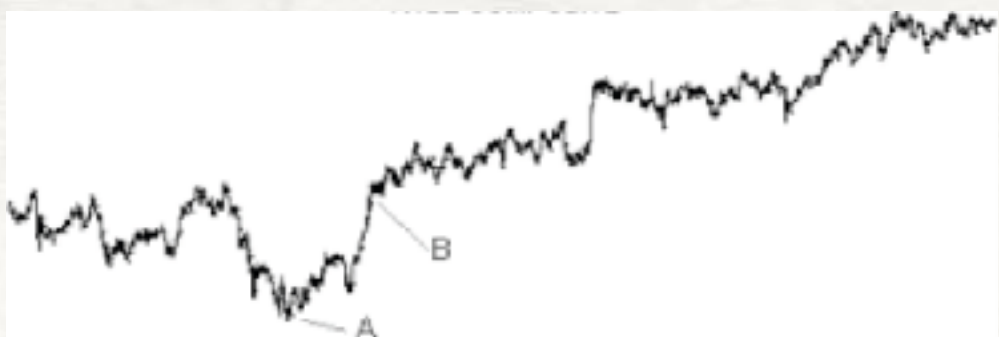
$$K_D = 0.95$$

$$K_P^o = 4$$

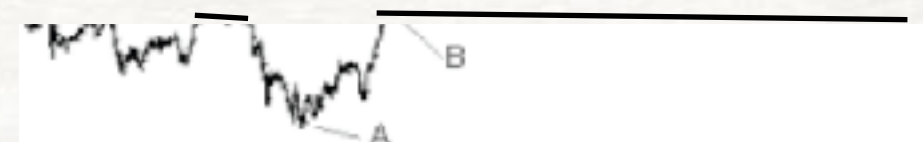
$$K_N^o = 2$$



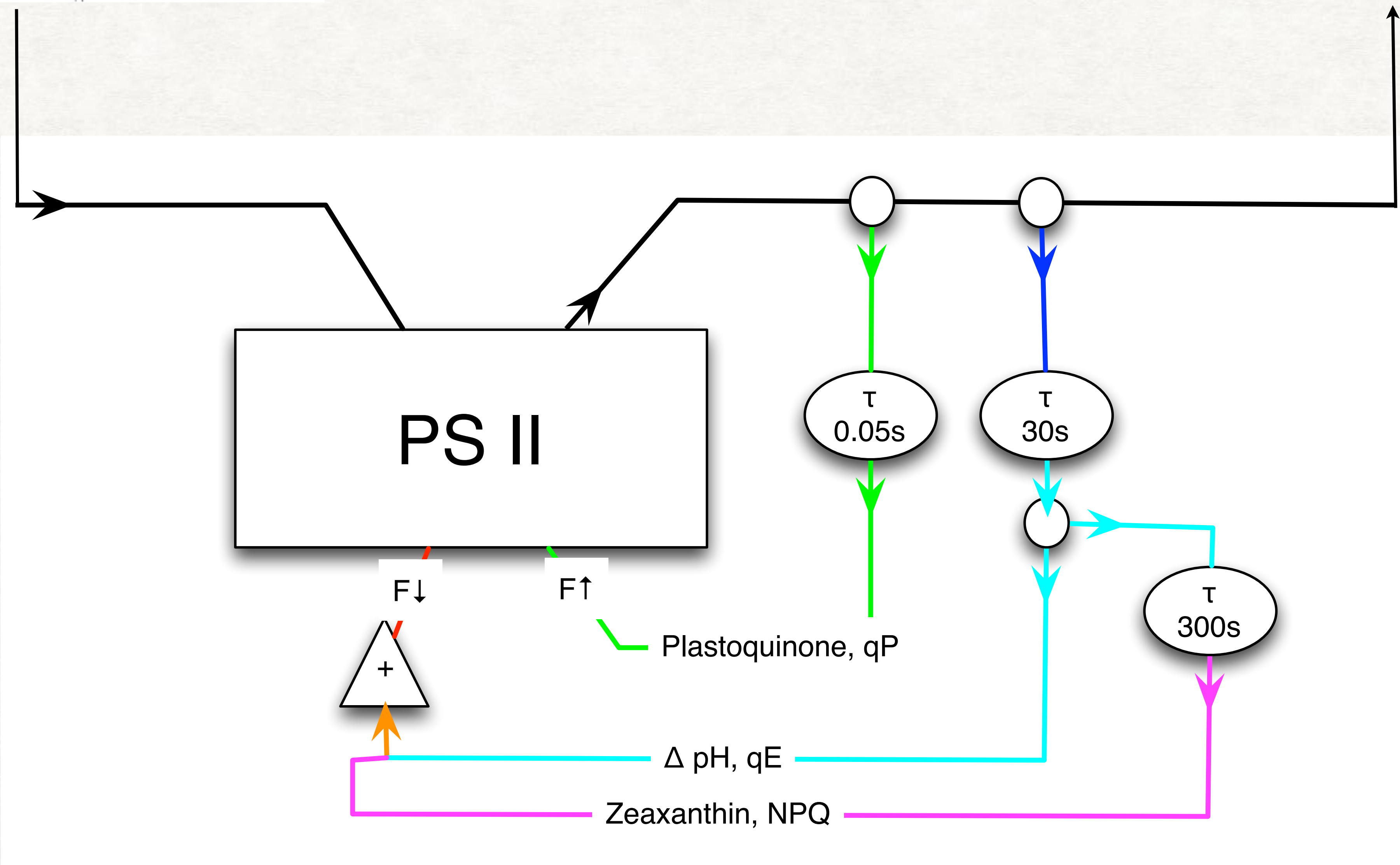
Input Light



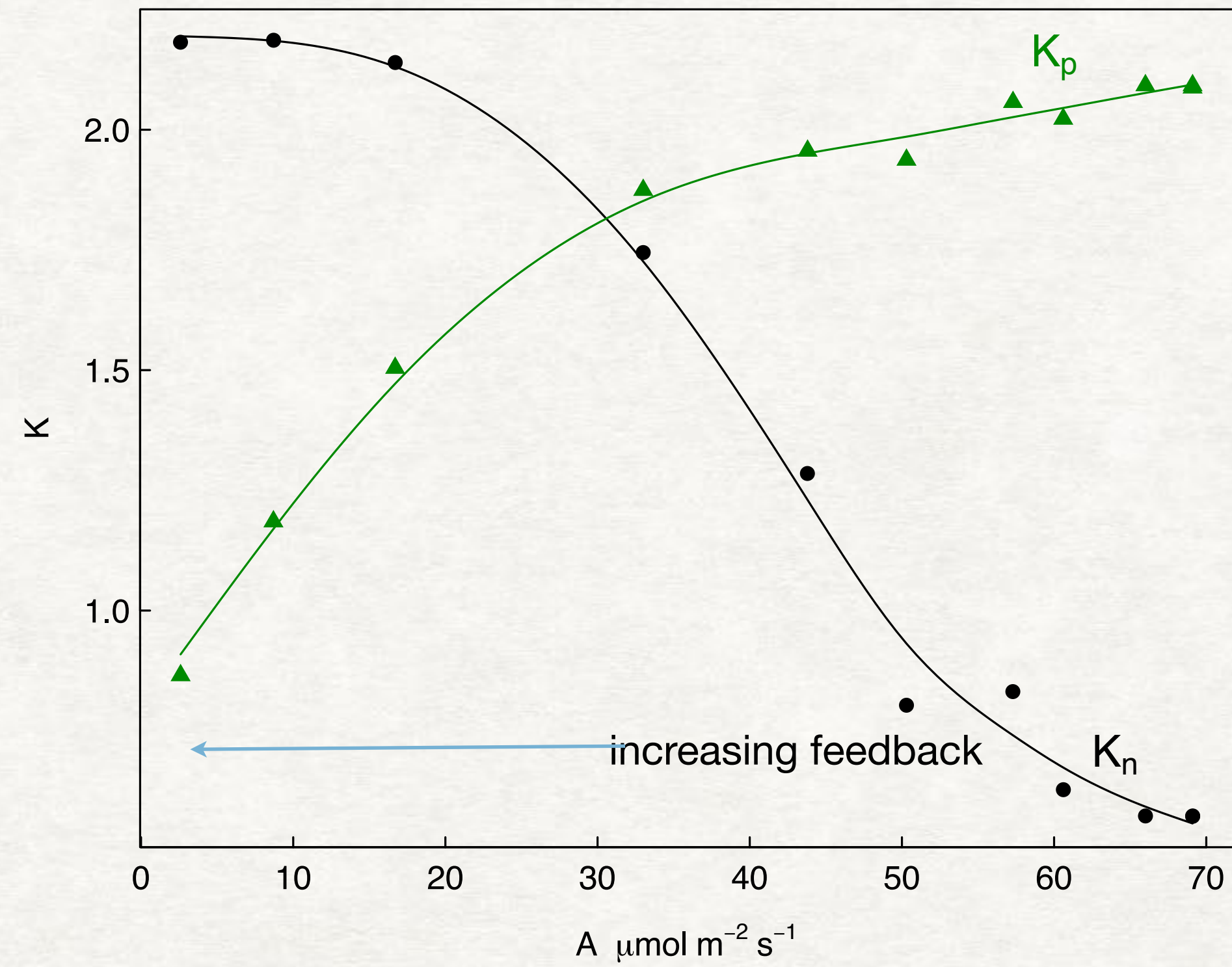
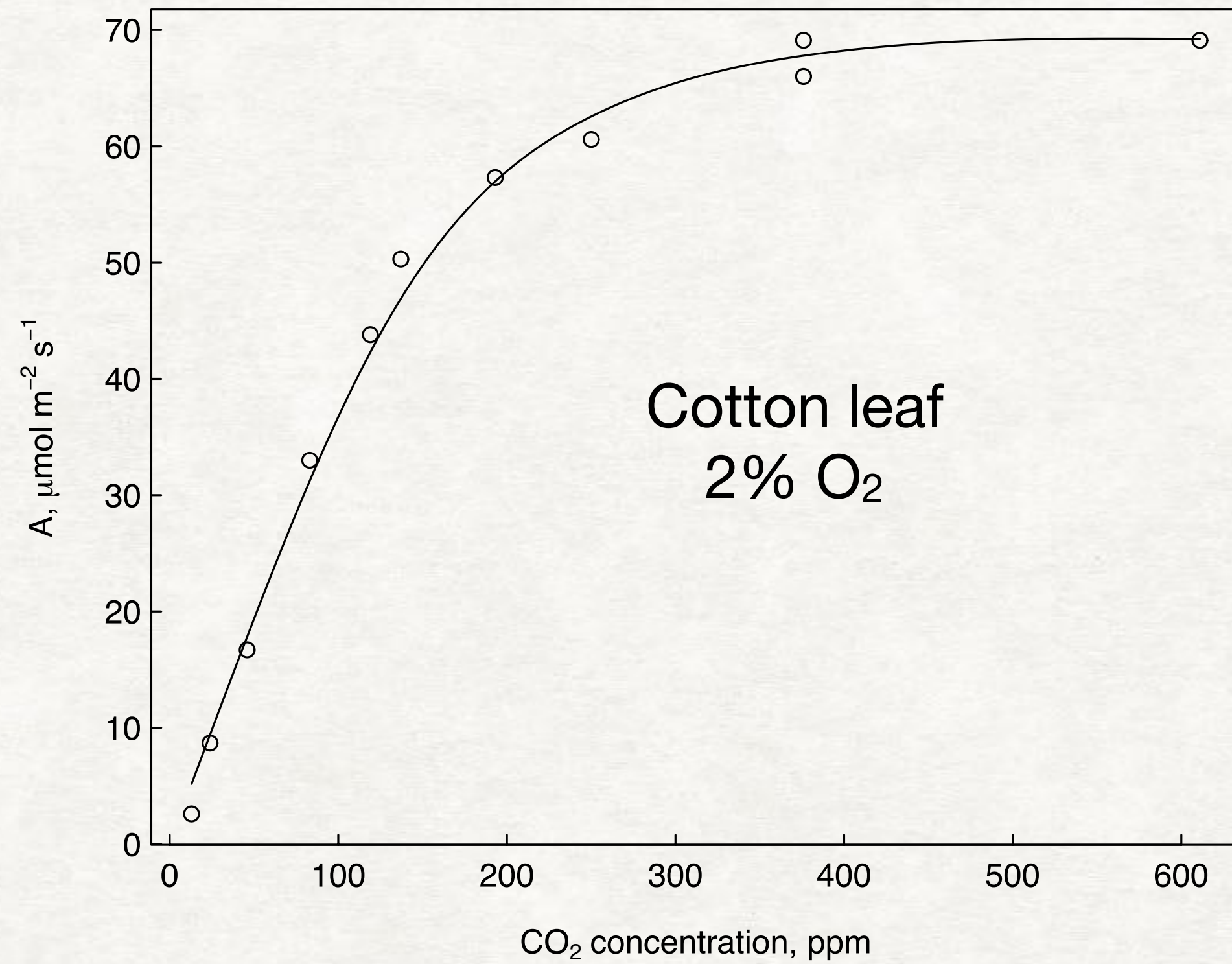
Output to carbon fixation



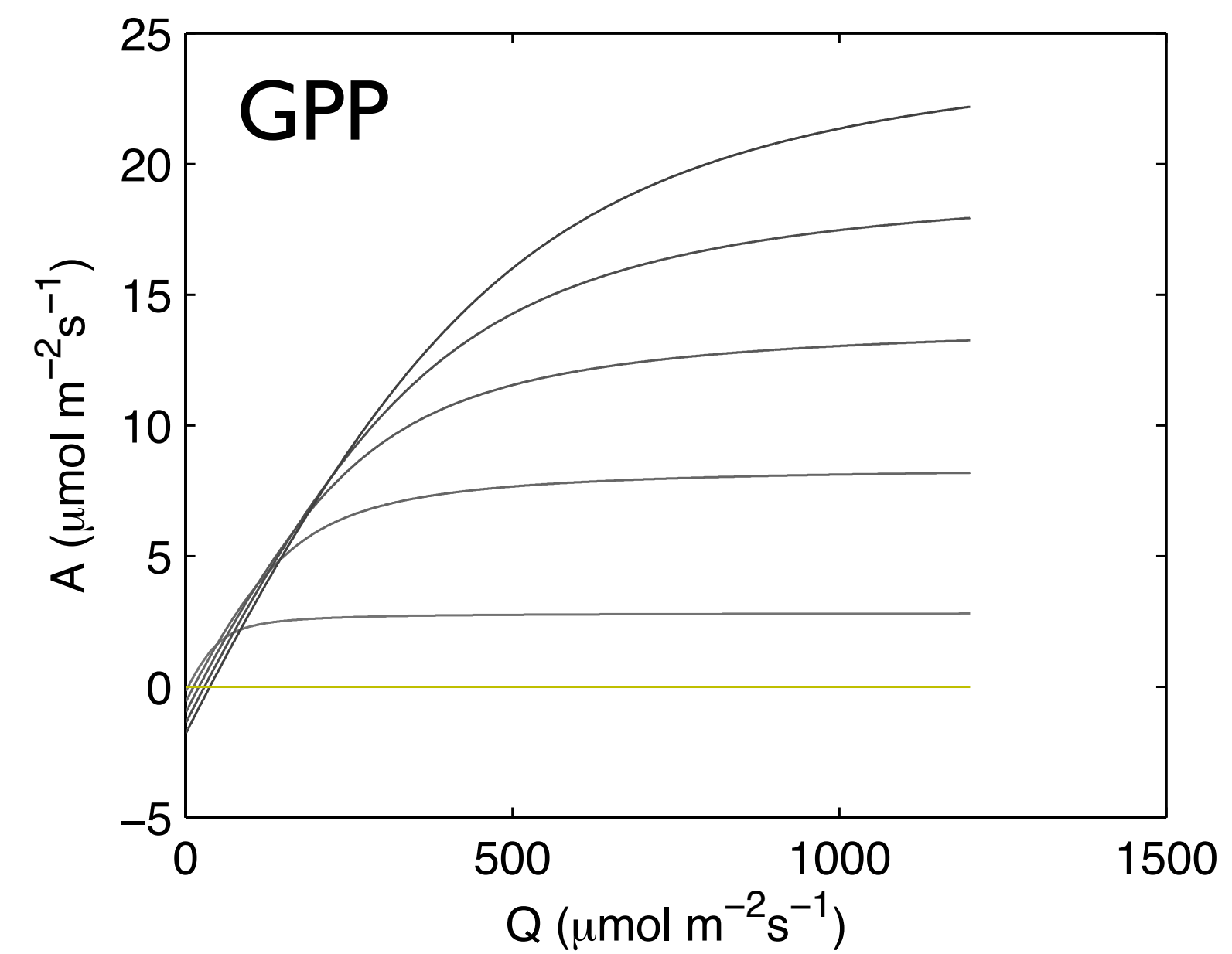
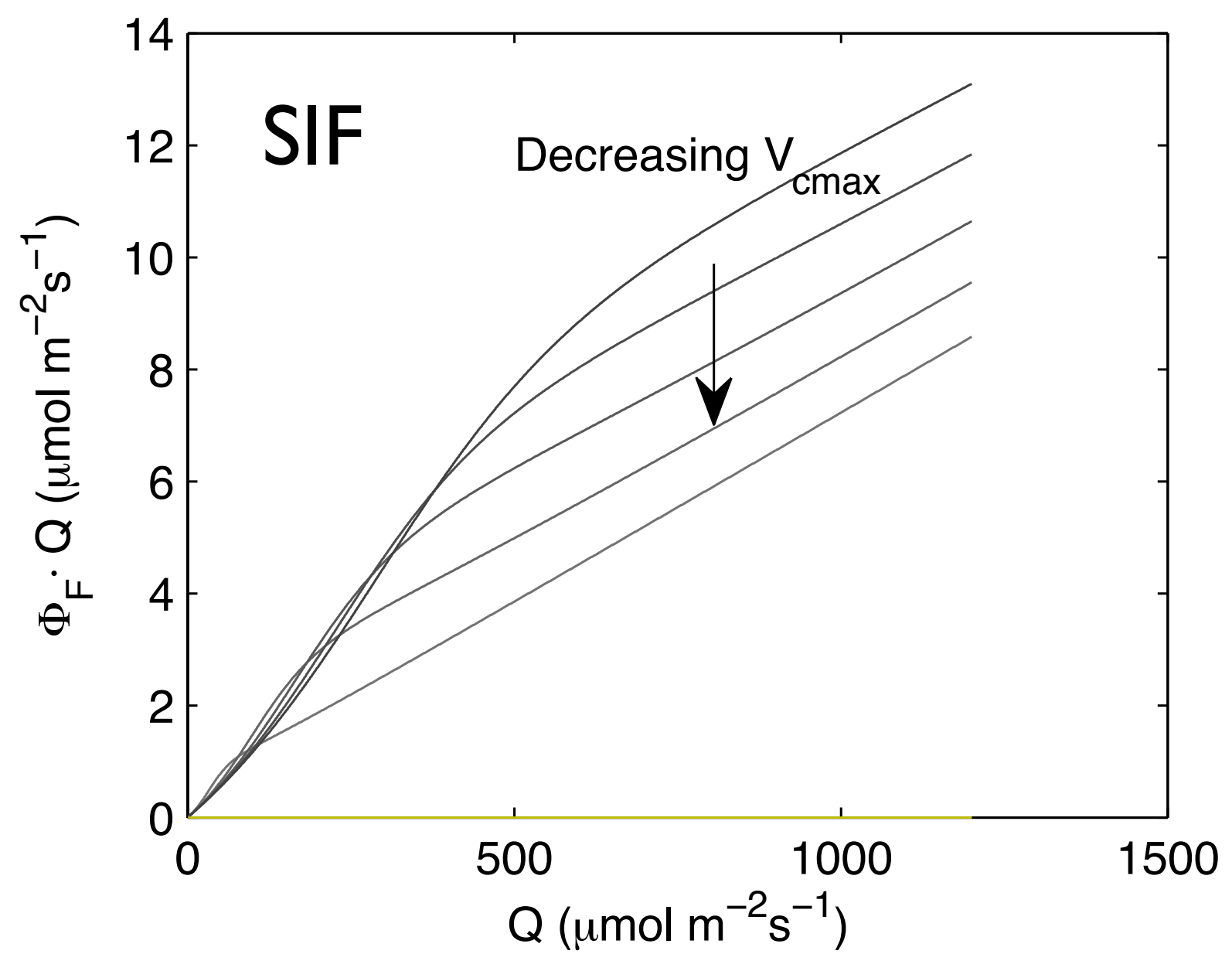
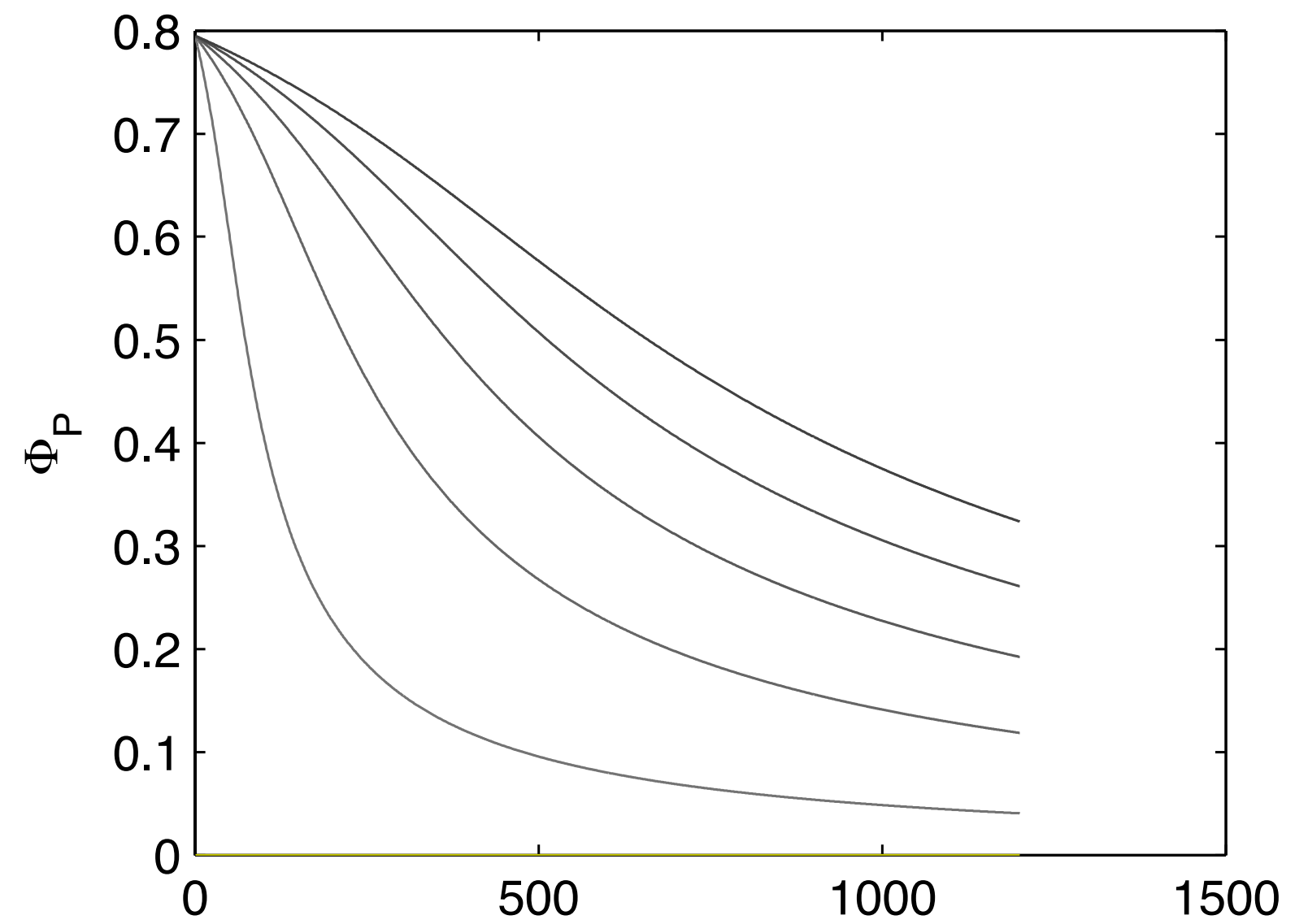
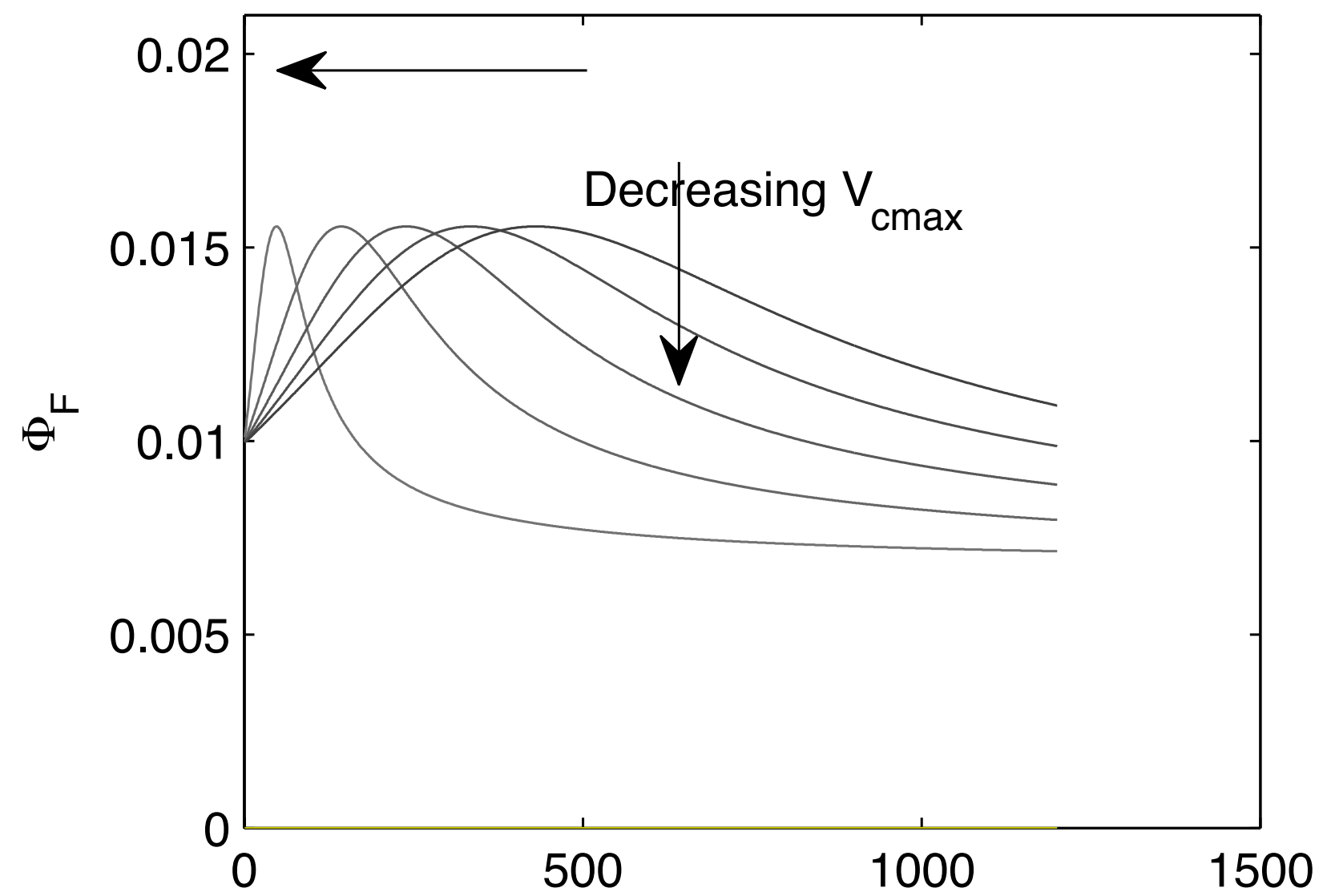
Think of Photosystem 2 as a sophisticated IC.



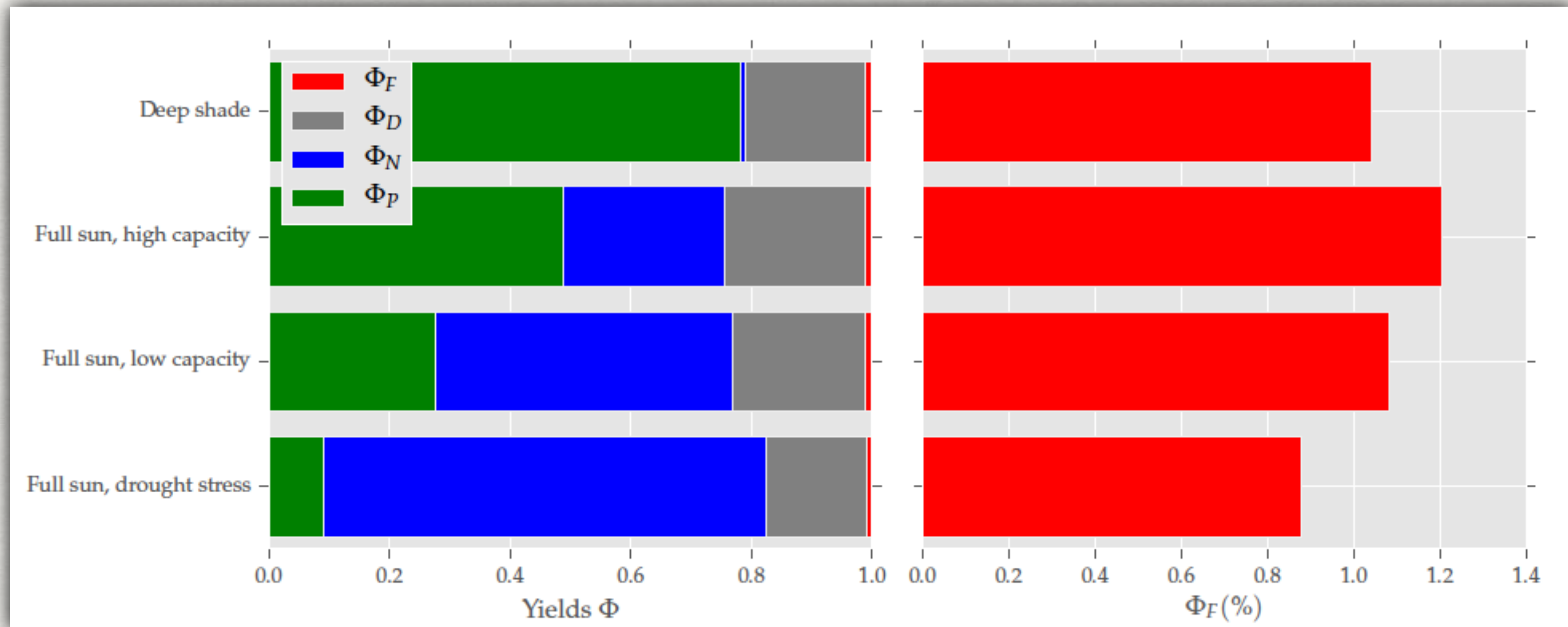
The Pattern of Feedback



$$\Phi_P = \frac{K_P}{K_P + K_N + K_D + K_F}$$

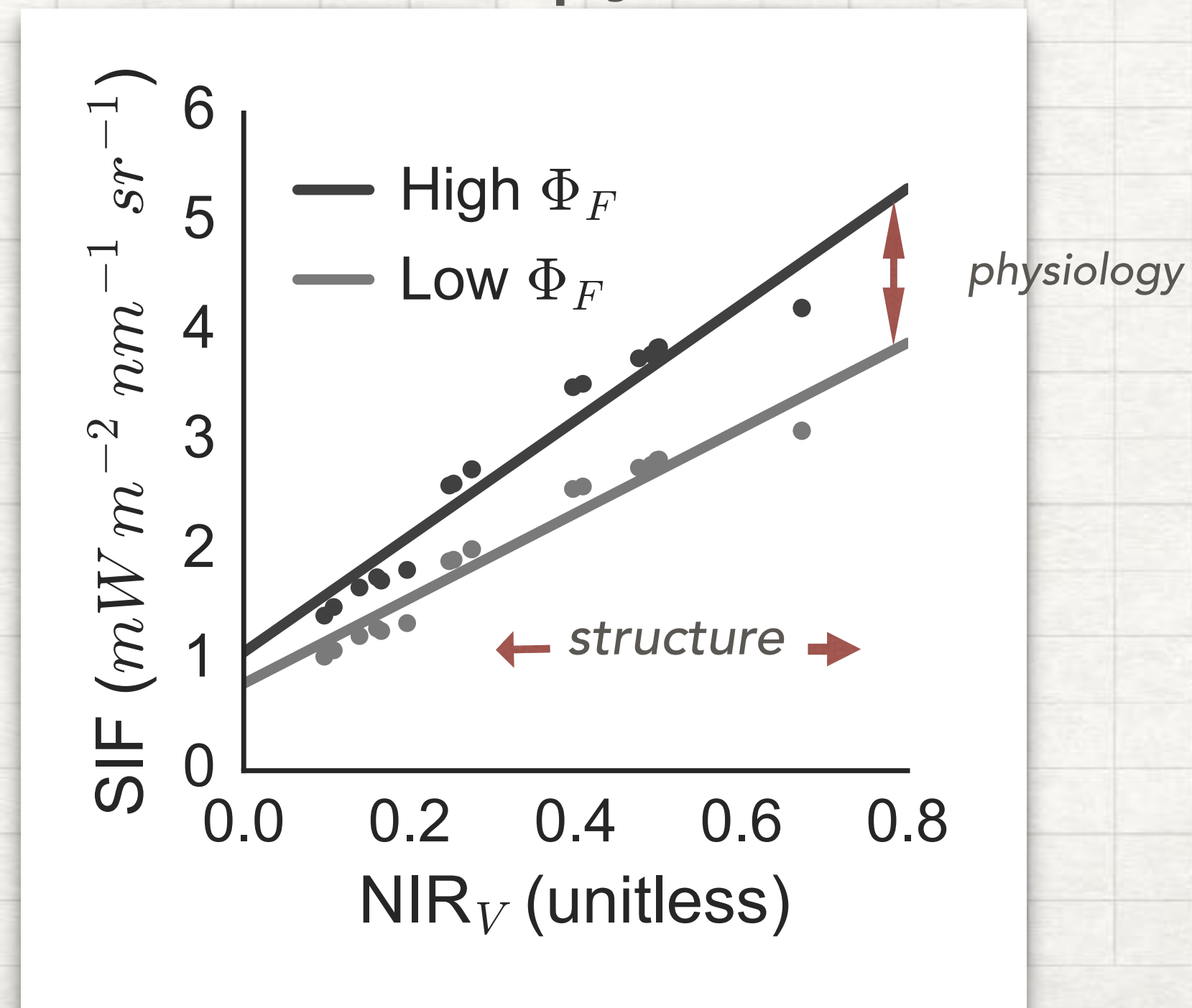


The Yield of SIF (Φ_F) is linked to Photochemistry (Φ_P) and Non-Photochemical Quenching (Φ_N)

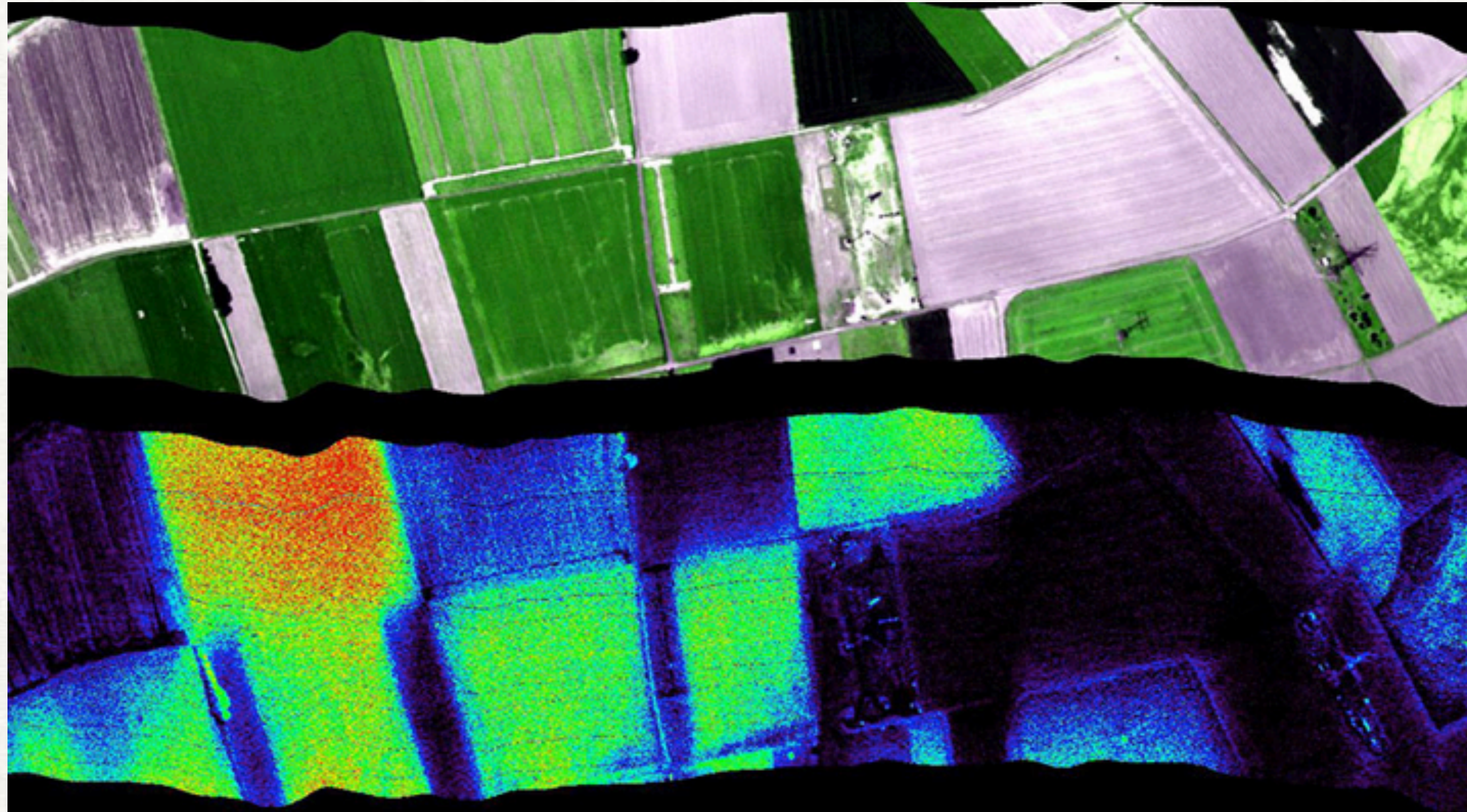


Summing up:

- We see that physiology works in the right direction, but it probably isn't sufficient to account for observed variation in SIF.
- Could it be the canopy structure?



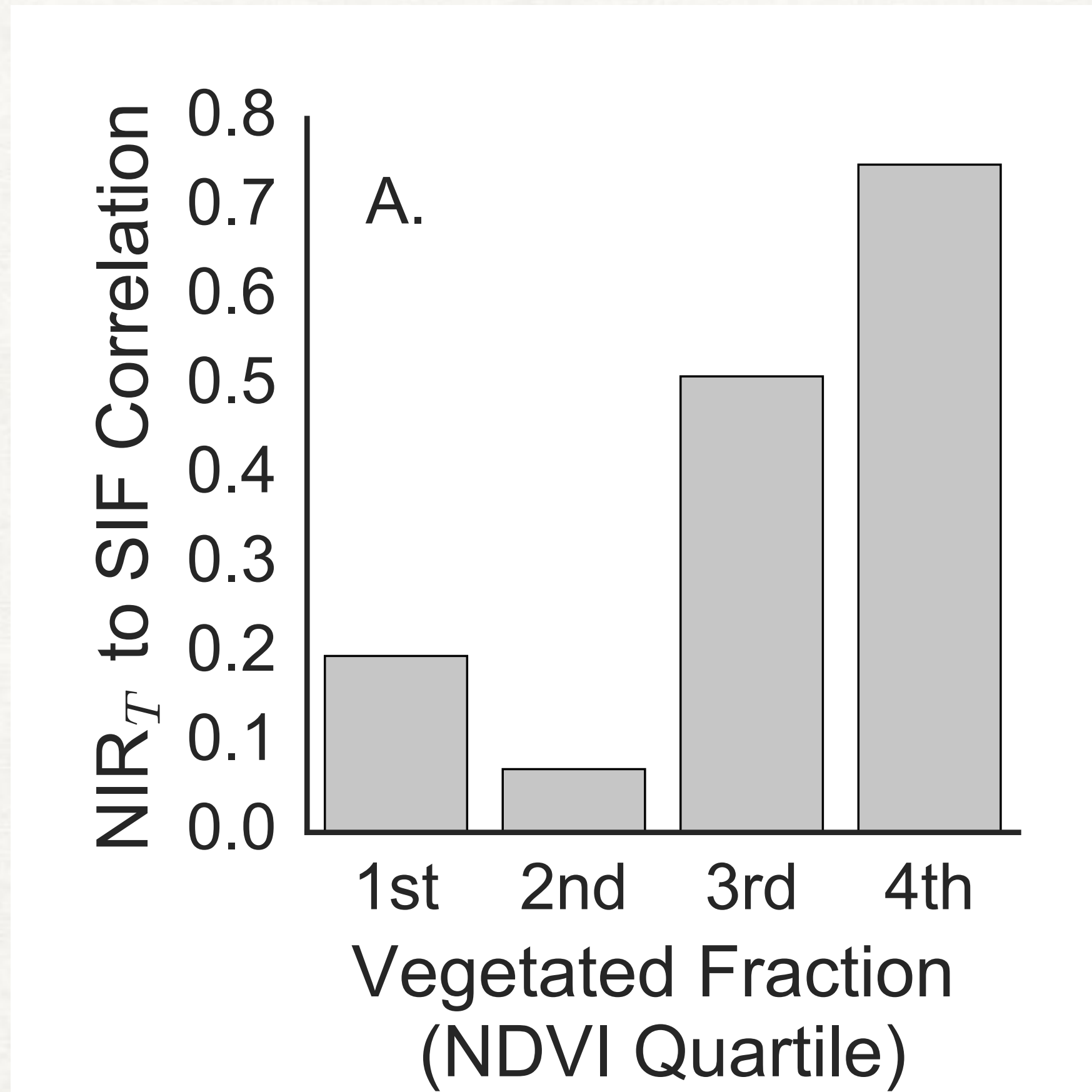
- SIF (bottom) is specific to vegetation, while reflectance (top) sees all surfaces of the scene.



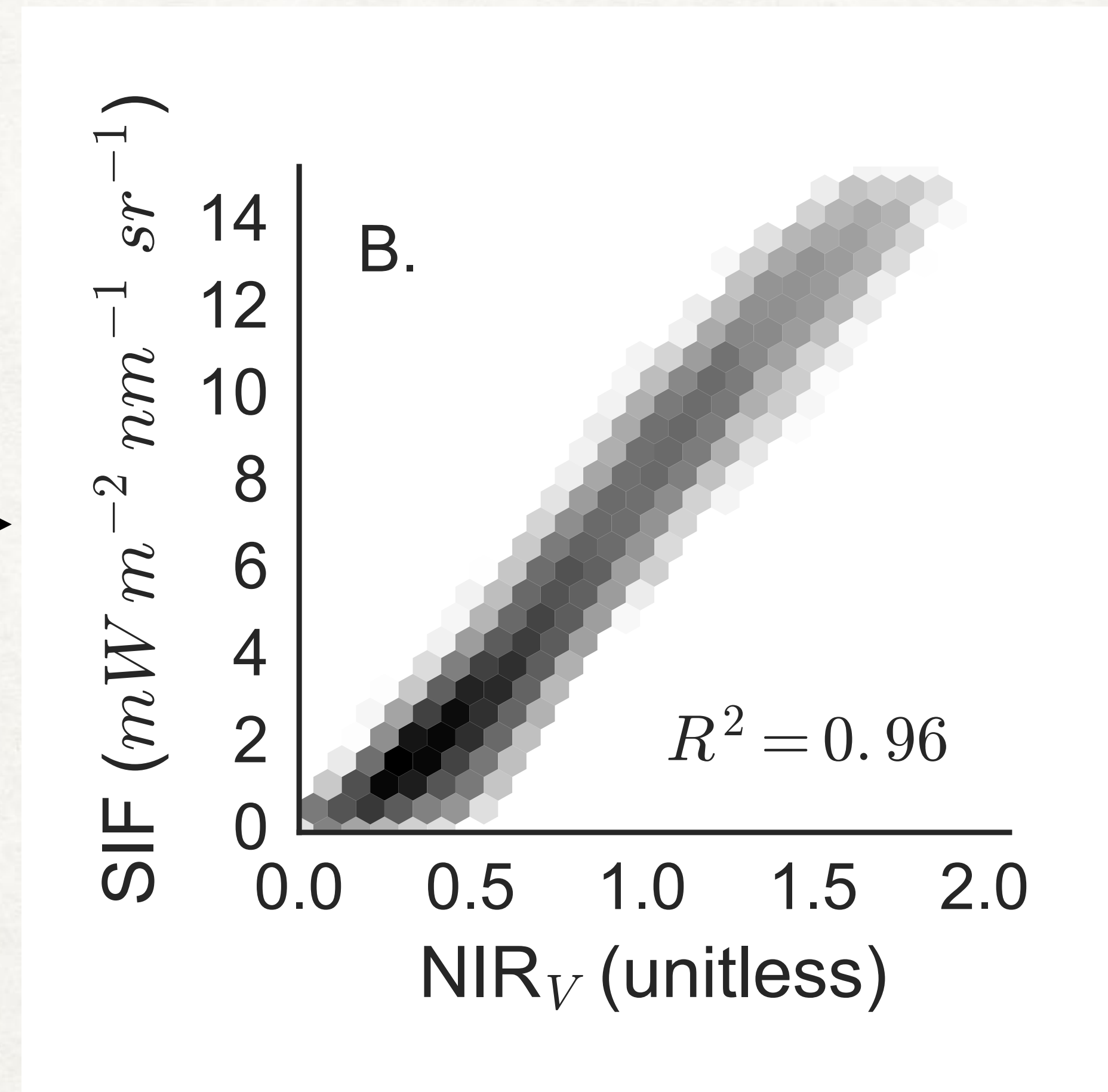
- SIF seems to be proportional to the NIR reflectance of the canopy (excluding the soil).



NDVI x NIR Radiance = NIR_V (vegetation)?



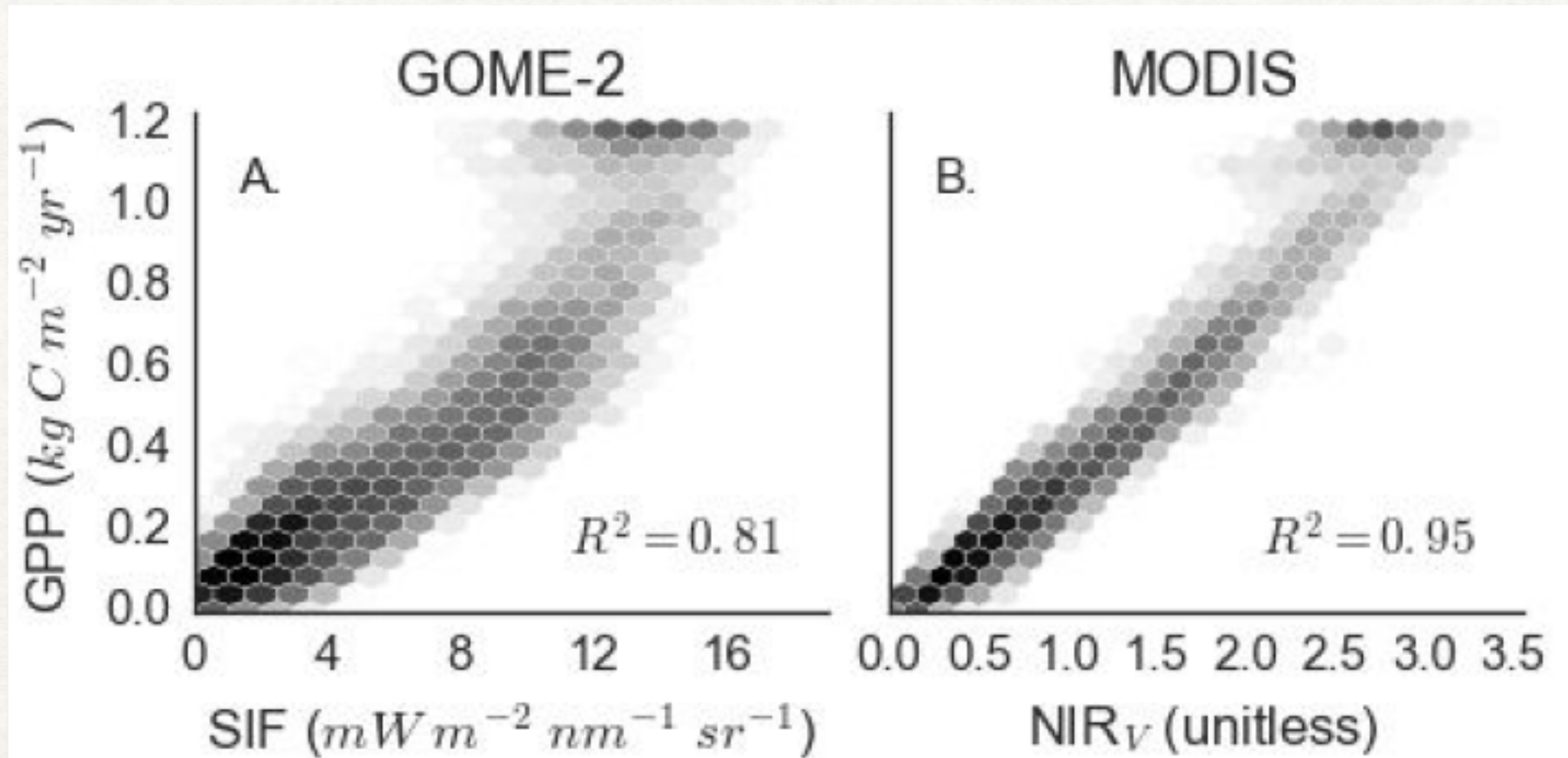
→
NIR_T * NDVI





CAN WE USE NIR_V TO ESTIMATE GPP?

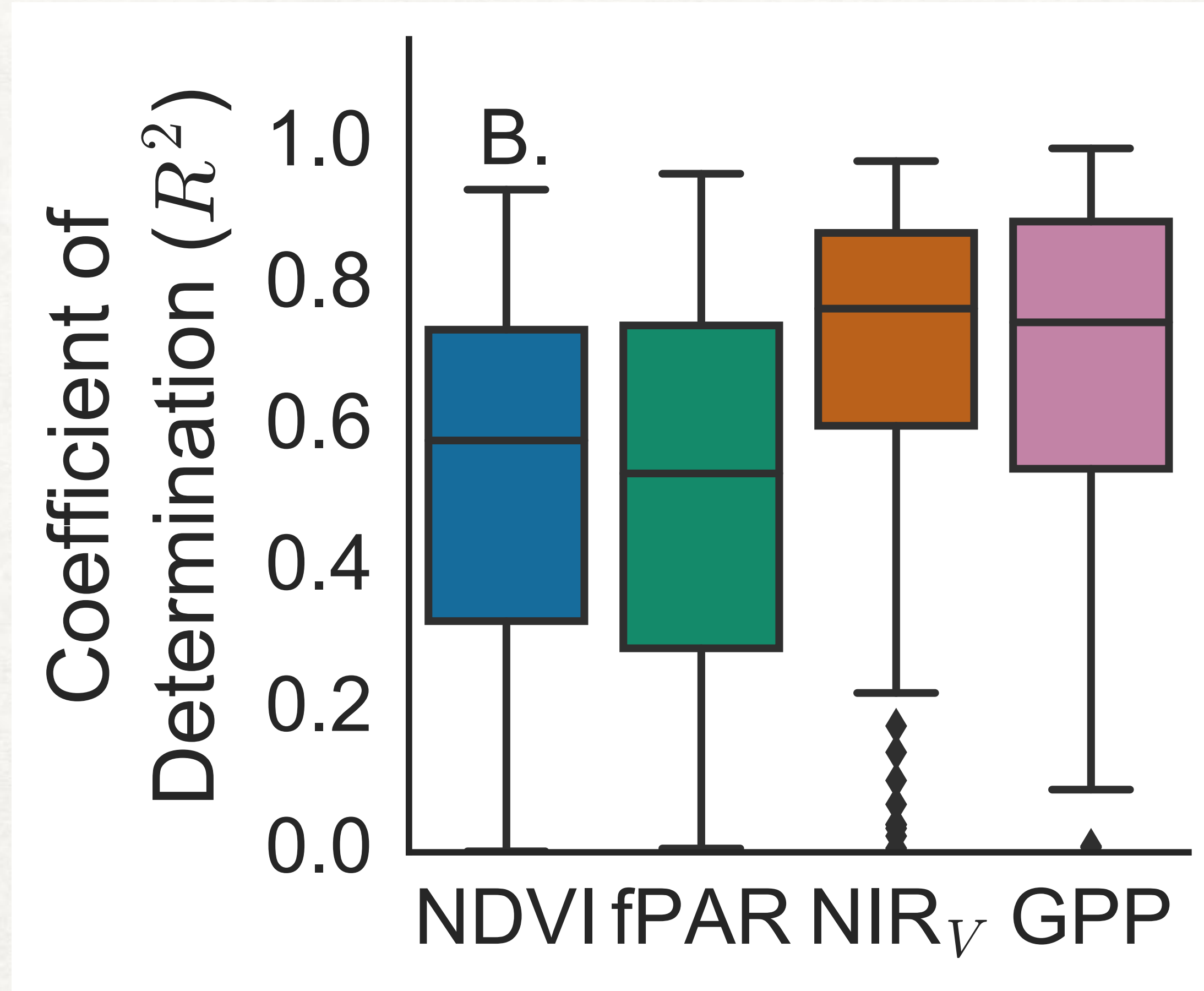
Monthly MPI-GPP at 0.5° vs SIF (GOME-2) or NIR_V (MODIS)



Badgley, G., Field, C. B., & Berry, J. A. (2017). Canopy near-infrared reflectance and terrestrial photosynthesis. *Science Advances*, 3(3), e1602244. <http://doi.org/10.1126/sciadv.1602244>



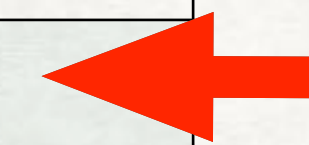
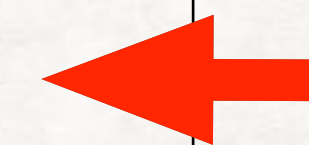
MODIS (0.5 km) used to evaluate NIR_V at 106 flux sites.





Values for the month of peak productivity

Land Cover Class	NIR _v (unitless)	fPAR (unitless)	GPP (gC m ⁻² d ⁻¹)
Deciduous Broadleaf	0.28	0.72	11.18
Cropland	0.27	0.62	12.57
Mixed Forest	0.22	0.68	9.35
Grassland	0.19	0.55	6.61
Wetland	0.19	0.58	6.41
Evergreen Broadleaf	0.17	0.70	7.87
Evergreen Needleleaf	0.17	0.67	7.85
Savanna	0.15	0.45	5.62
Woody Savanna	0.14	0.42	6.23
Open Shrubland	0.13	0.48	3.85
Closed Shrubland	0.10	0.46	5.30



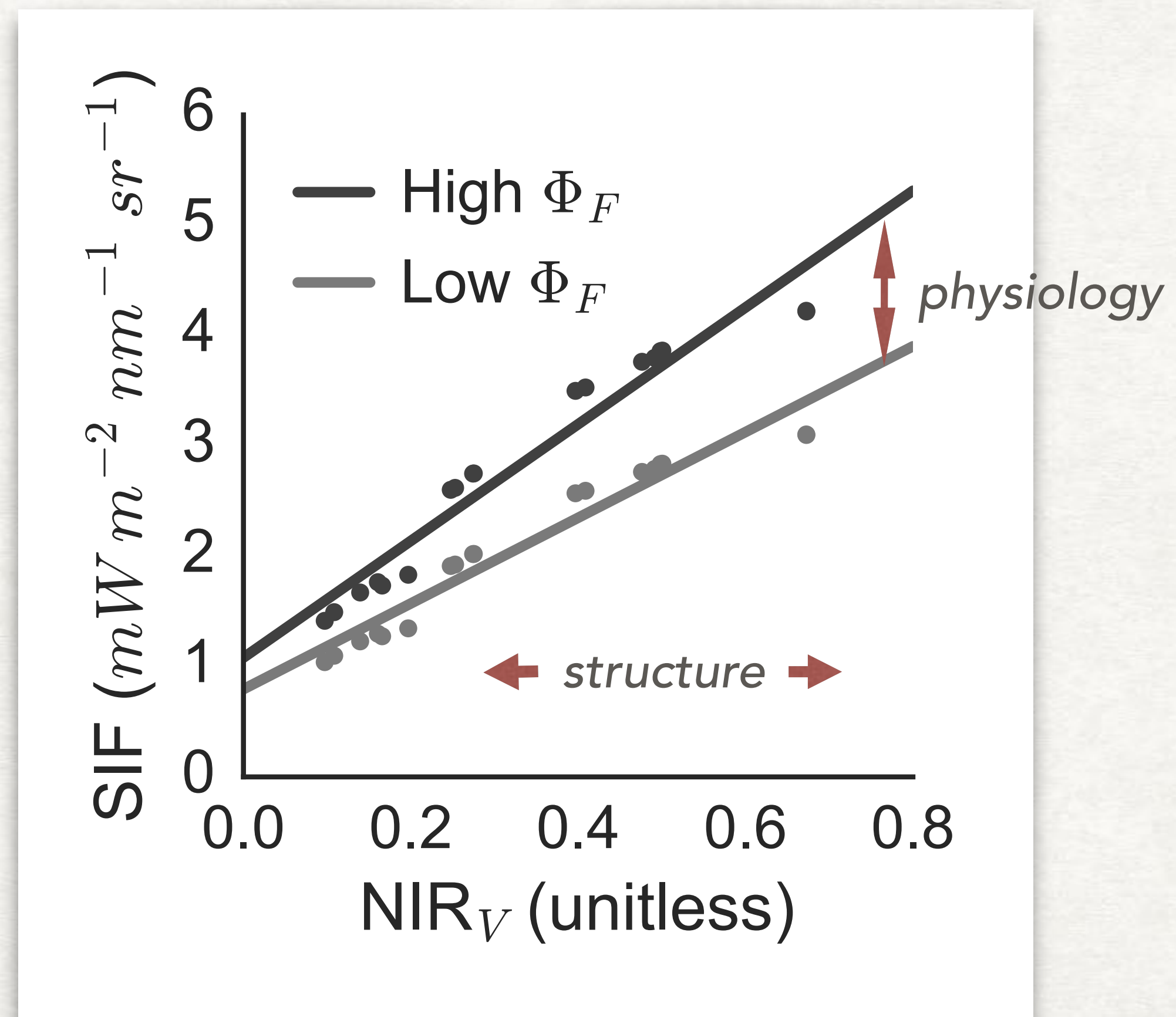
$$J_{\text{SIF}} = J_{\text{PAR}}(f_{\text{PAR}} \cdot \Phi_F \cdot \Omega_C)$$

↑

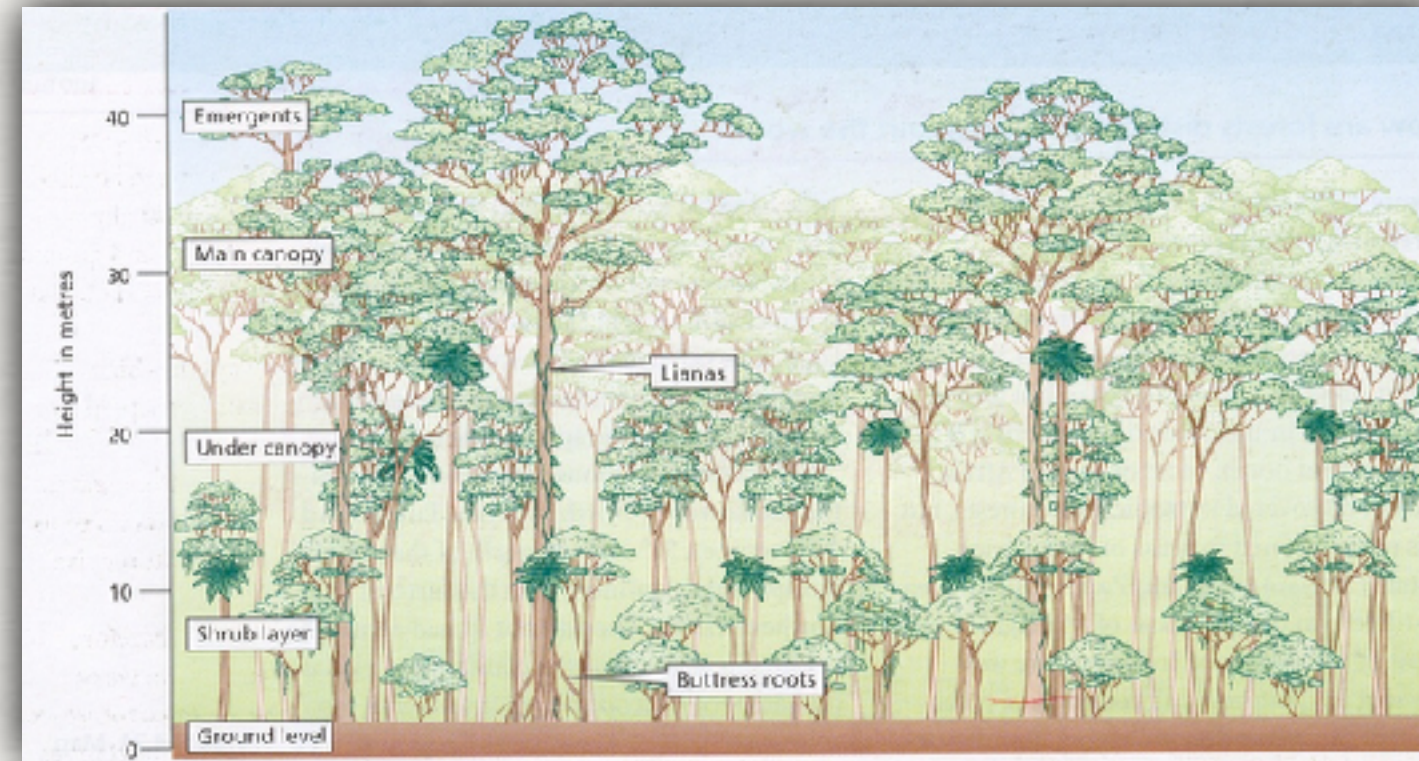
escape probability

The correlation between SiF and GPP is based on:

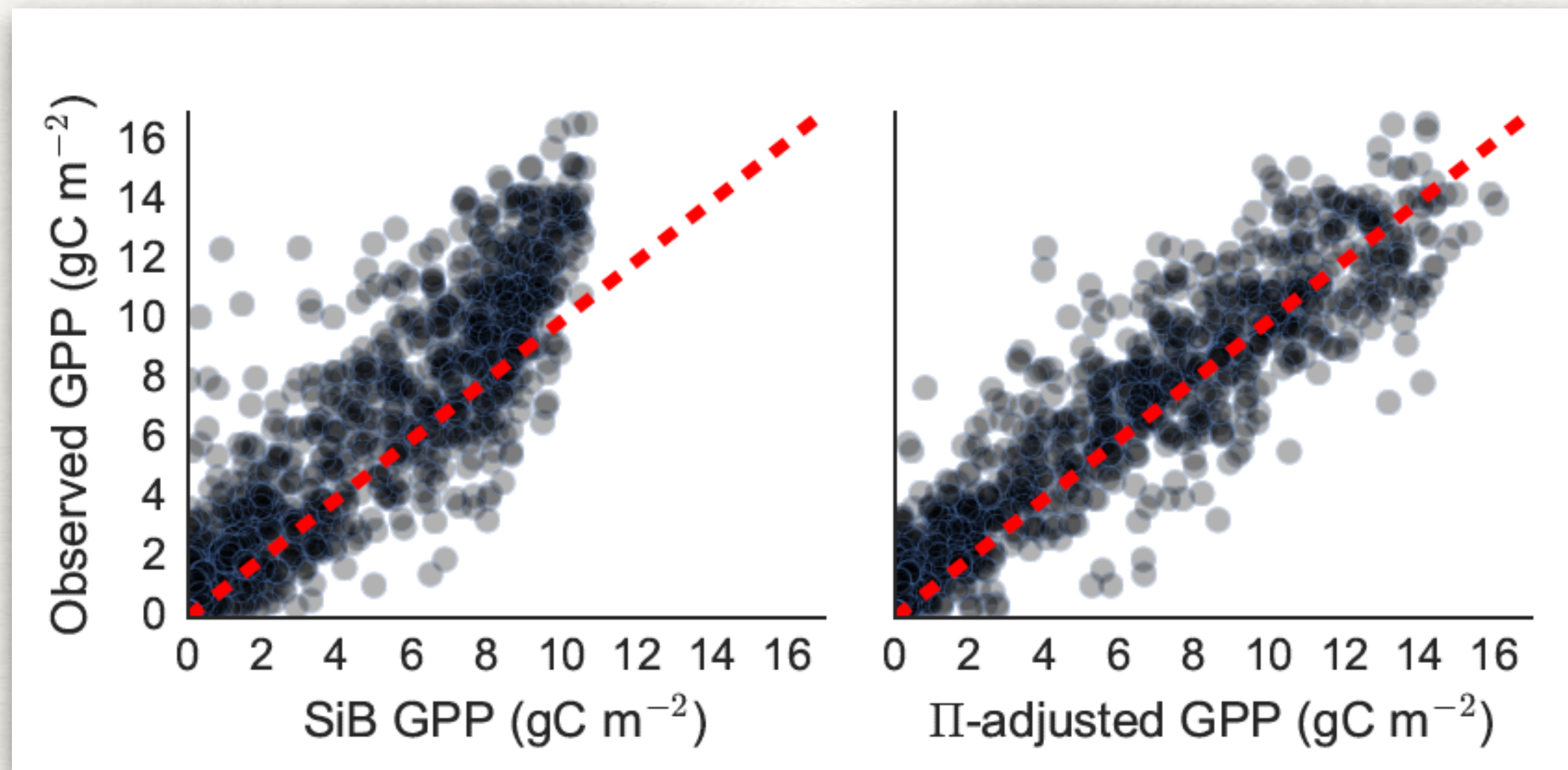
- Physiological control of the yield of fluorescence (Φ_F)
- Structural properties of the canopy that effect leaf display



Canopies are very diverse

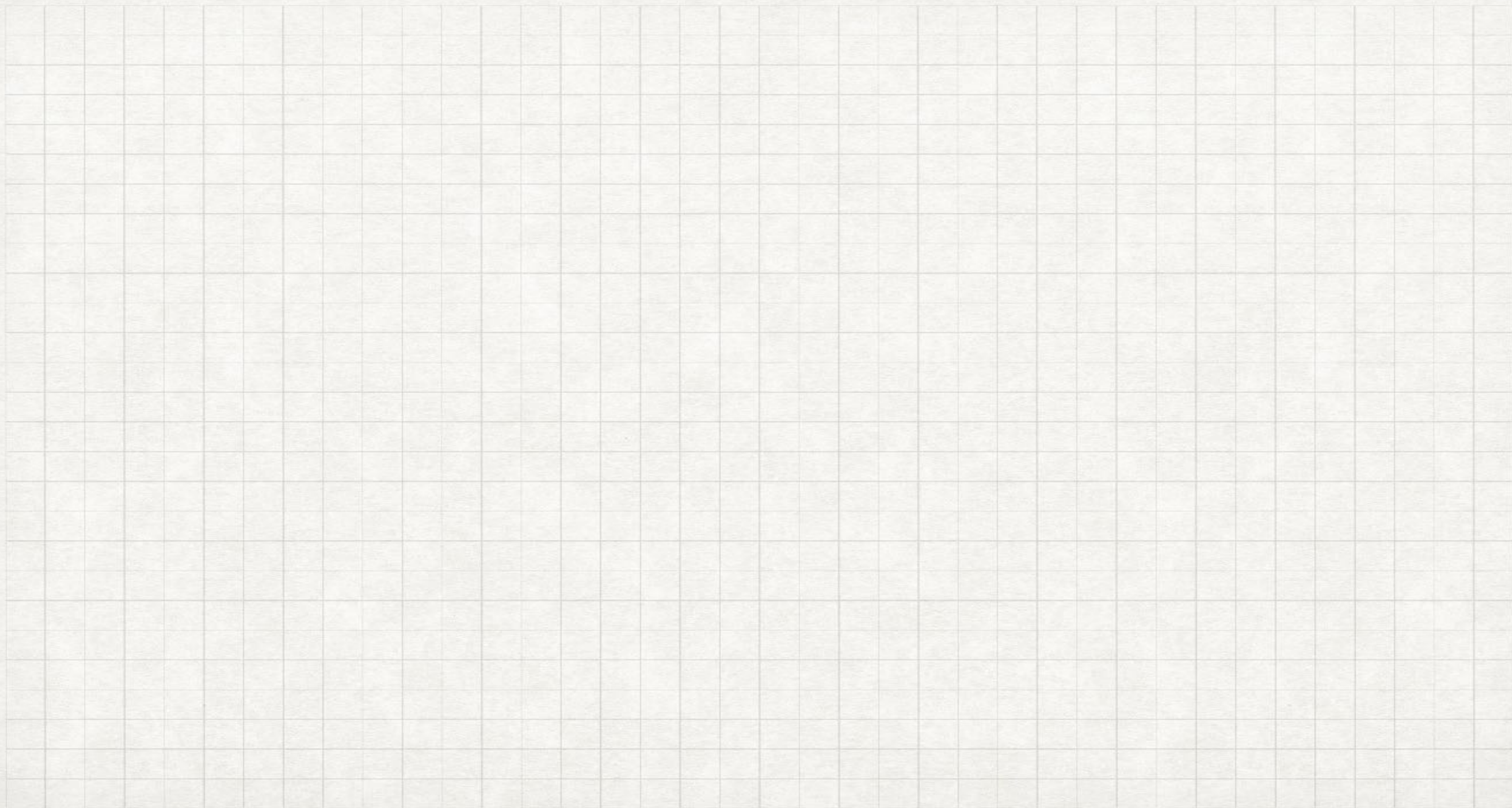


Scale π in SIB by SIF or NIR_v

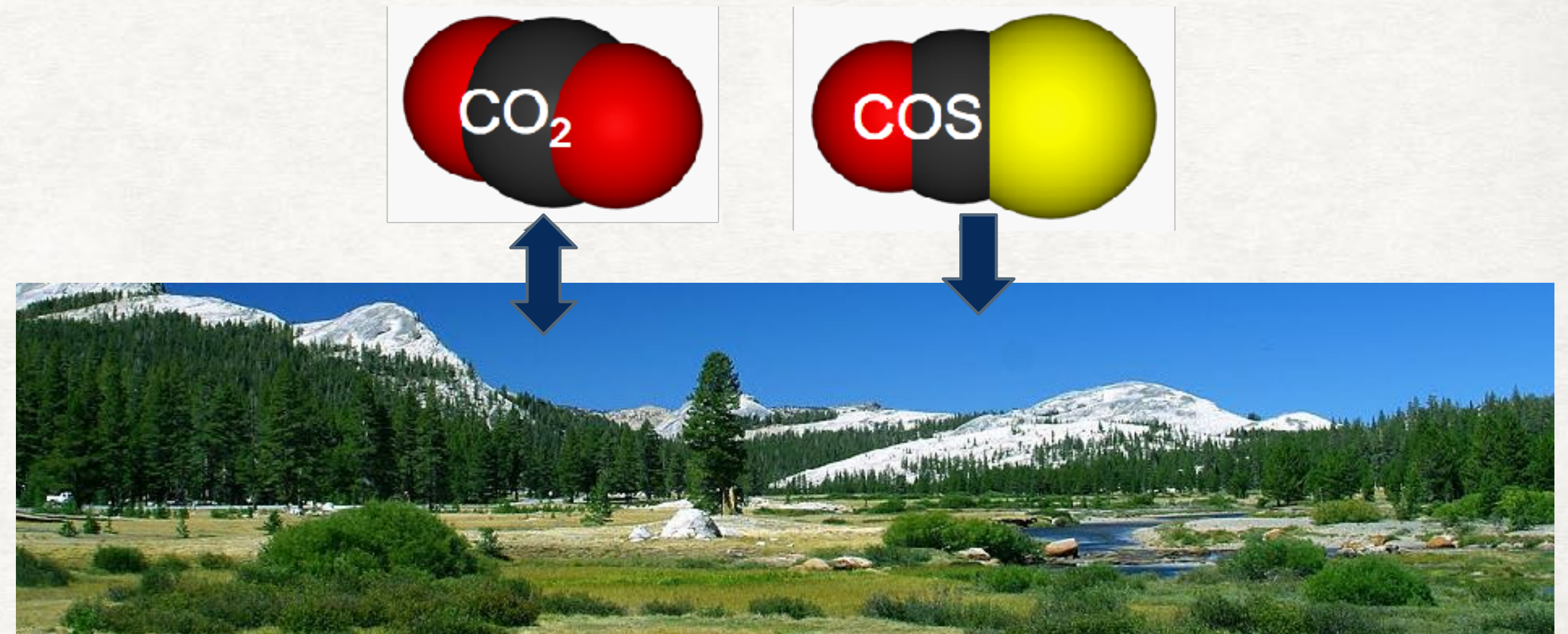
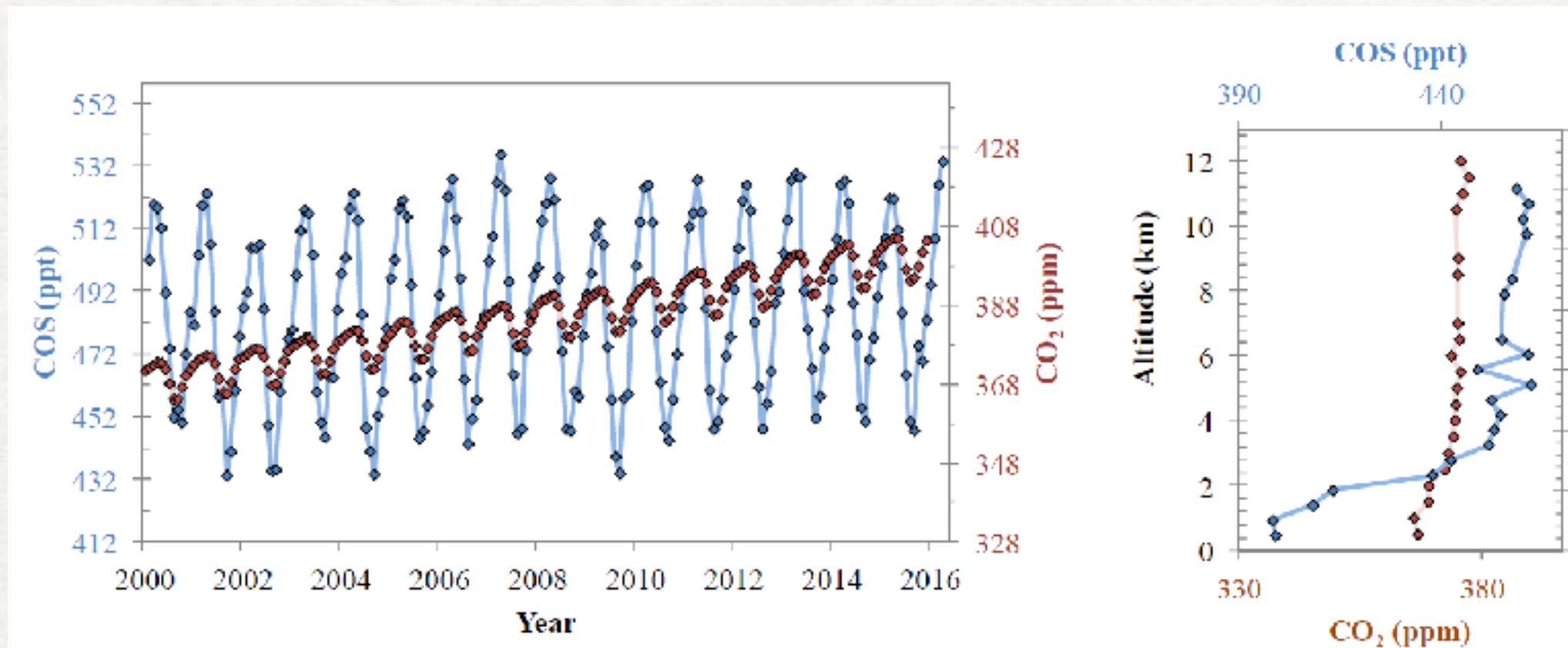


Summing up:

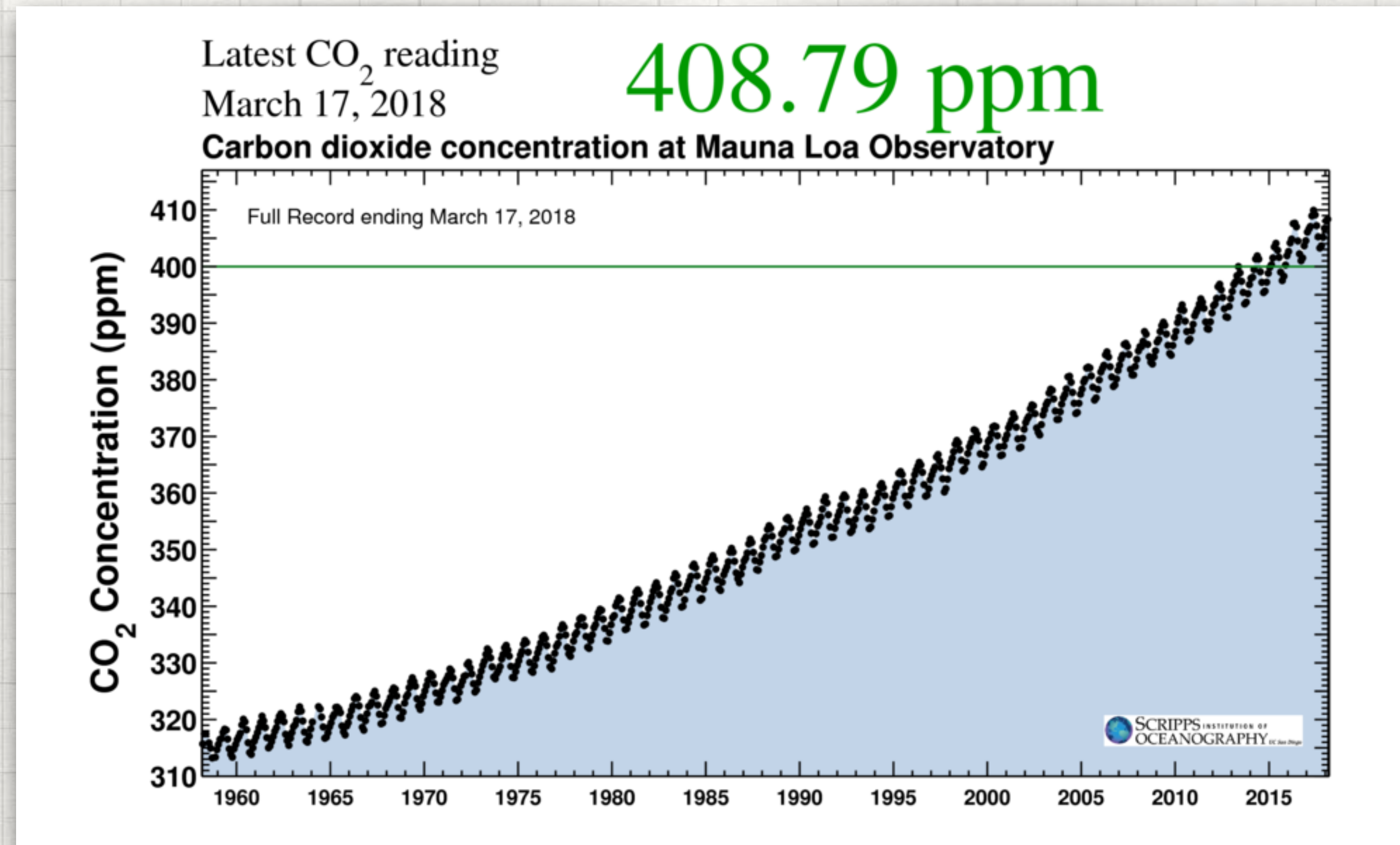
- Canopy structure seems to play a large role in SIF.
- This leaves us with the question of how does canopy structure have such a large control on the rate of photosynthesis.
- We are doing explicit 3-d modeling of leaf illumination in canopies of different structure and asking what would be the optimal distribution of photosynthetic capacity among the leaves.
- Preliminary results are promising - Sun leaves have more capacity and contribute more SIF.
- Leaf angle distribution is also important - using functions like "spherical" obscure important canopy properties.



COS - Another perspective on the Carbon Cycle

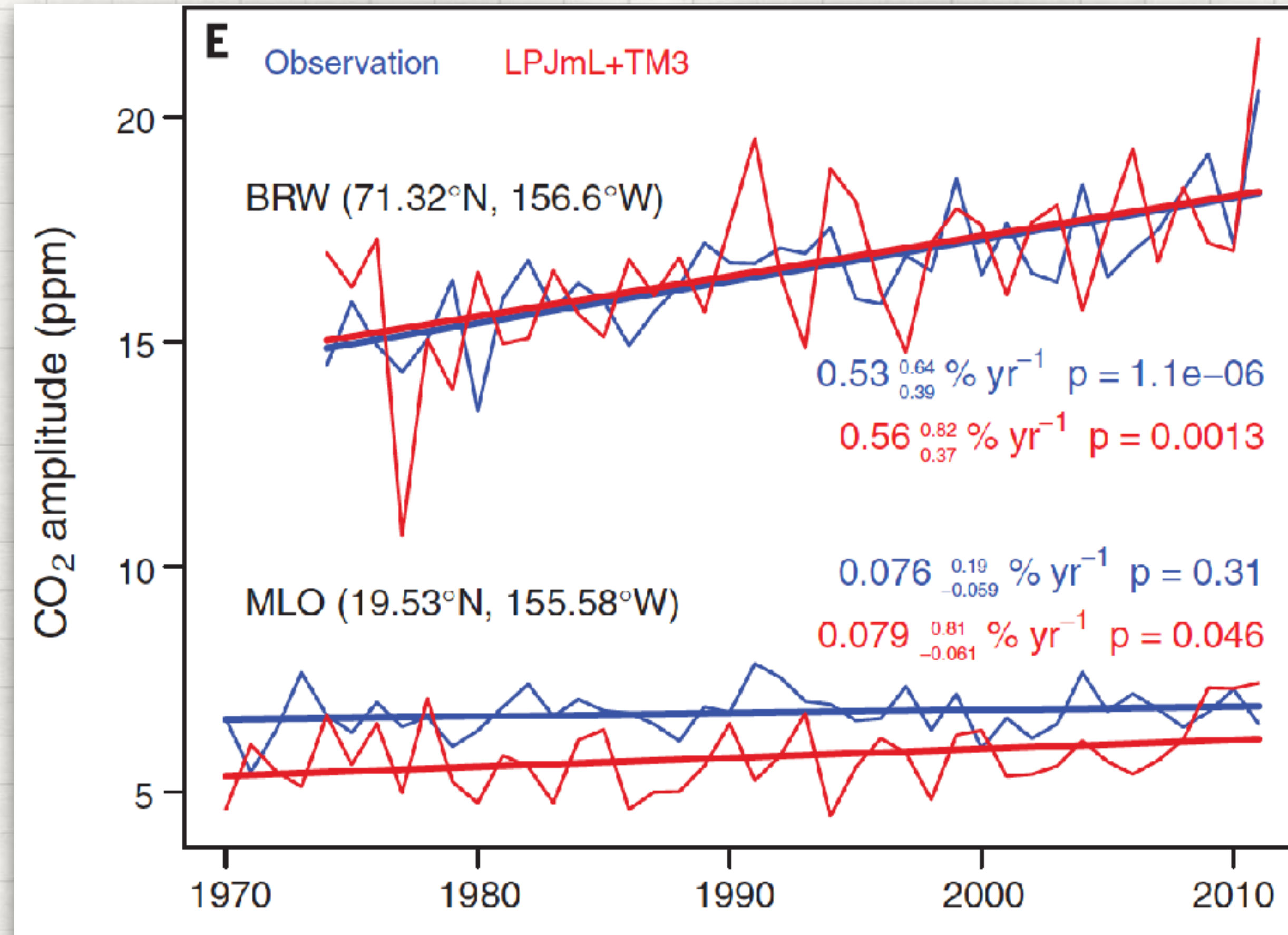


Enhanced Seasonal Exchange of CO₂ by Northern Ecosystems Since 1960



Graven, H. D., Keeling, R. F., Piper, S. C., Patra, P. K., Stephens, B. B., Wofsy, S. C., et al. (2013). Enhanced Seasonal Exchange of CO₂ by Northern Ecosystems Since 1960. *Science*. <http://doi.org/10.1126/science.1239207>

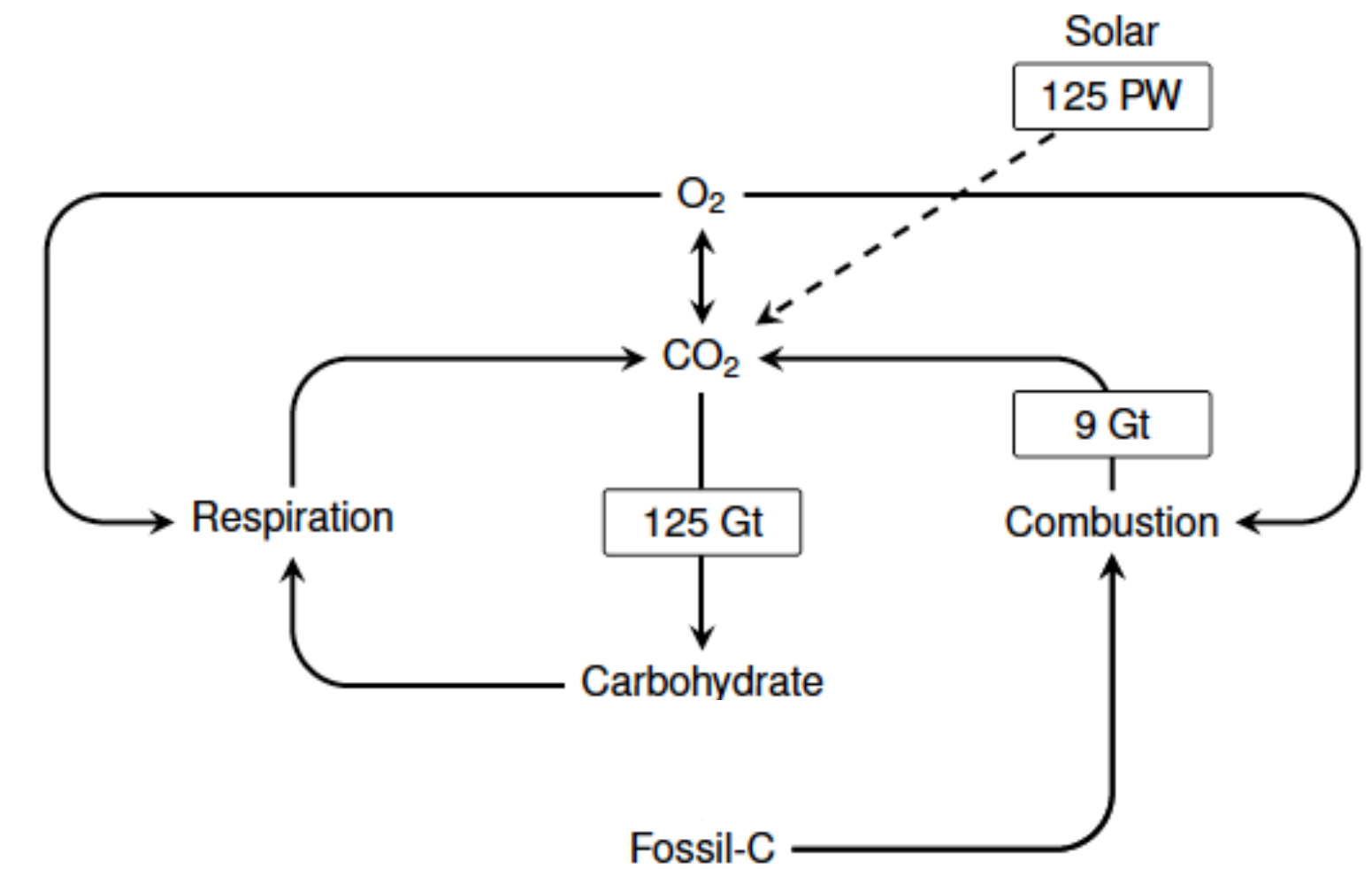
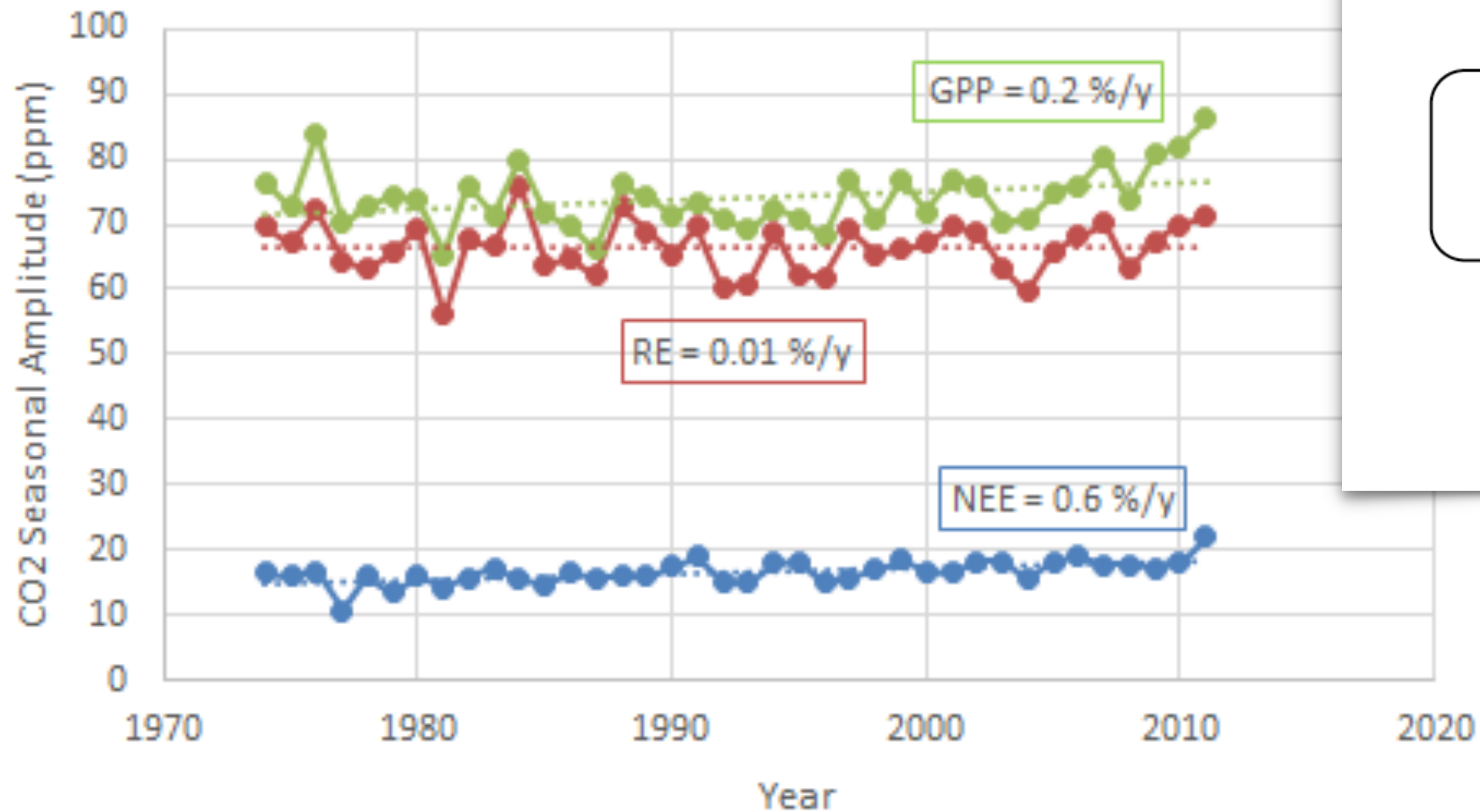
The peak to peak seasonal variation is increasing at Pt Barrow



Graven, H. D., Keeling, R. F., Piper, S. C., Patra, P. K., Stephens, B. B., Wofsy, S. C., et al. (2013). Enhanced Seasonal Exchange of CO₂ by Northern Ecosystems Since 1960. *Science*. <http://doi.org/10.1126/science.1239207>

Forkel, M., Carvalhais, N., Rödenbeck, C., Keeling, R., Heimann, M., Thonicke, K., et al. (2016). Enhanced seasonal CO₂ exchange caused by amplified plant productivity in northern ecosystems. *Science*, 351(6274), 696–699. <http://doi.org/10.1126/science.aac4971>

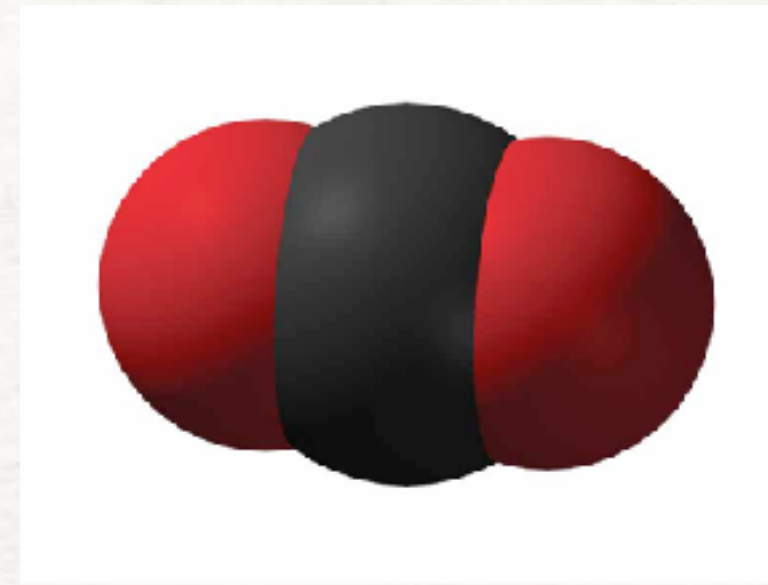
The observed response is the difference between photosynthesis and respiration.



- Respiration and photosynthesis respond to different factors and must be modeled separately.
- We have two variables and one observation.

What do we know about COS?

CO₂

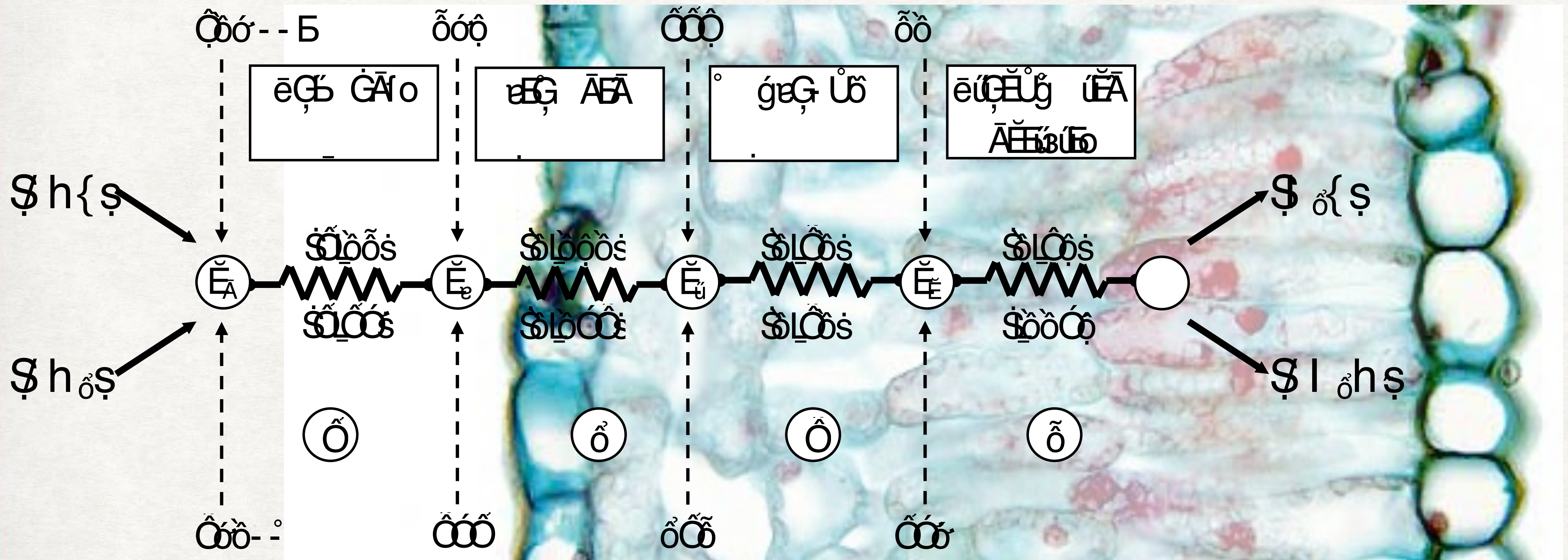


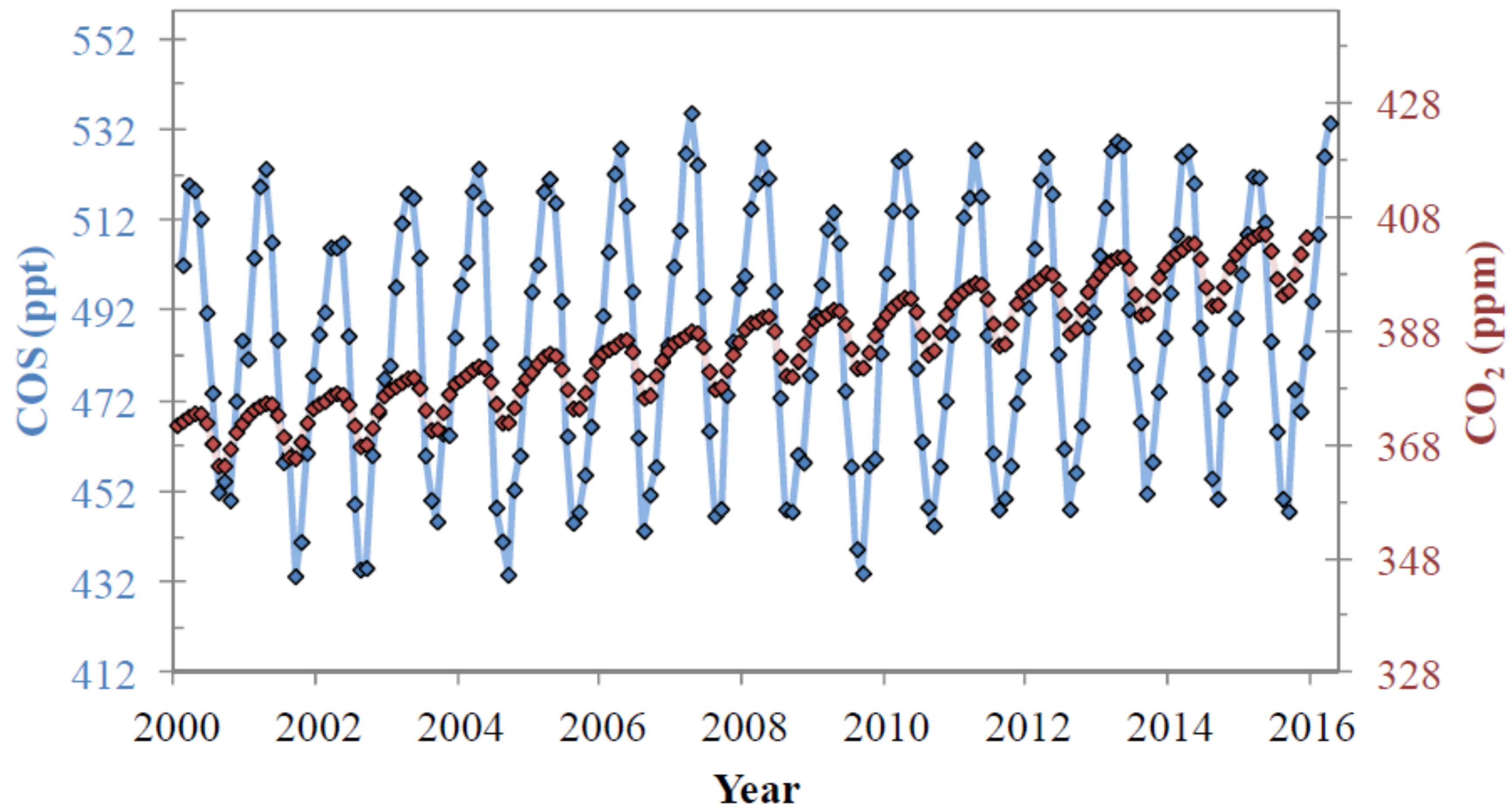
OCS or COS

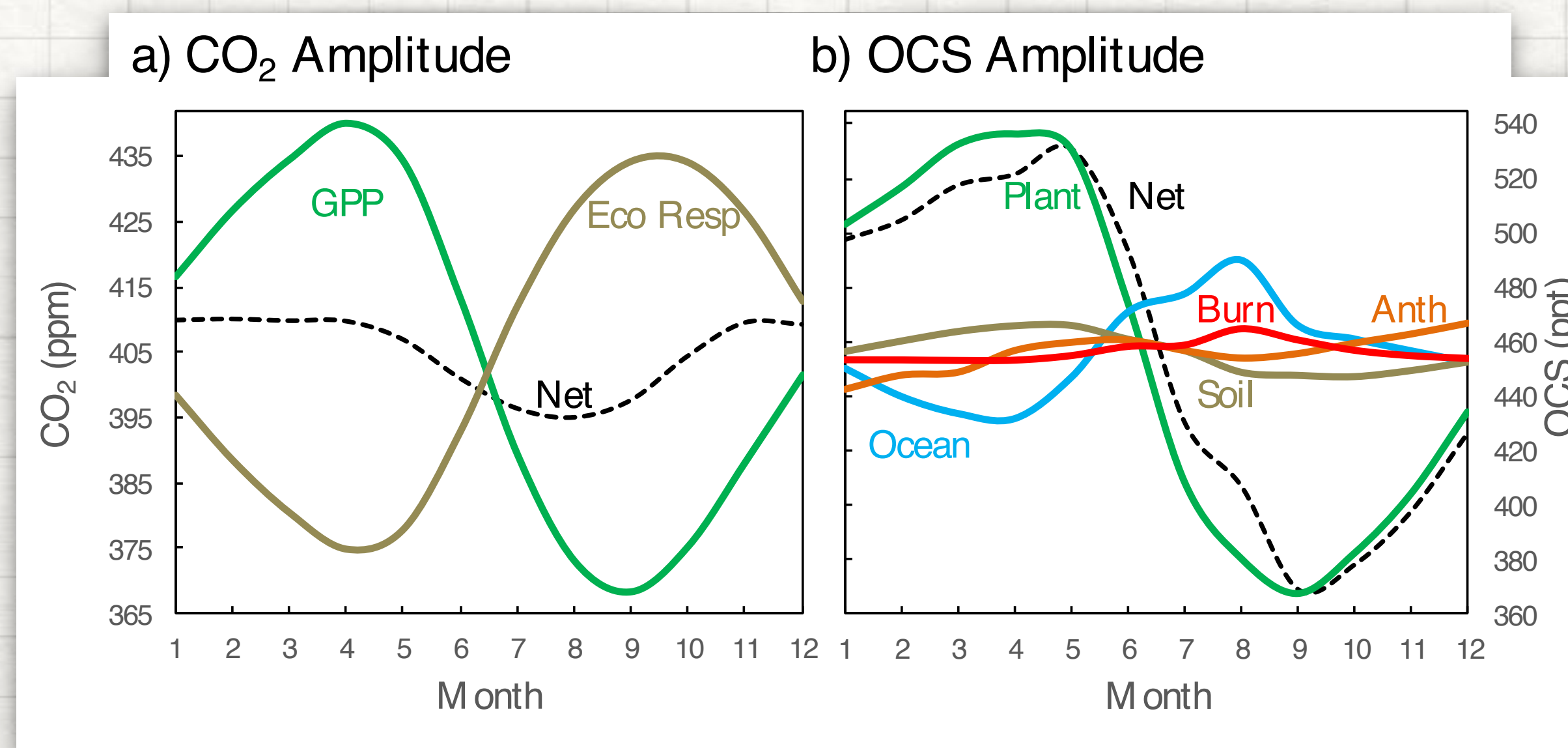


atmospheric concentration	400 ppm	500 ppt
turnover time	~7 years	~2 years
net atmospheric change	5% per year	30% per year
key enzymes	Rubisco + respiration	Carbonic-anhydrase

Leaf Uptake OCS follows the same path as CO₂

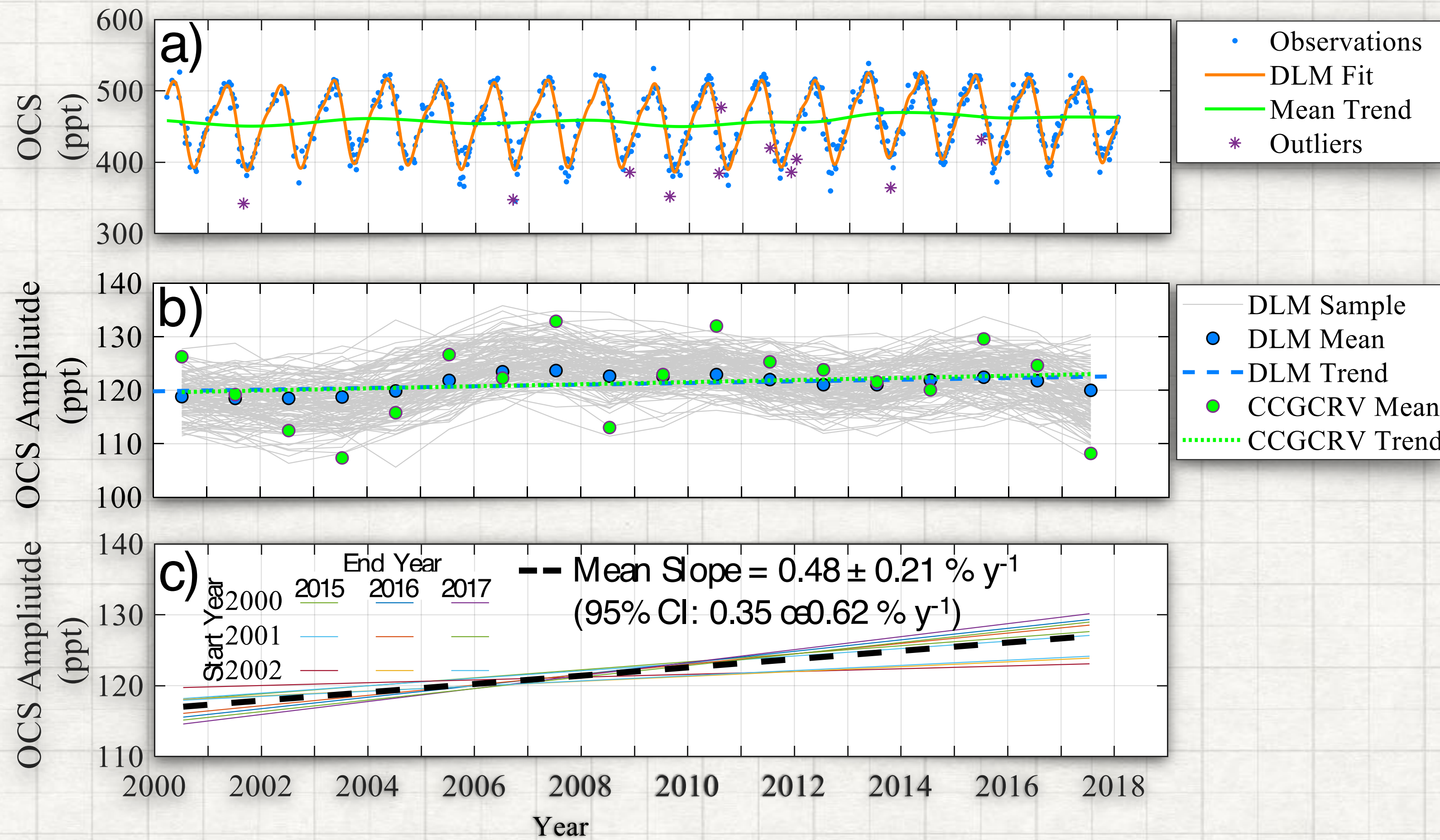


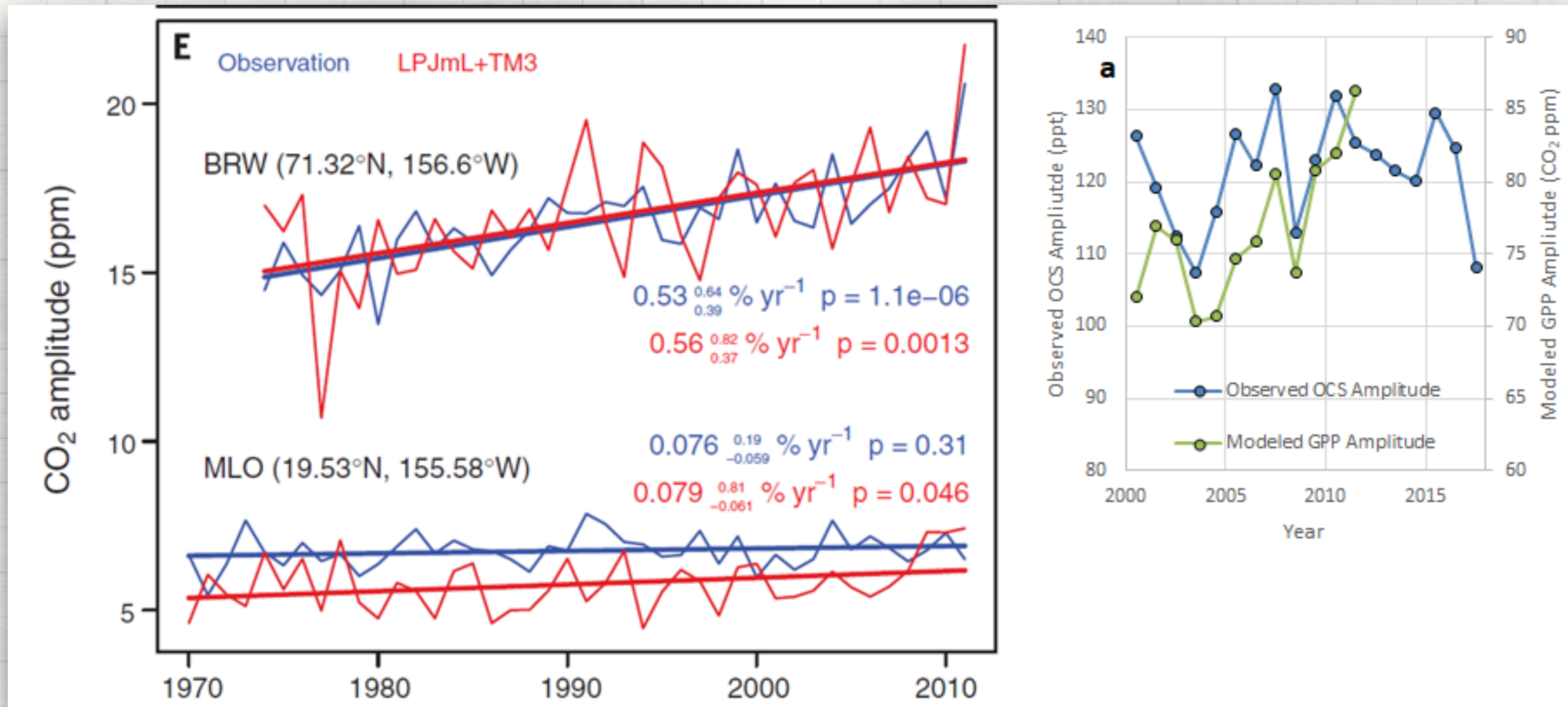




Global 3-D atmospheric transport simulations extracted at Barrow. Budget components at the Barrow site were obtained from global gridded source and sink inventories which were simulated in 3-D atmospheric chemical transport models (Kuai et al. 2015, Forkel et al. 2016). The dominant regional budget components influencing the CO₂ amplitude at Barrow are gross primary production (green) and the offsetting ecosystem respiration source (brown) which result in a relatively small net signal (black) **(A)**. For OCS, the plant sink (green) is the dominant influence on the Barrow seasonality which results in a similar net sink (black) **(B)**. Other OCS sources and sinks were also accounted for including soil (brown), biomass burning (red), industry (orange), and ocean (blue) budget components. Detailed methods are in the supporting online materials.

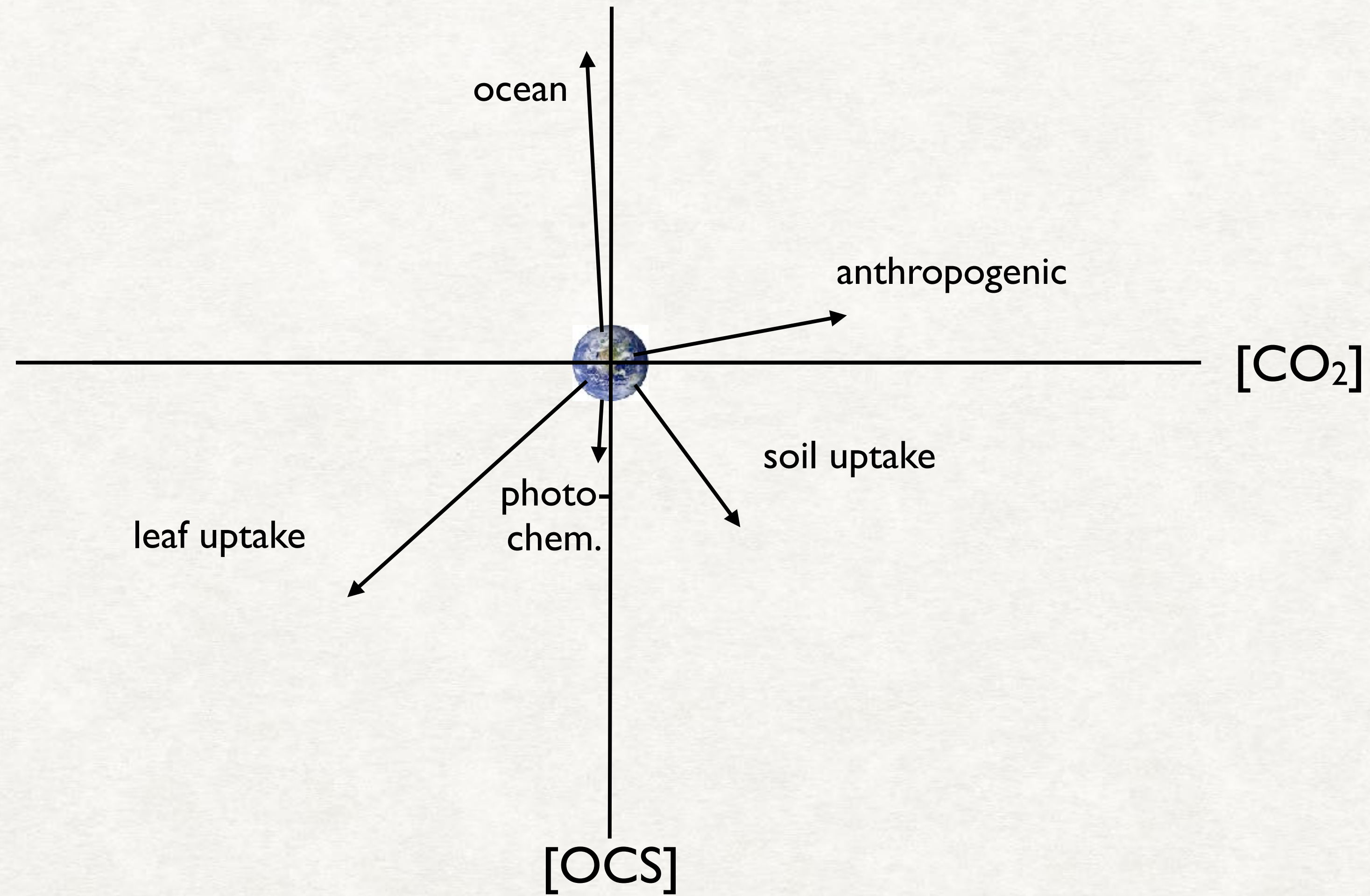
COS Seasonal Variation at Pt Barrow



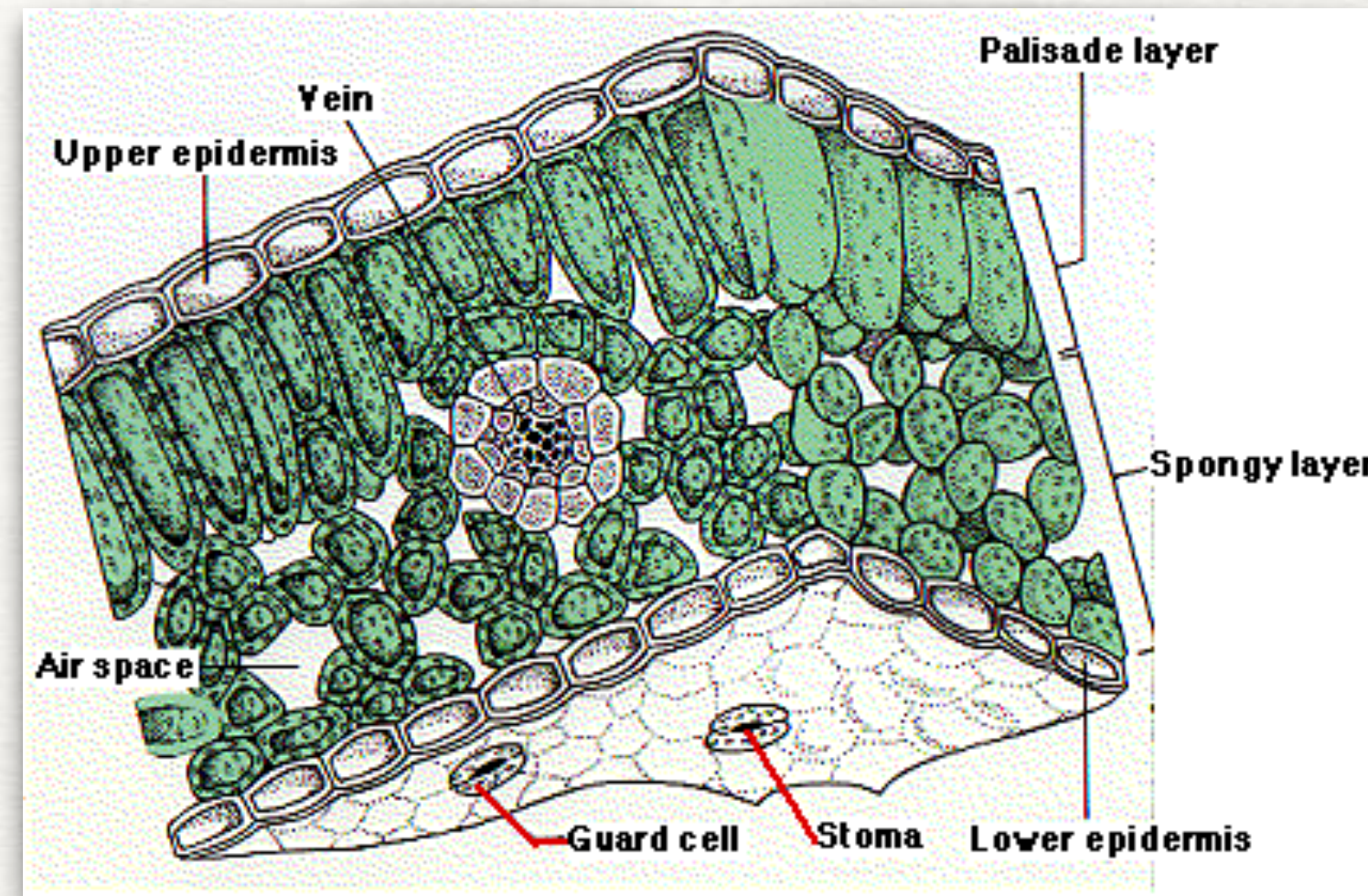


The measured COS seasonal variation (blue) matches the modeled GPP (green) fairly well. Forkel et al., are probably getting photosynthesis about right. It is increasing.

The atmosphere [CO₂] integrates the impact of all exchanges



Modeling Photosynthesis



Solving for Stomatal Conductance (g).

The approach takes advantage of the observation that g is generally highly correlated with the rate of Photosynthesis (A_n), and that we have reasonably good models of Photosynthesis.

$$g = m \cdot A_n \frac{h_s}{c_s} + b$$

Where: h_s and c_s are the relative humidity and CO_2 concentrations at the leaf surface, respectively; m and b are regression coefficients, and:

$$A_n = f(\text{PAR}, \text{Temp}, \text{CO}_2, \text{Stress}, V_{max})$$

Solving for CO_2 Assimilation (A_n)

C_3 : Farquhar *et al.*, assumed that photosynthesis could approach the capacity of the limiting step and that the identity of this limiting step could switch as environmental conditions change.

$$A_n \leq \min \begin{cases} J_e \\ J_c \\ J_s \end{cases} - R_d$$

$$J_e \text{ (Light limited rate)} = \alpha_p \epsilon_m Q_p \frac{(p_i - \Gamma^*)}{(p_i + 2\Gamma^*)}$$

$$J_c \text{ (Carboxylase limited rate)} = \frac{V_m (p_i - \Gamma^*)}{p_i + K_{app}}$$

$$J_s \text{ (Regeneration limited rate)} \approx \frac{V_m}{2}$$

$$R_d \approx 0.015 V_m$$

A_n	Assimilation ($\text{mol m}^{-2} \text{s}^{-1}$)
R_d	Mitochondrial respiration ($\text{mol m}^{-2} \text{s}^{-1}$)
p_i	partial pressure of CO_2 (Pa)
Q_p	light flux ($\text{mol m}^{-2} \text{s}^{-1}$)
Γ^*	CO_2 compensation point (Pa)
α_p	absorptance
ϵ_m	quantum yield (mol/mol)
V_m	Rubisco activity ($\text{mol m}^{-2} \text{s}^{-1}$)
K_{app}	Rubisco apparent K_m (Pa)
	$= K_m (1 + \frac{[O_2]}{K_i})$
Γ^*	$= \frac{[O_2]}{2\tau}$
	τ ← Rubisco specificity factor

A coupled model of the global cycles of carbonyl sulfide and CO₂: A possible new window on the carbon cycle

Joe Berry,¹ Adam Wolf,² J. Elliott Campbell,³ Ian Baker,⁴ Nicola Blake,⁵ Don Blake,⁵
A. Scott Denning,⁴ S. Randy Kawa,⁶ Stephen A. Montzka,⁷ Ulrike Seibt,⁸ Keren Stimler,⁹
Dan Yakir,⁹ and Zhengxin Zhu⁶

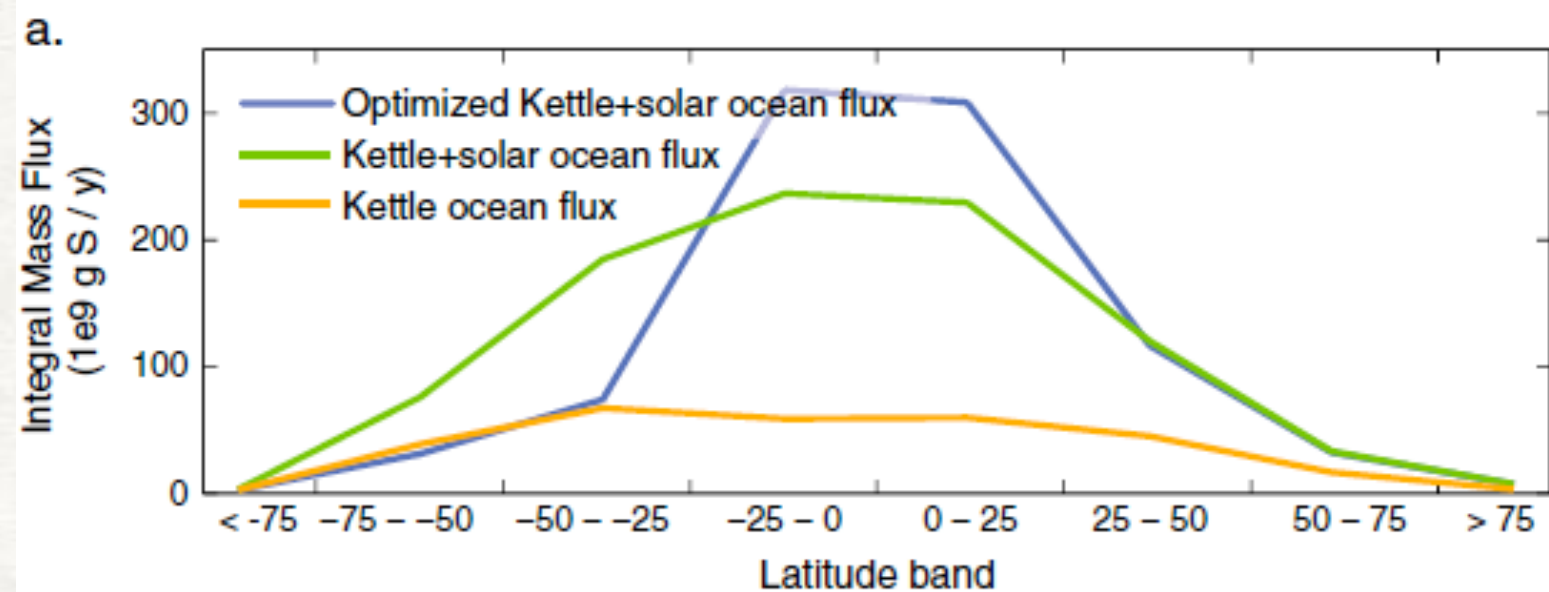
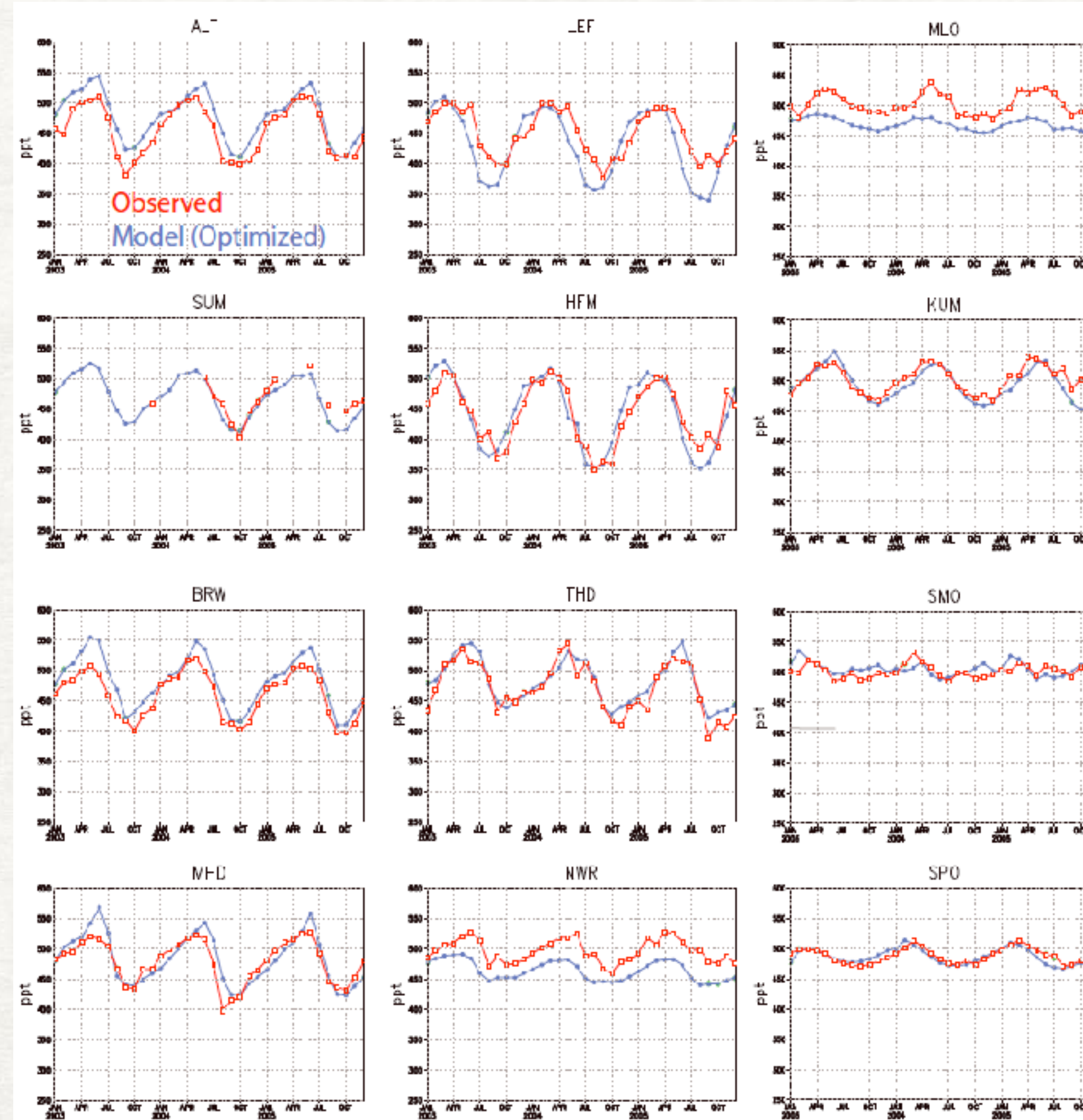


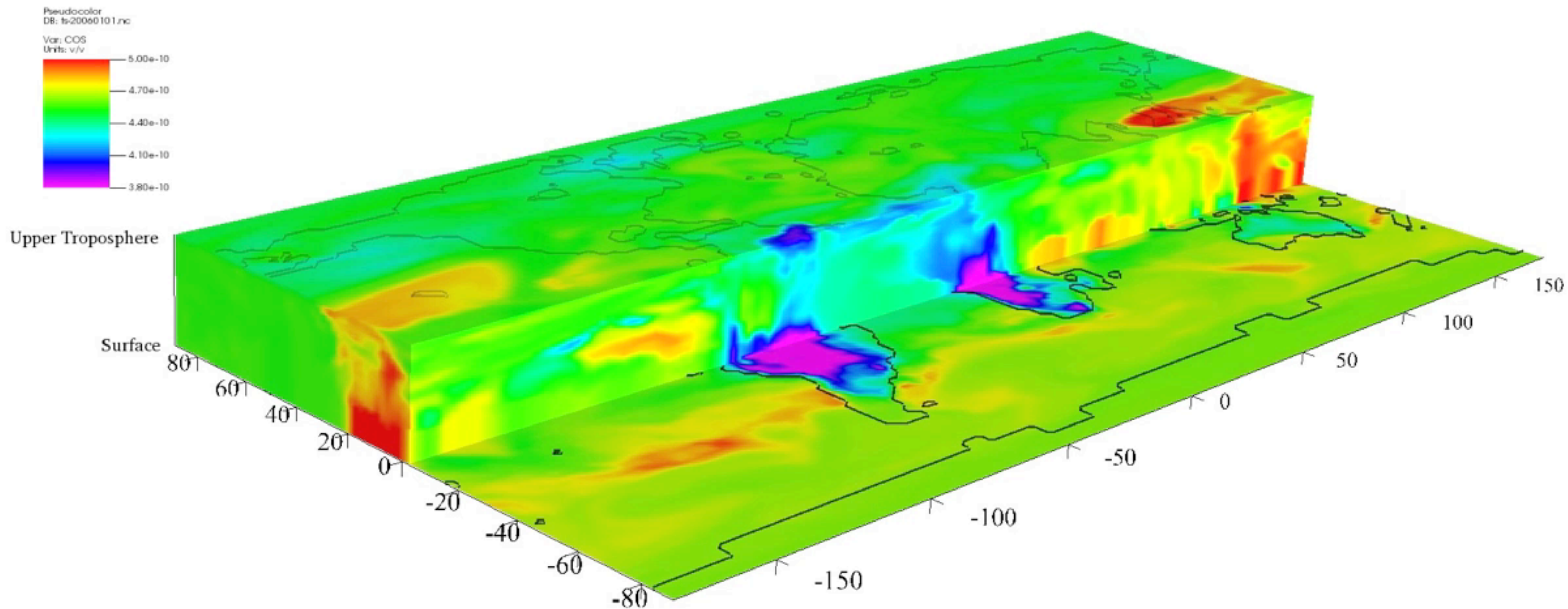
Table 1. A Compilation of the Global Sources and Sinks Used for PCTM Simulations of Atmospheric COS^a

Sources	<i>Kettle et al.</i> , 2002	This Study
Direct COS Flux From Oceans	39	39
Indirect COS Flux as DMS From Oceans	81	81
Indirect COS Flux as CS ₂ From Oceans	156	156
Direct Anthropogenic Flux	64	64
Indirect Anthropogenic Flux From CS ₂	116	116
Indirect Anthropogenic Flux From DMS	0.5	0.5
Biomass Burning	11	136
Additional (Photochemical) Ocean Flux		600
<i>Sinks</i>		
Destruction by OH Radical	-94	-101
Uptake by Canopy	-238	-738
Uptake by Soil	-130	-355
Net Total	-5	-2.5

^aUnits are 1.0×10^9 g of sulfur. Fluxes changed in this study are highlighted with bold type.

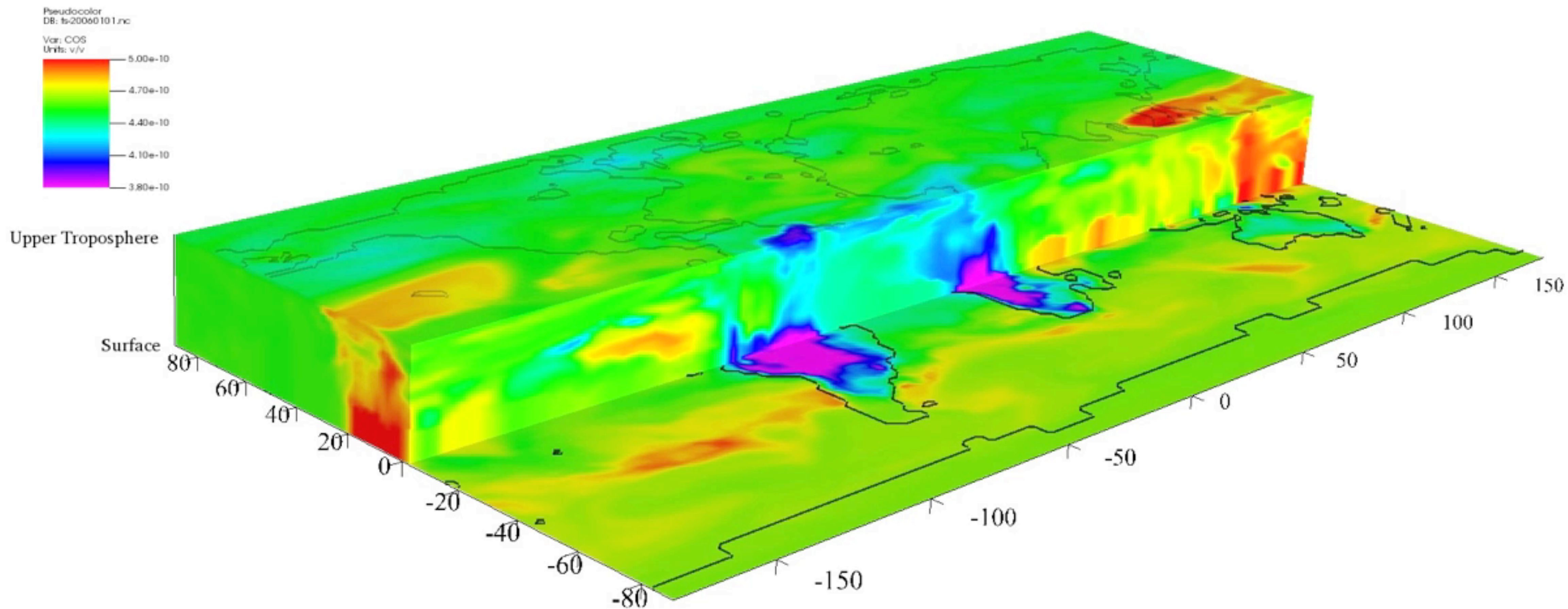


[OCS] vertical slice at Equator



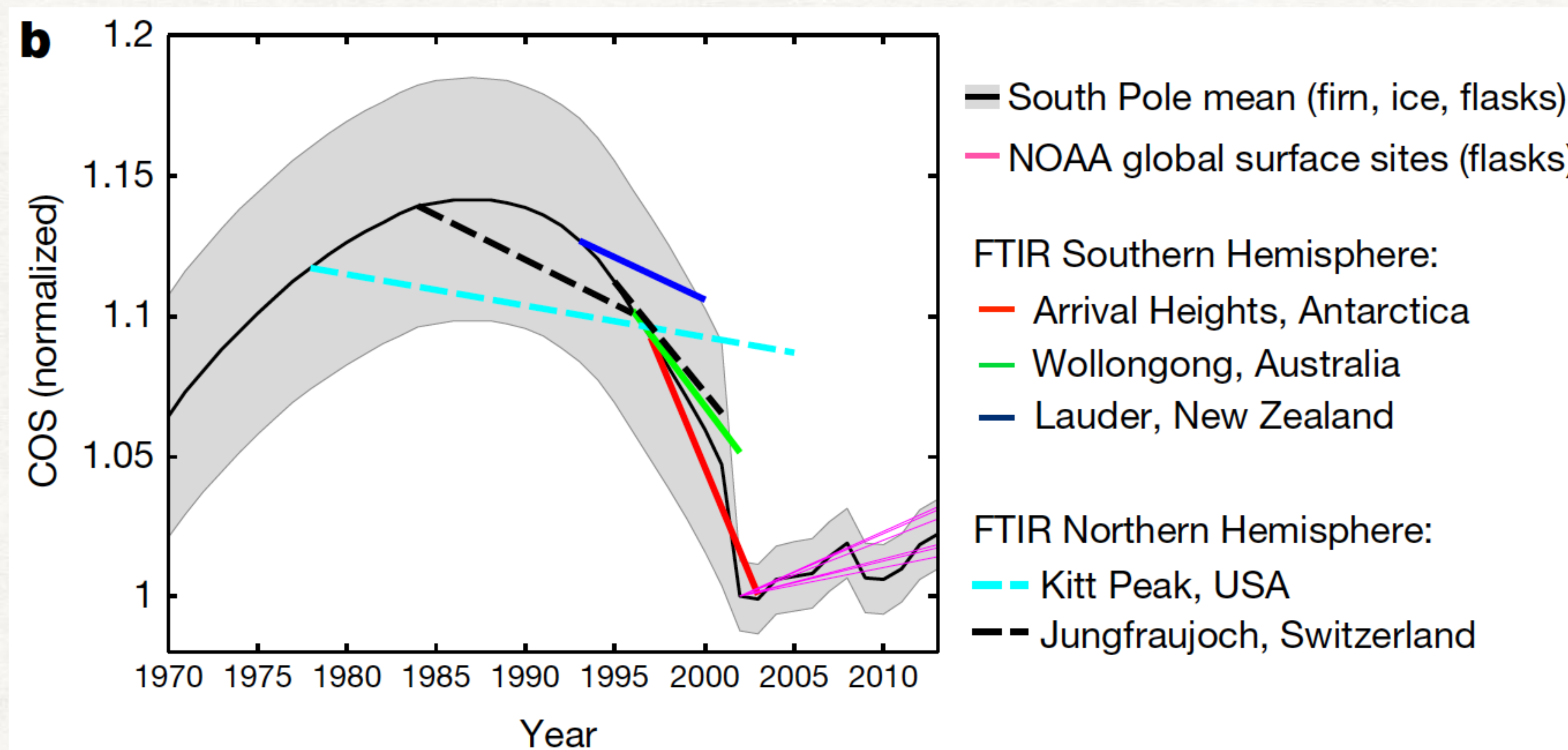
Jim Stineciper and Elliott Campbell (unpublished)

[OCS] vertical slice at Equator



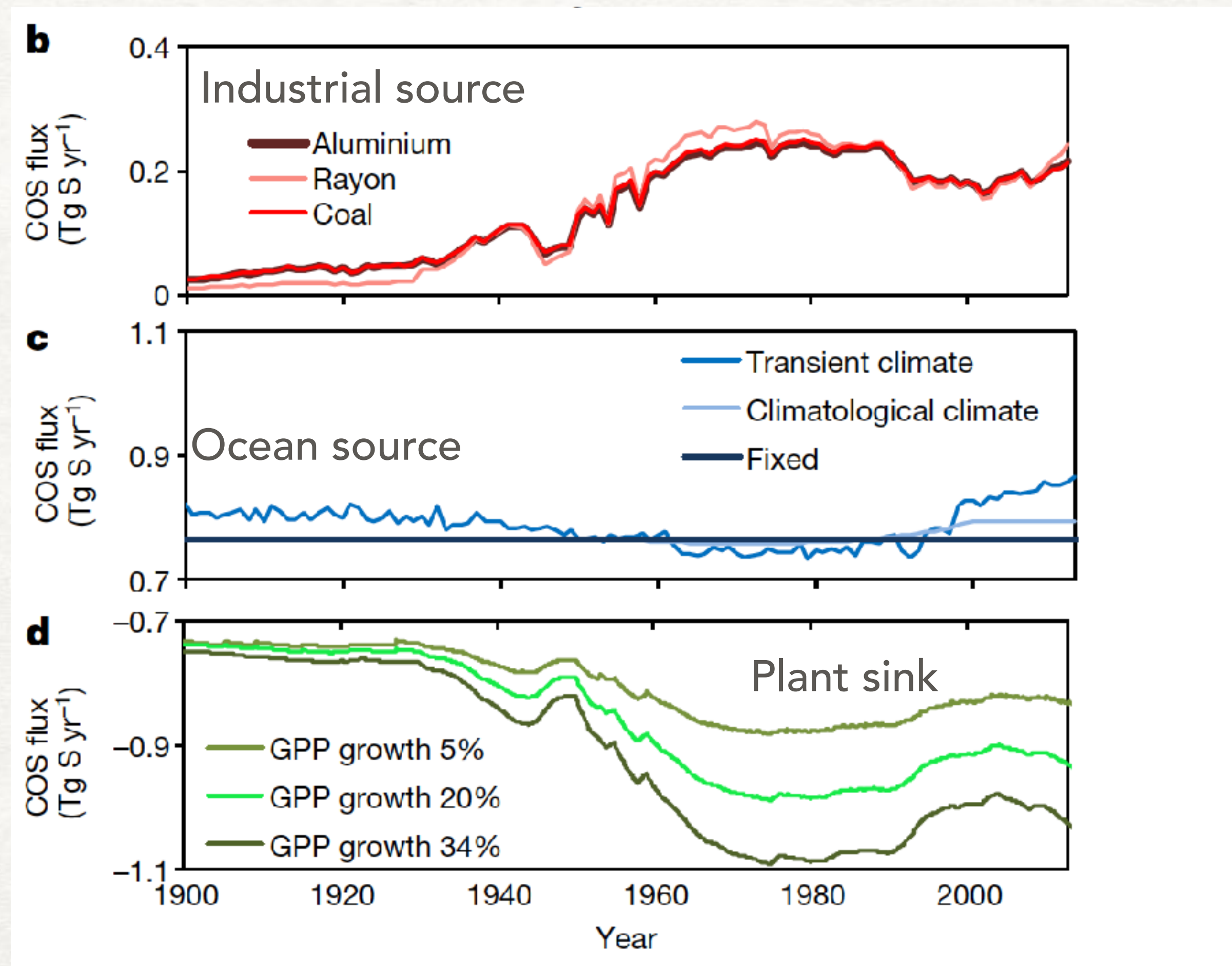
Jim Stinecipher and Elliott Campbell (unpublished)

A history of [COS] can be obtained from Antarctic ice

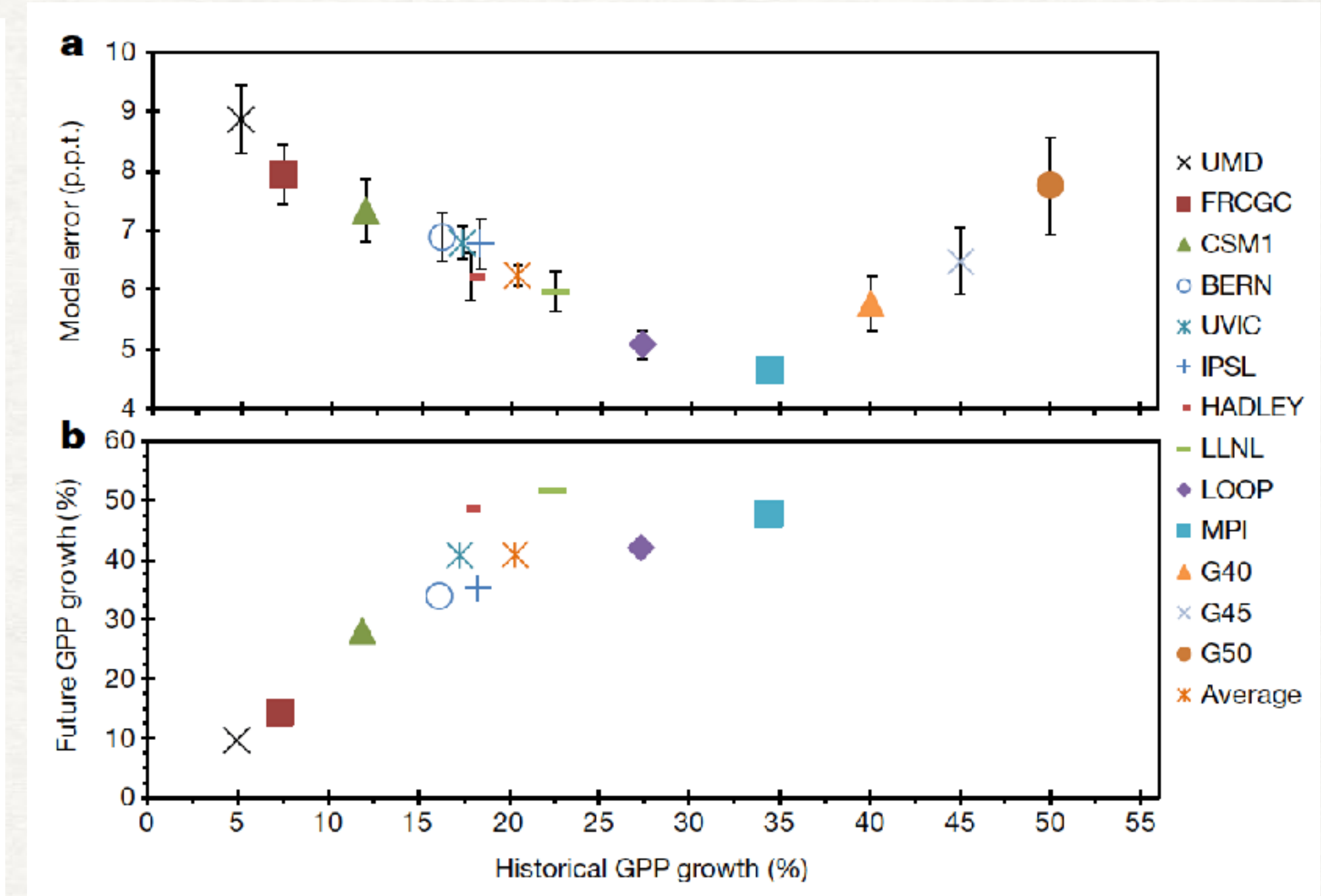
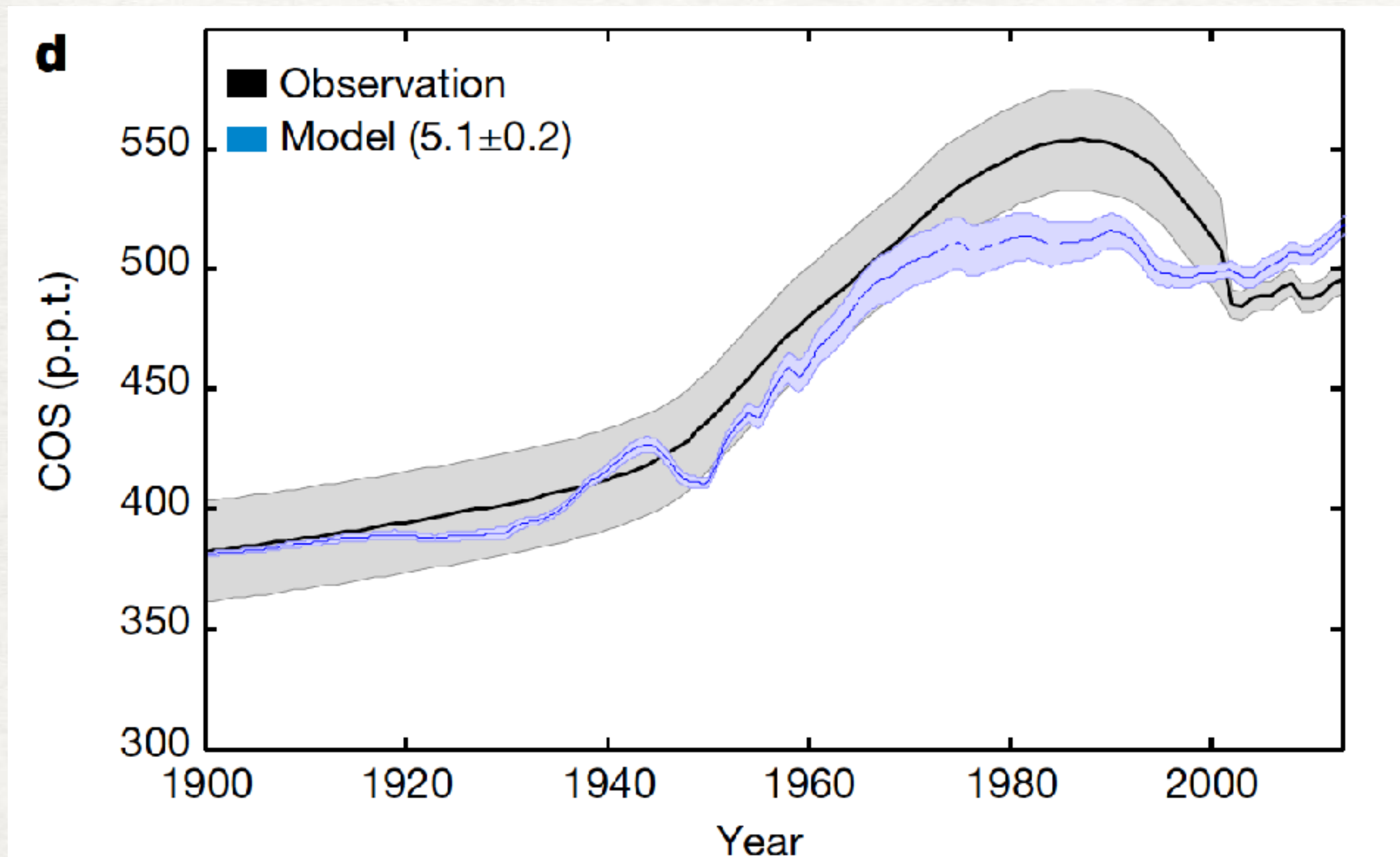


Can this tell us something about past photosynthesis?

Fluxes that contribute to the changing global [CO₂]



Find the GPP to minimize the error in COS growth



Models that predict moderate stimulation of photosynthesis by CO₂ best predict the observed COS growth

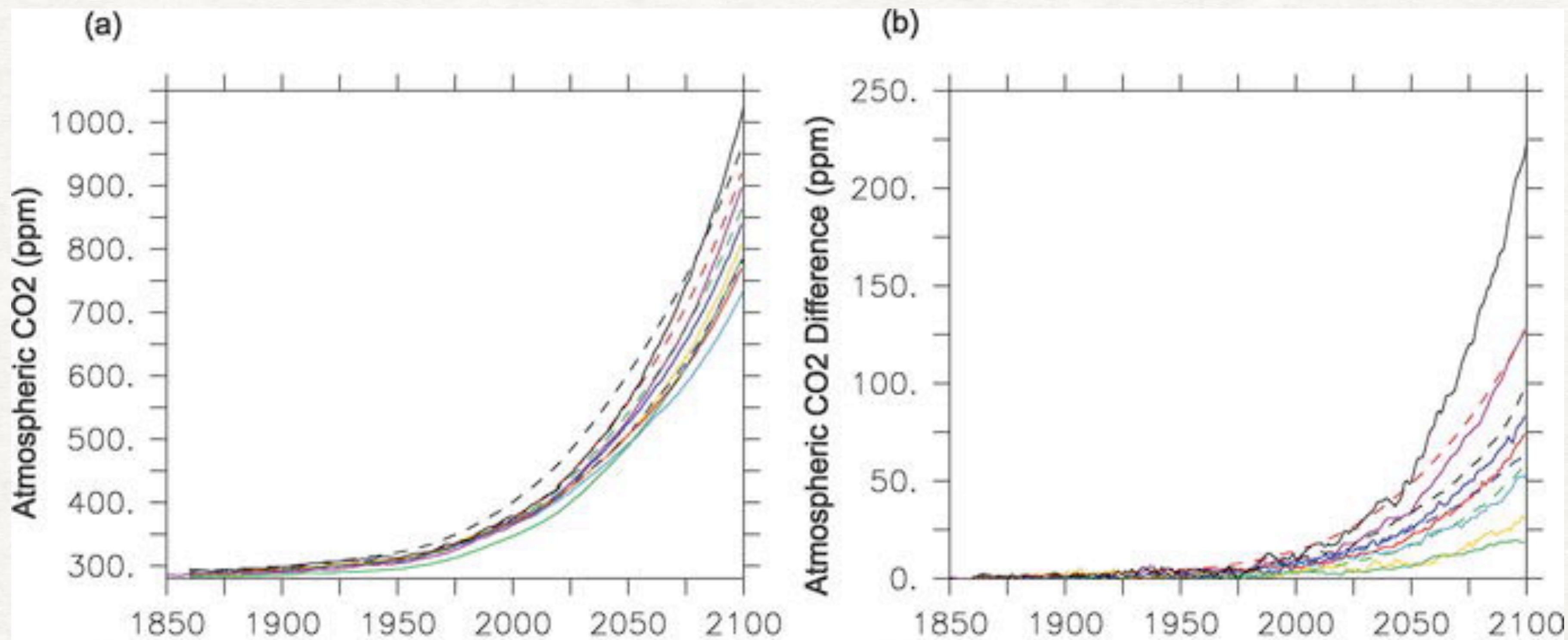
LETTER

doi:10.1038/nature22030

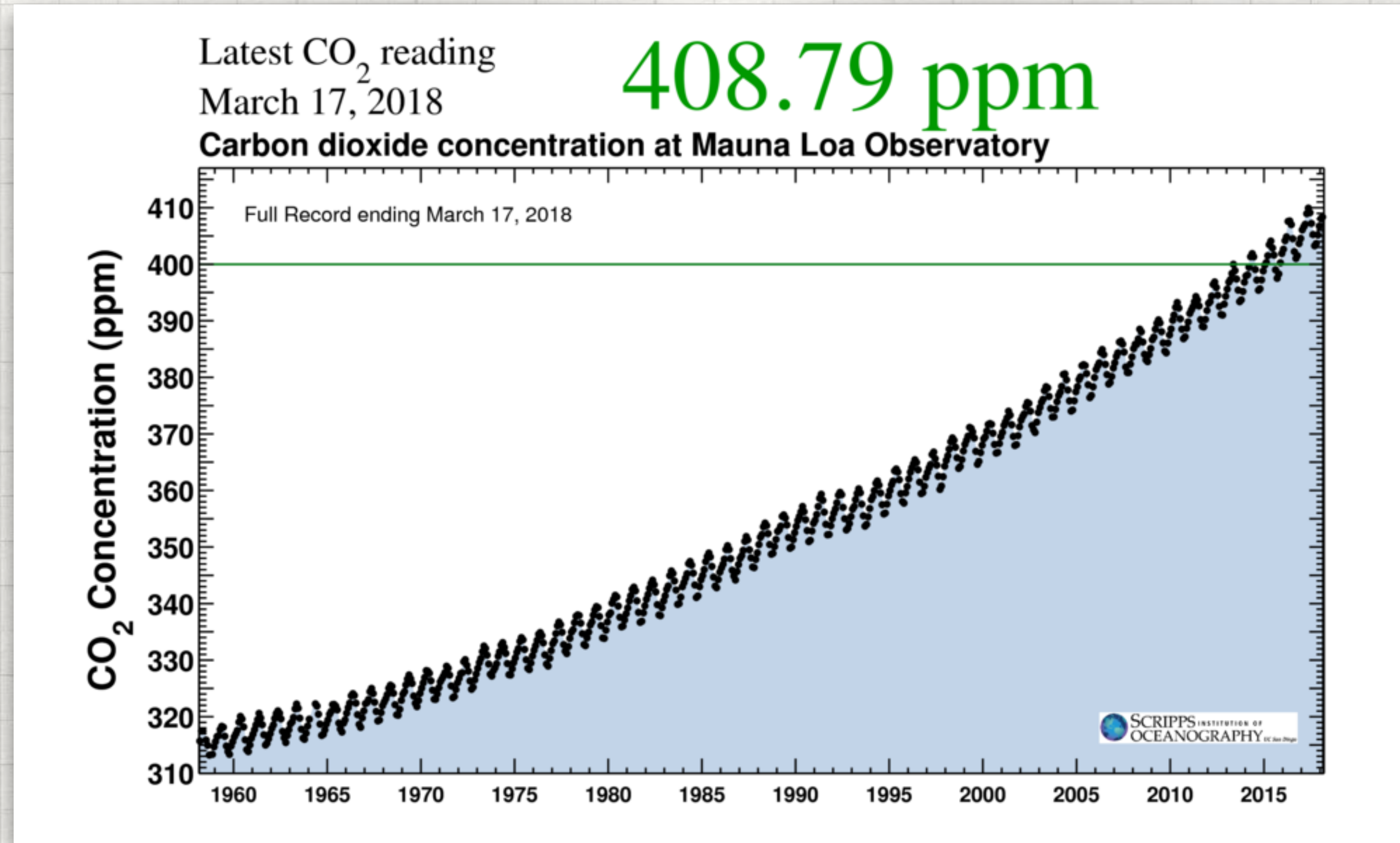
Large historical growth in global terrestrial gross primary production

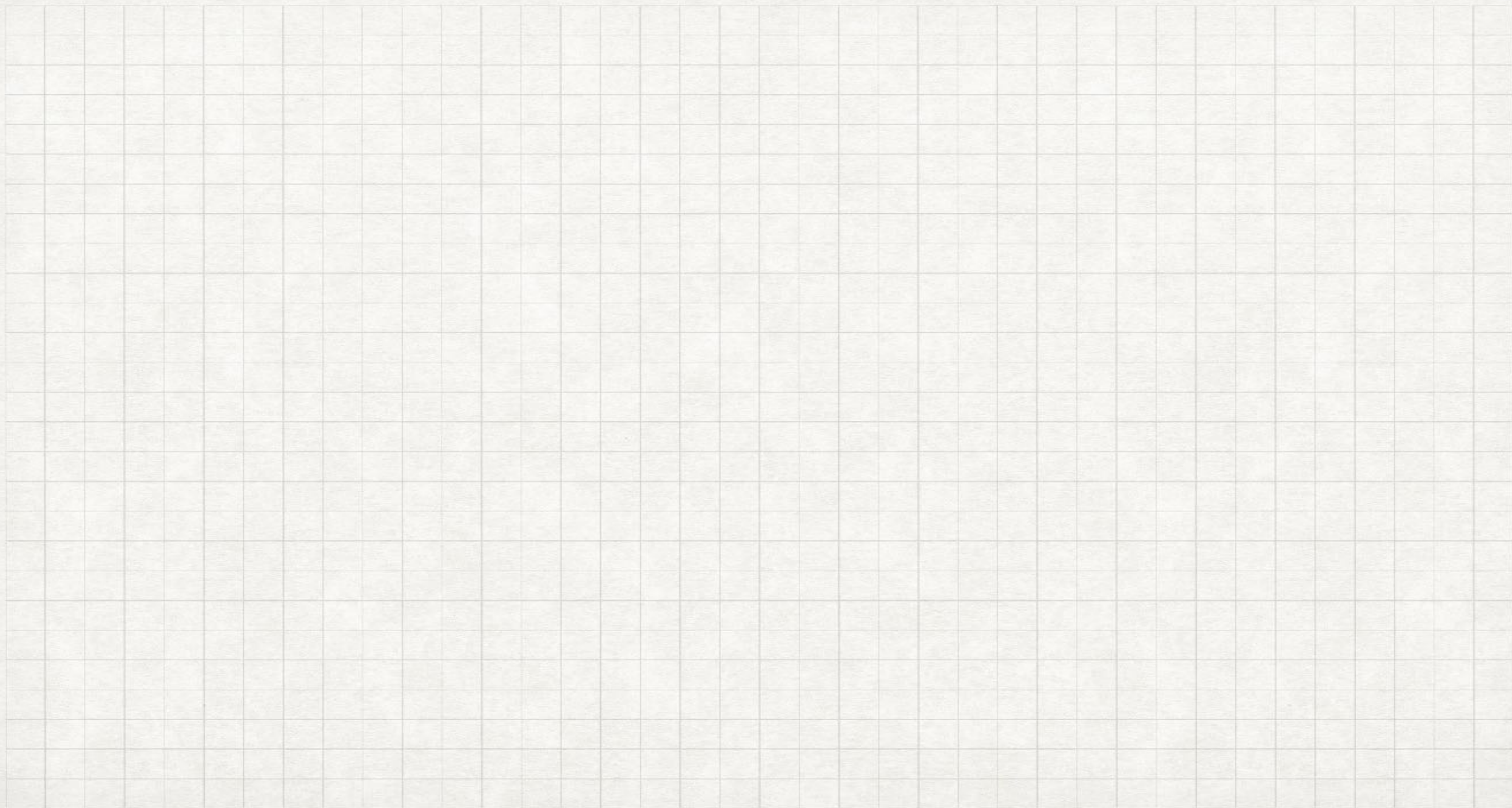
J. E. Campbell¹, J. A. Berry², U. Seibt³, S. J. Smith⁴, S. A. Montzka⁵, T. Launois^{6†}, S. Belviso⁶, L. Bopp^{6†} & M. Laine⁷

Growth in terrestrial gross primary production (GPP)—the amount of carbon dioxide that is ‘fixed’ into organic material through the photosynthesis of land plants—may provide a negative feedback for climate change^{1,2}. It remains uncertain, however, to what extent this growth has occurred since the industrial revolution. Here we use a combination of ground-based measurements and global GPP estimates for the total atmospheric burden of CO₂. The Antarctic record derived from measurements of air trapped in Antarctic ice and firn (granular snow deposited in previous years), and from ambient air samples—is consistent with independent long-term data from ground-based measurements and global GPP estimates (Fig. 1). The Antarctic record shows a steady increase in CO₂ concentration from 1850 to 2010, with a sharp increase after 1950. This increase is consistent with the increase in CO₂ concentration shown in Fig. 1(a).

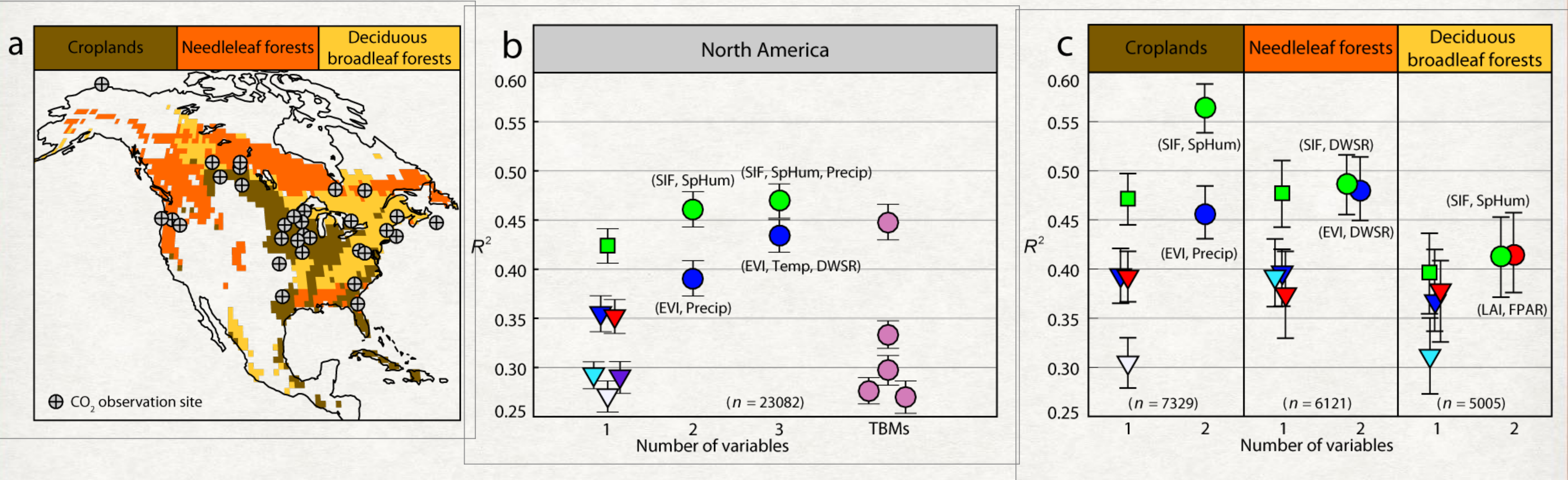


Ok, We have a feedback mechanism. Is it adequate to control raising [CO₂]?





SIF alone explains more variation in regional CO₂ inversions than most TBMs (terrestrial biosphere models)



Yoichi Shiga & Anna Michalak (Carnegie)