

Breaking with tradition: Dealing with unavoidable but imperfectly observed physics in ice-sheet models and sea-level rise

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Nov. 28, 2018

The Future of Earth System Modeling: Polar Climates. Caltech

Please note: I work for Penn State University, and help UN IPCC, NRC, etc., But I am not representing them, just me.

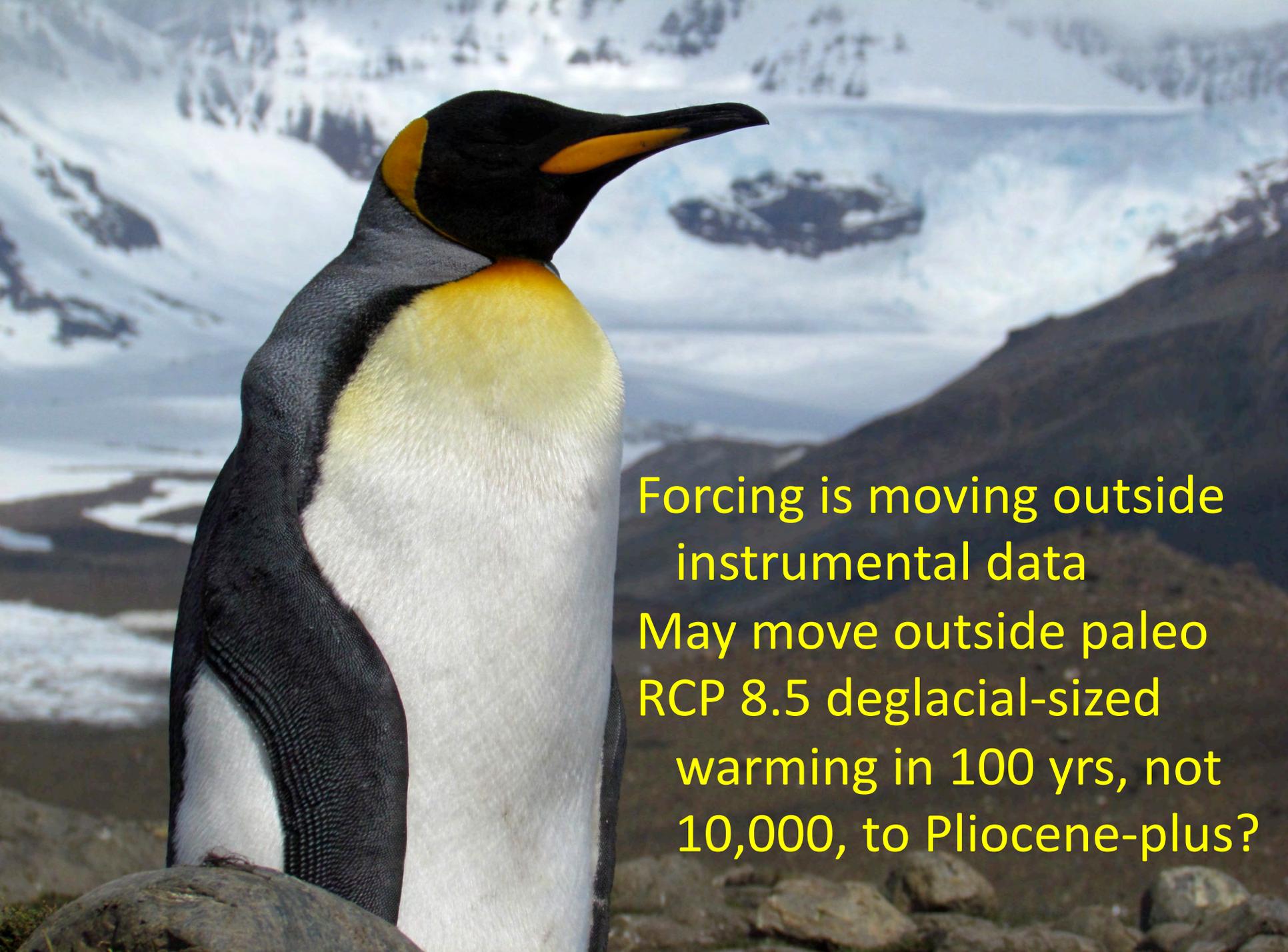
G. Comer
Foundation



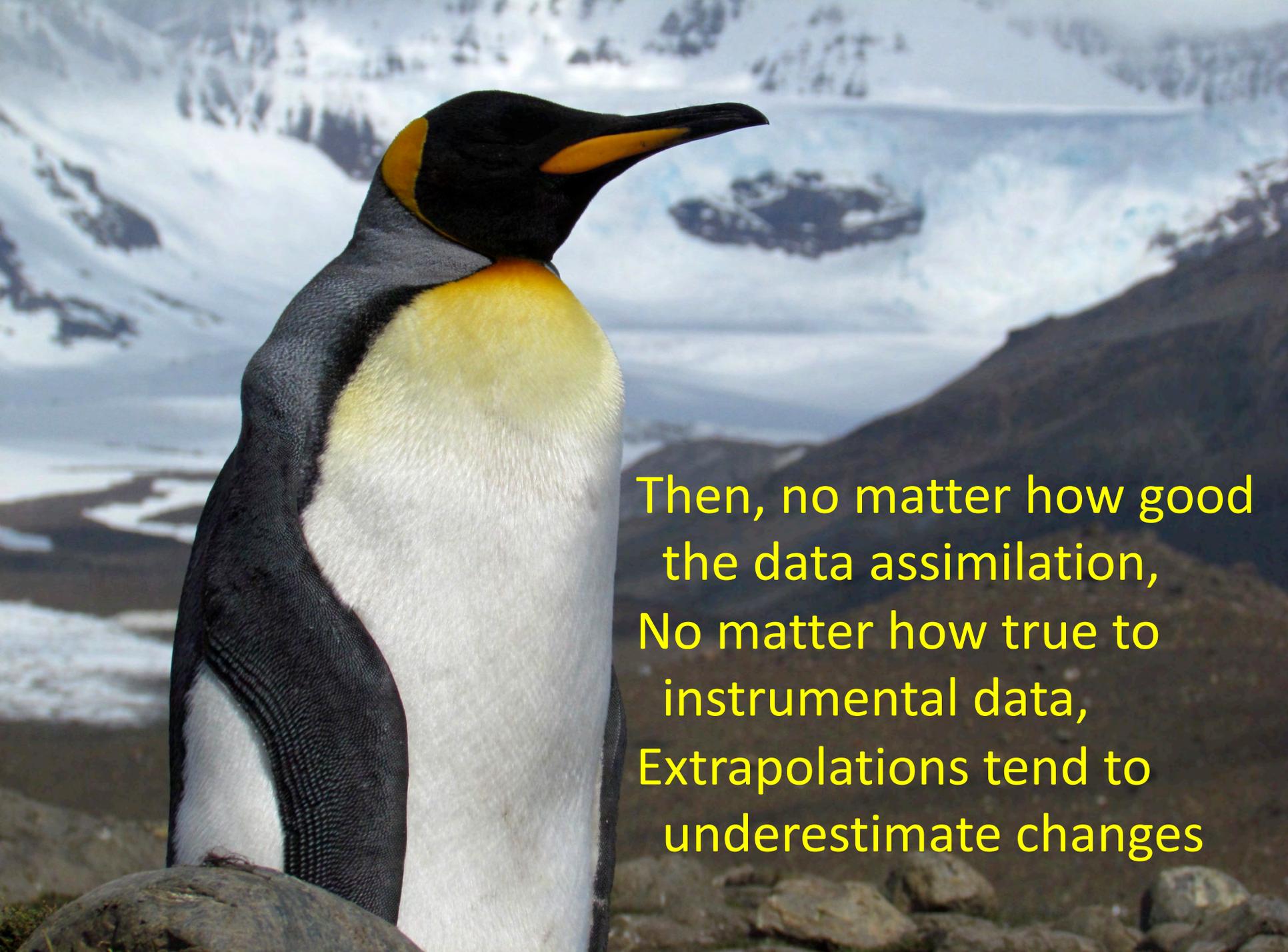
CRESIS

PENNSTATE





Forcing is moving outside
instrumental data
May move outside paleo
RCP 8.5 deglacial-sized
warming in 100 yrs, not
10,000, to Pliocene-plus?



Then, no matter how good
the data assimilation,
No matter how true to
instrumental data,
Extrapolations tend to
underestimate changes

Fig. 1, Shoji, H. & A. Higashi, 1978, A deformation mechanism map of ice, J.Glac. 21(85), 419-427

Deformation maps

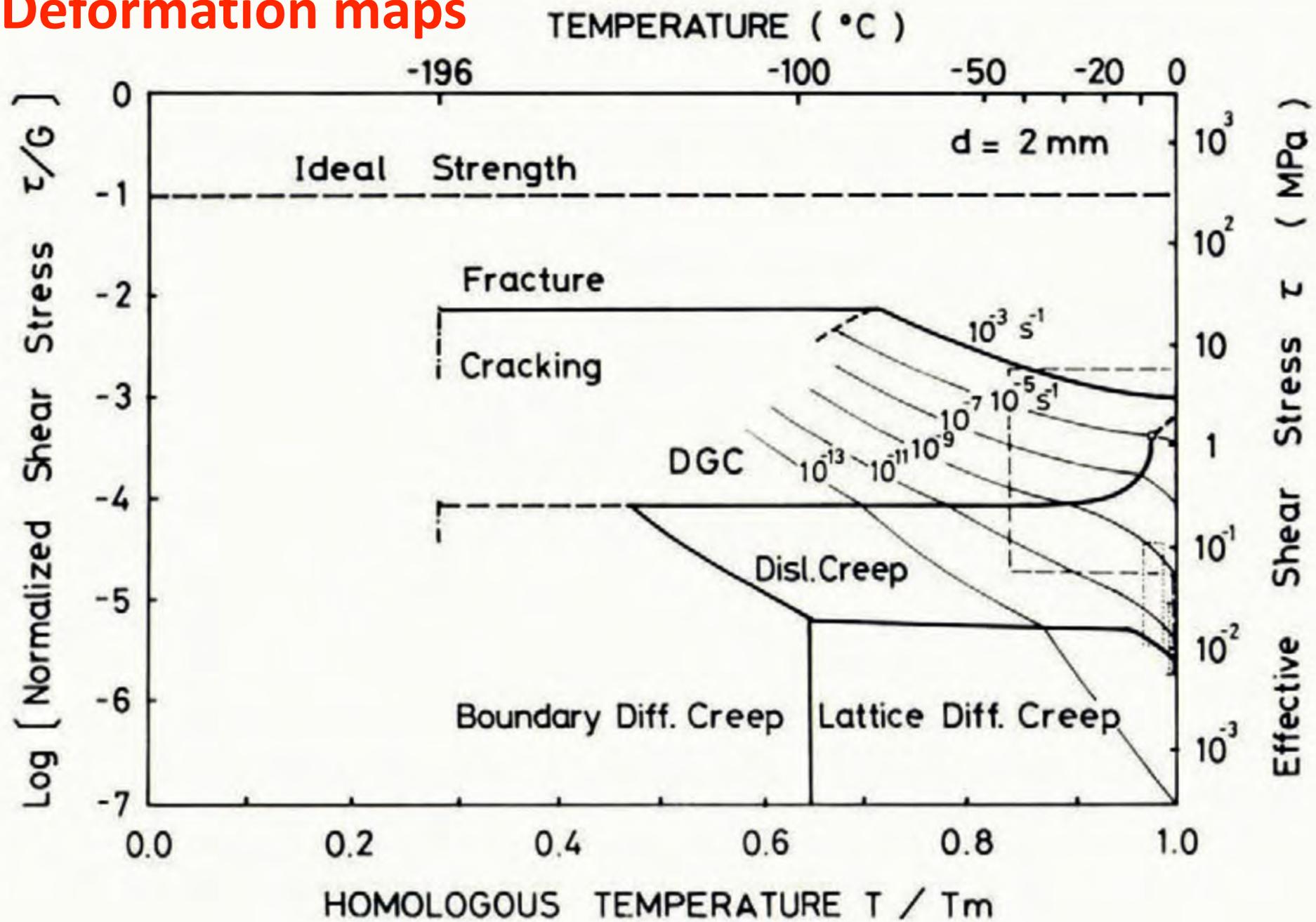
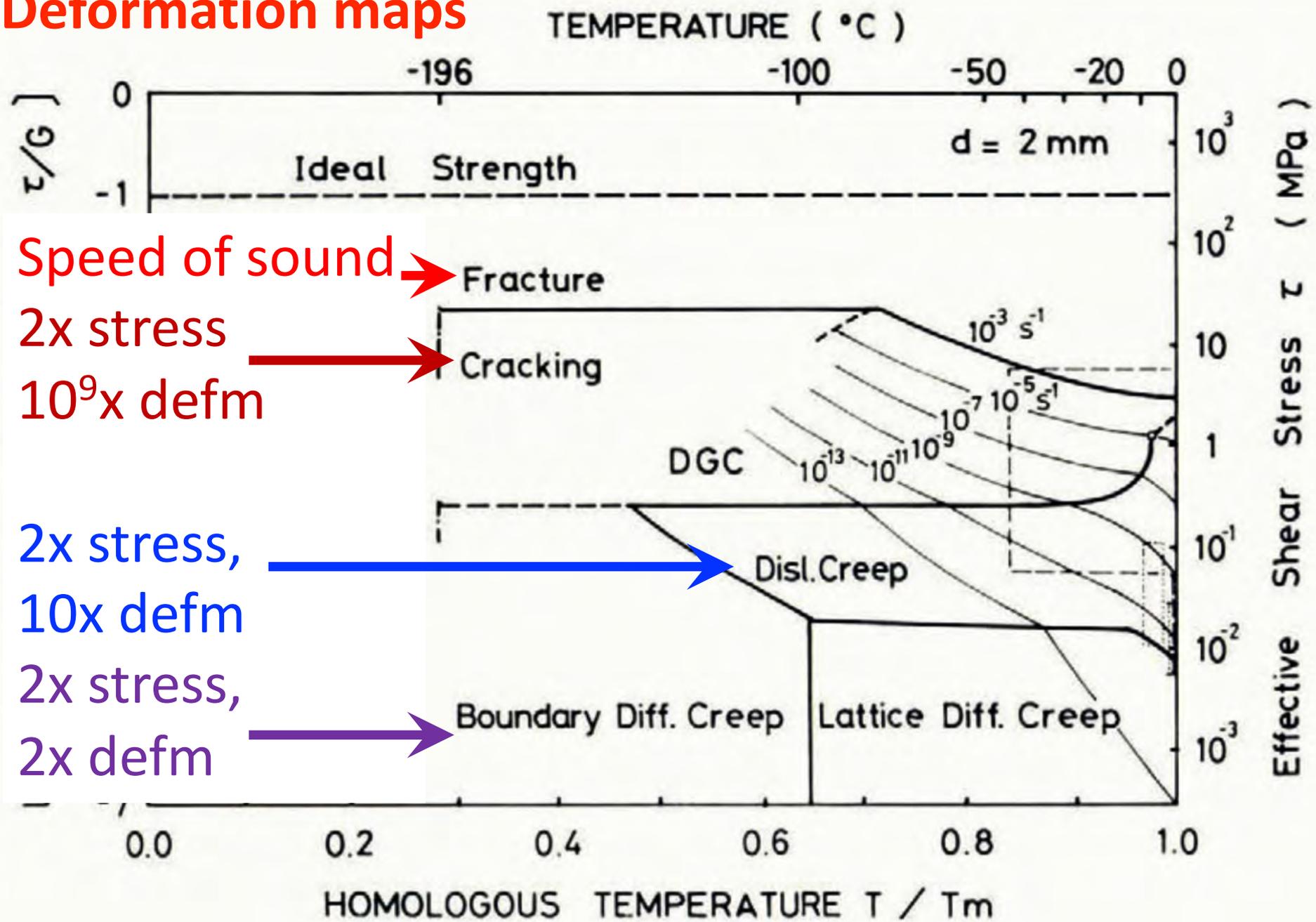


Fig. 1, Shoji, H. & A. Higashi, 1978, A deformation mechanism map of ice, J.Glac. 21(85), 419-427

Deformation maps



Speed of sound →

2x stress →

10⁹x defm →

2x stress, →

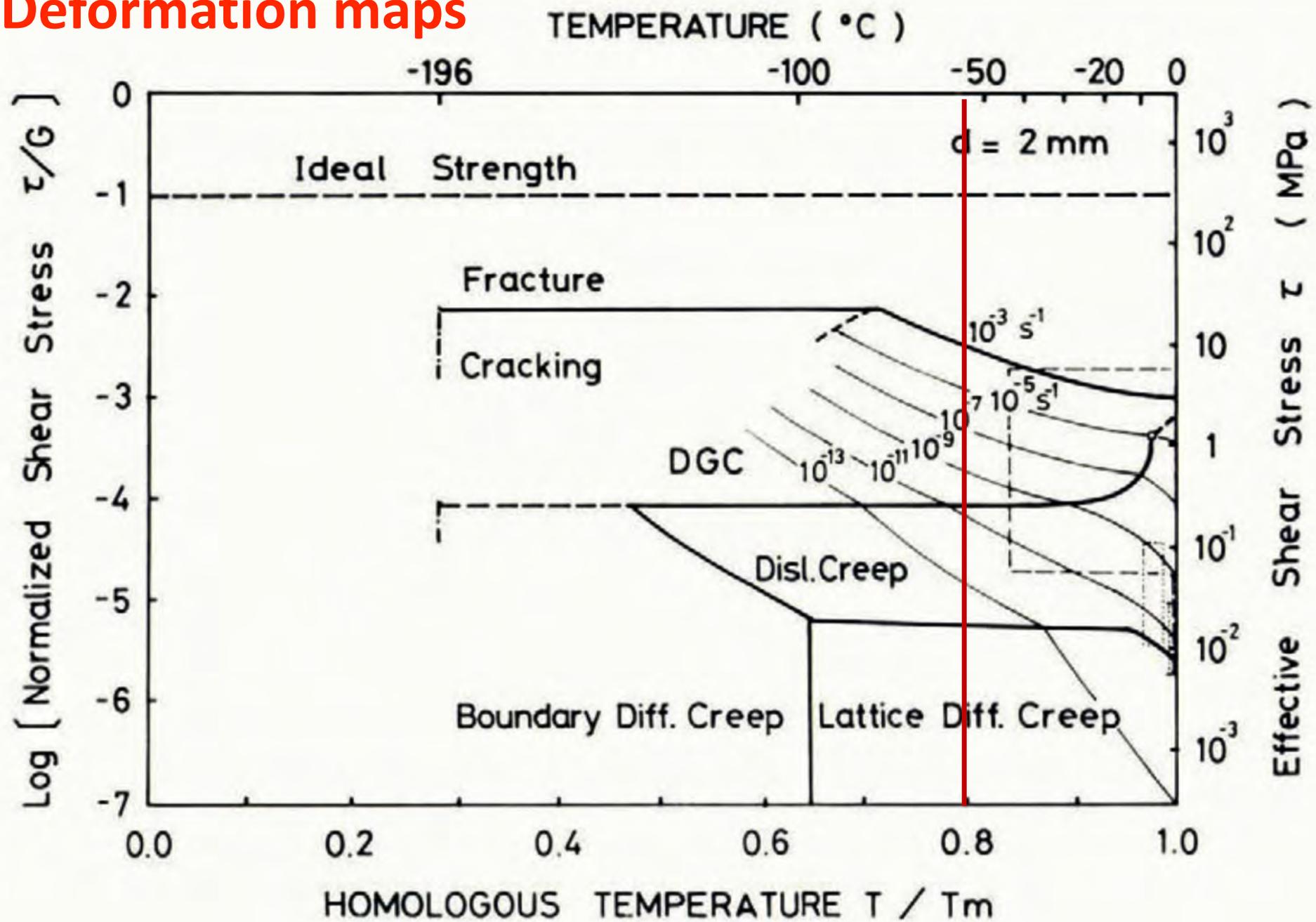
10x defm →

2x stress, →

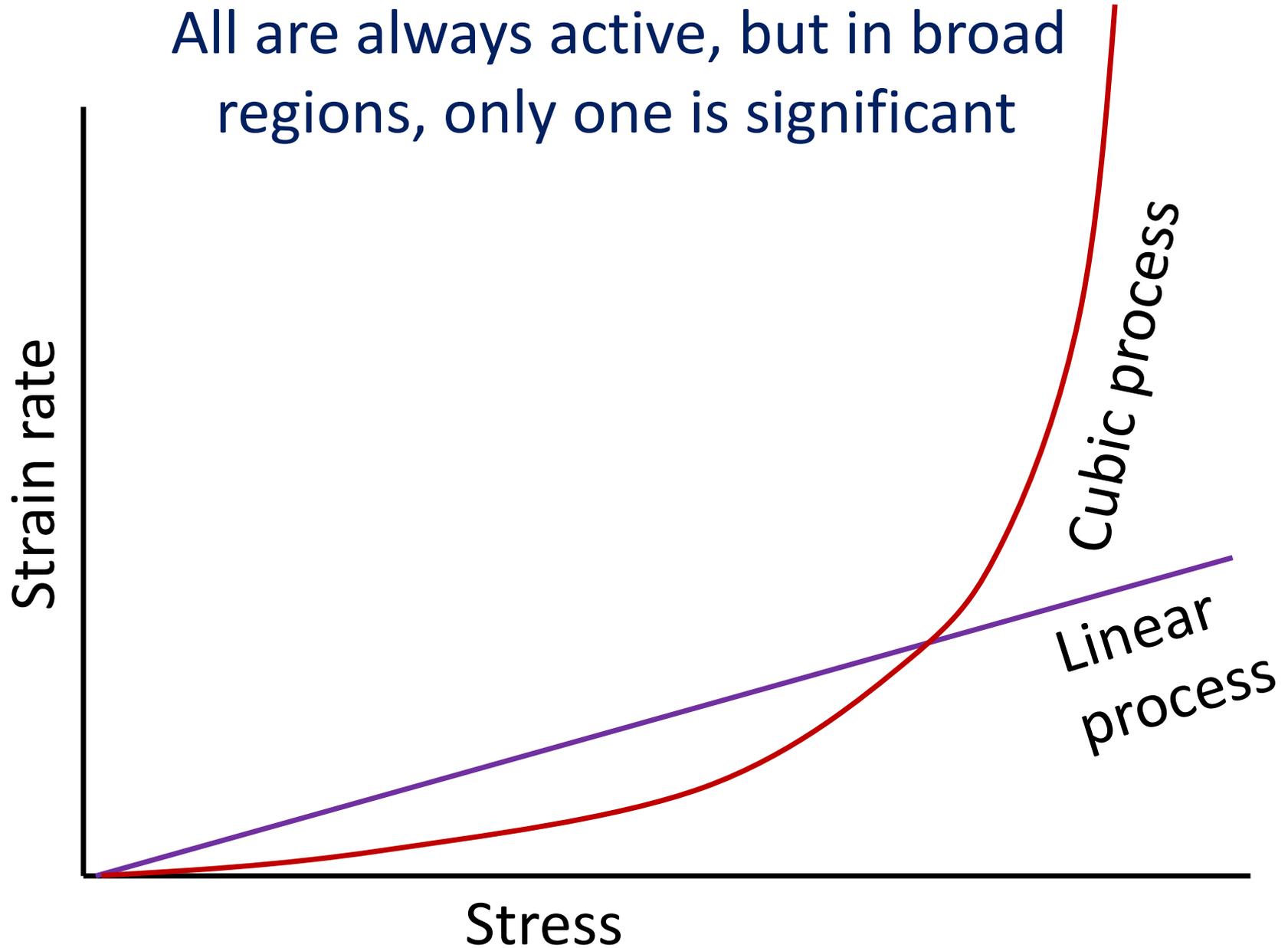
2x defm →

Fig. 1, Shoji, H. & A. Higashi, 1978, A deformation mechanism map of ice, J.Glac. 21(85), 419-427

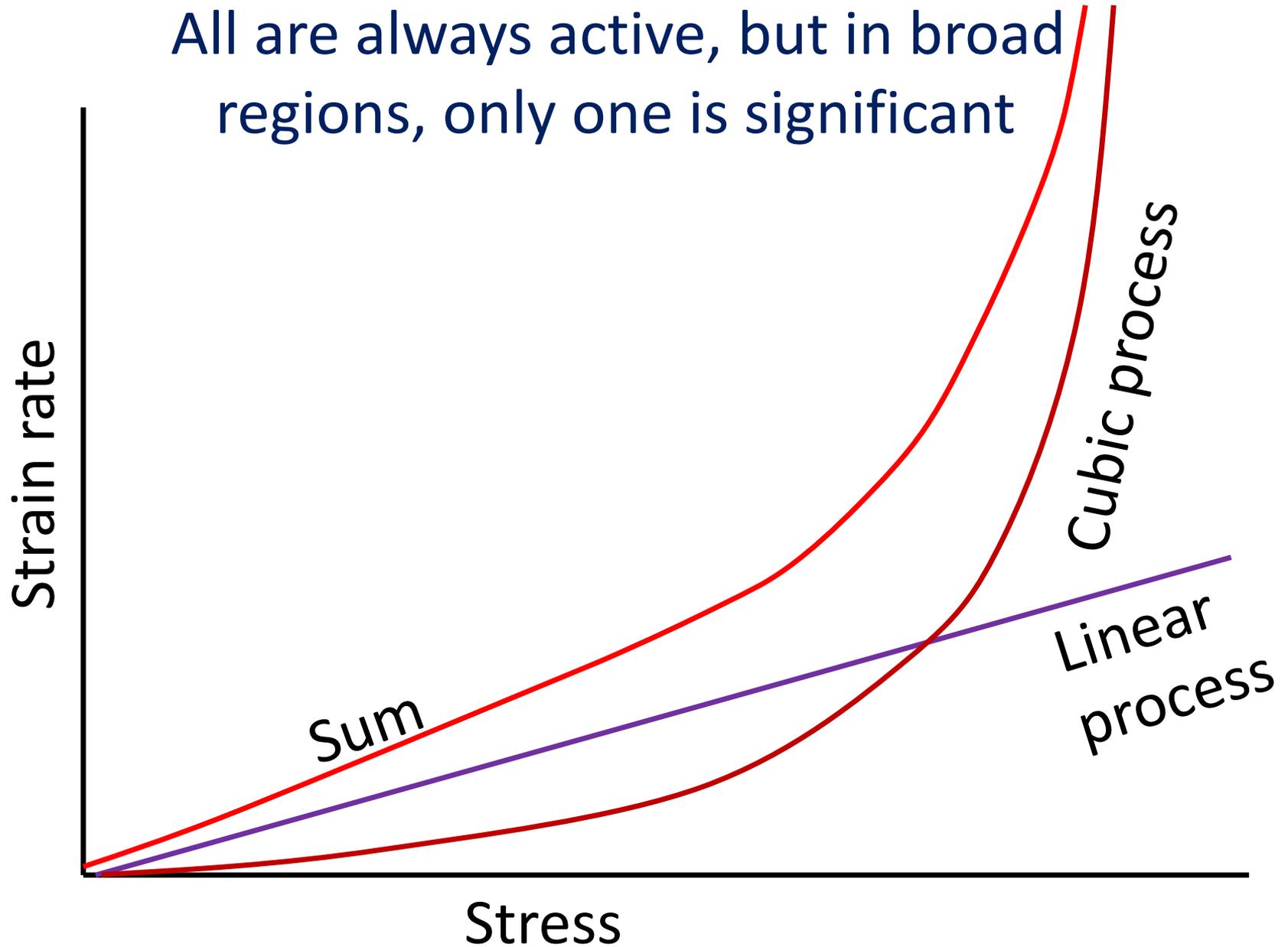
Deformation maps



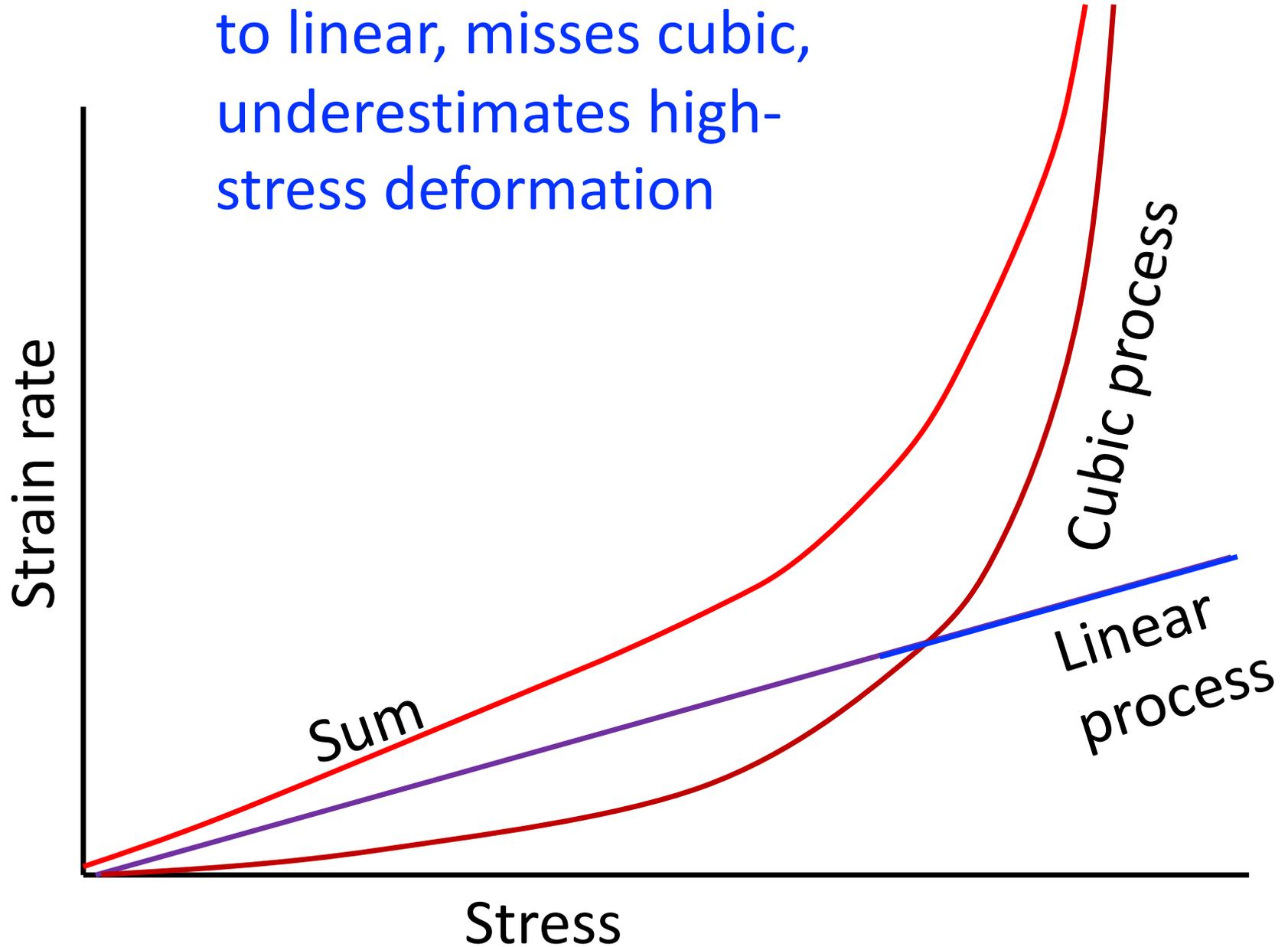
Cartoon of deformation mechanisms
All are always active, but in broad
regions, only one is significant



Cartoon of deformation mechanisms
All are always active, but in broad
regions, only one is significant

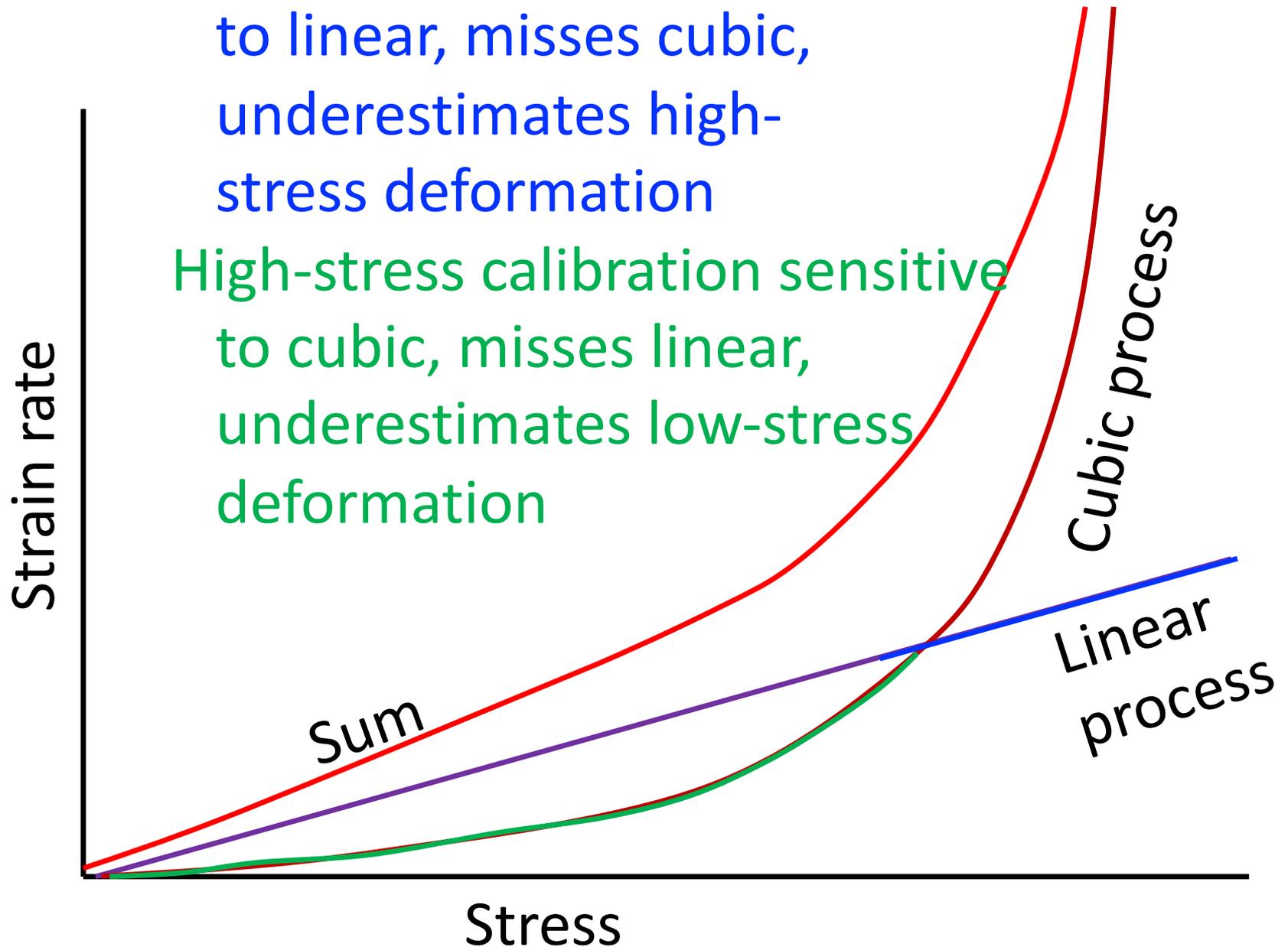


Low-stress calibration sensitive
to linear, misses cubic,
underestimates high-
stress deformation



Low-stress calibration sensitive
to linear, misses cubic,
underestimates high-
stress deformation

High-stress calibration sensitive
to cubic, misses linear,
underestimates low-stress
deformation

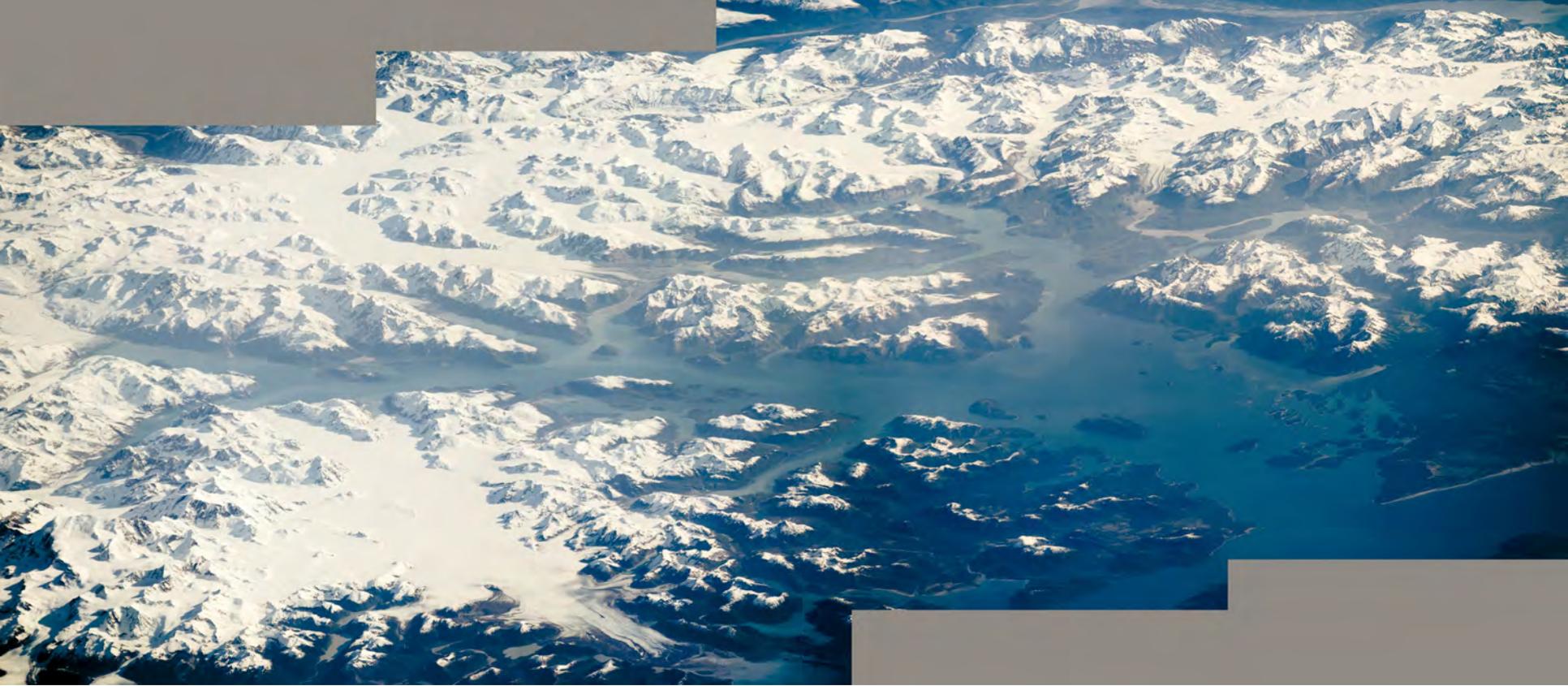


- Many deformation mechanisms
- Rates $\sim A_0 \exp[-Q/(RT)] \tau^n$
- $n=1$ (several processes) to ~ 30
(subcritical crack growth), wide range of Q
- For some (T, τ) one process dominates
- Data are sensitive to only that process
- If extrapolated too far, some other process always grows to dominate
- And model always underestimates rate

- Extrapolations underestimate rate
- Very clear for ice deformation
- Especially at onset of fracture
- Probably for basal sliding as well
- I'm happy to discuss (Iken limit, till deformation viscous-plastic, role of ice deformation in sliding; R-channels unclear but likely don't change answer)



If you want buildings,
bridges and ice sheets
to persist, then rule is:
**All extrapolations are
overly optimistic**



<https://earthobservatory.nasa.gov/images/88701/glacier-bay-national-park-preserve>

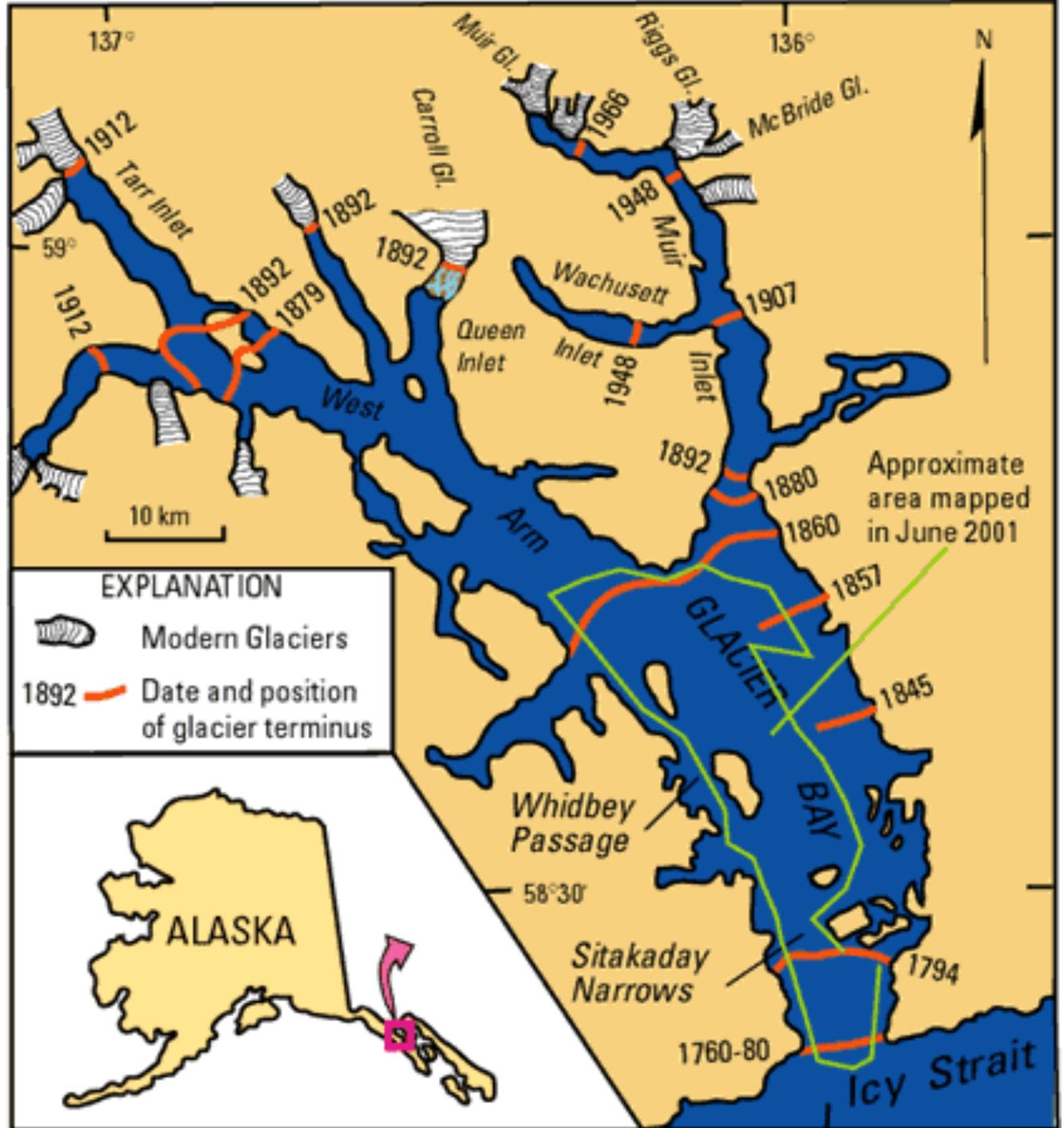


Glacier Bay

Icy Strait

<https://earthobservatory.nasa.gov/images/88701/glacier-bay-national-park-preserve>

<https://sound.waves.usgs.gov/2001/07/glacierbaymap.gif>



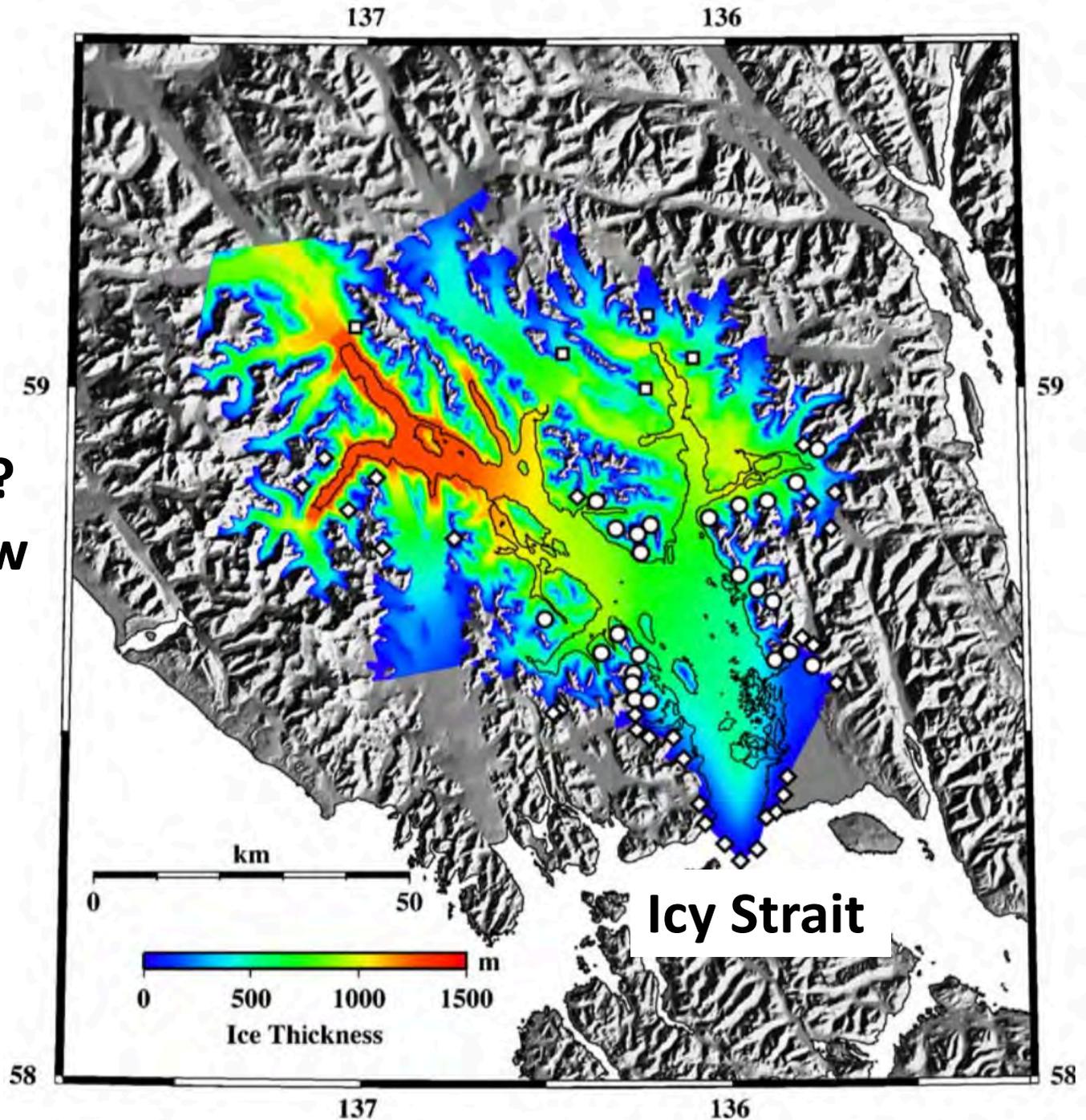


Muir Glacier, Alaska, August 13, 1941, photo by W.O. Field



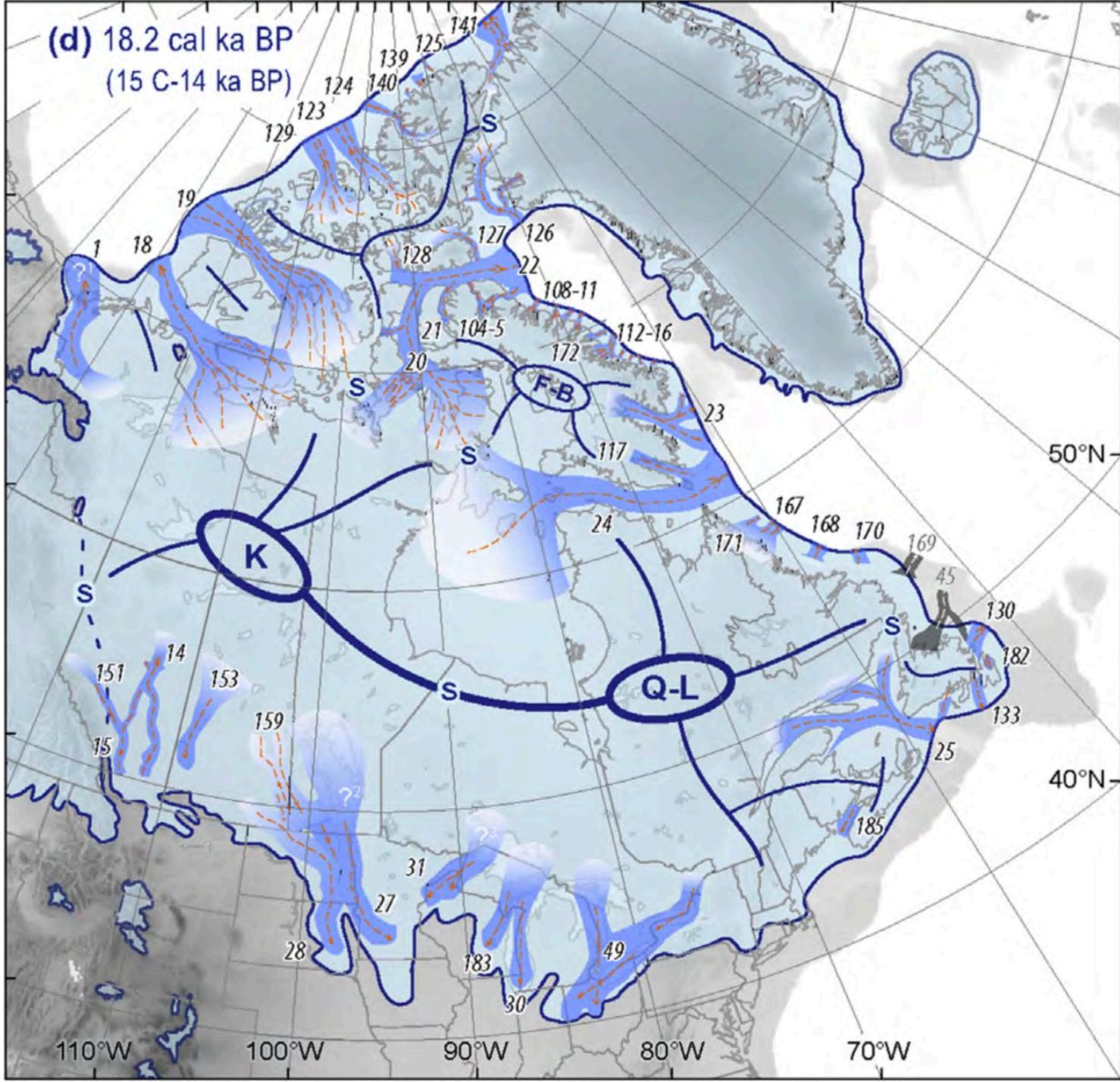
Muir Glacier, Alaska, August 31, 2004, photo by B.F. Molnia

**Δ Ice thickness
~1500 m max
NOT by flow
to Icy Strait
105 km retreat
Peak ~1 km/yr?
Shallow, narrow
compared to
polar outlets**

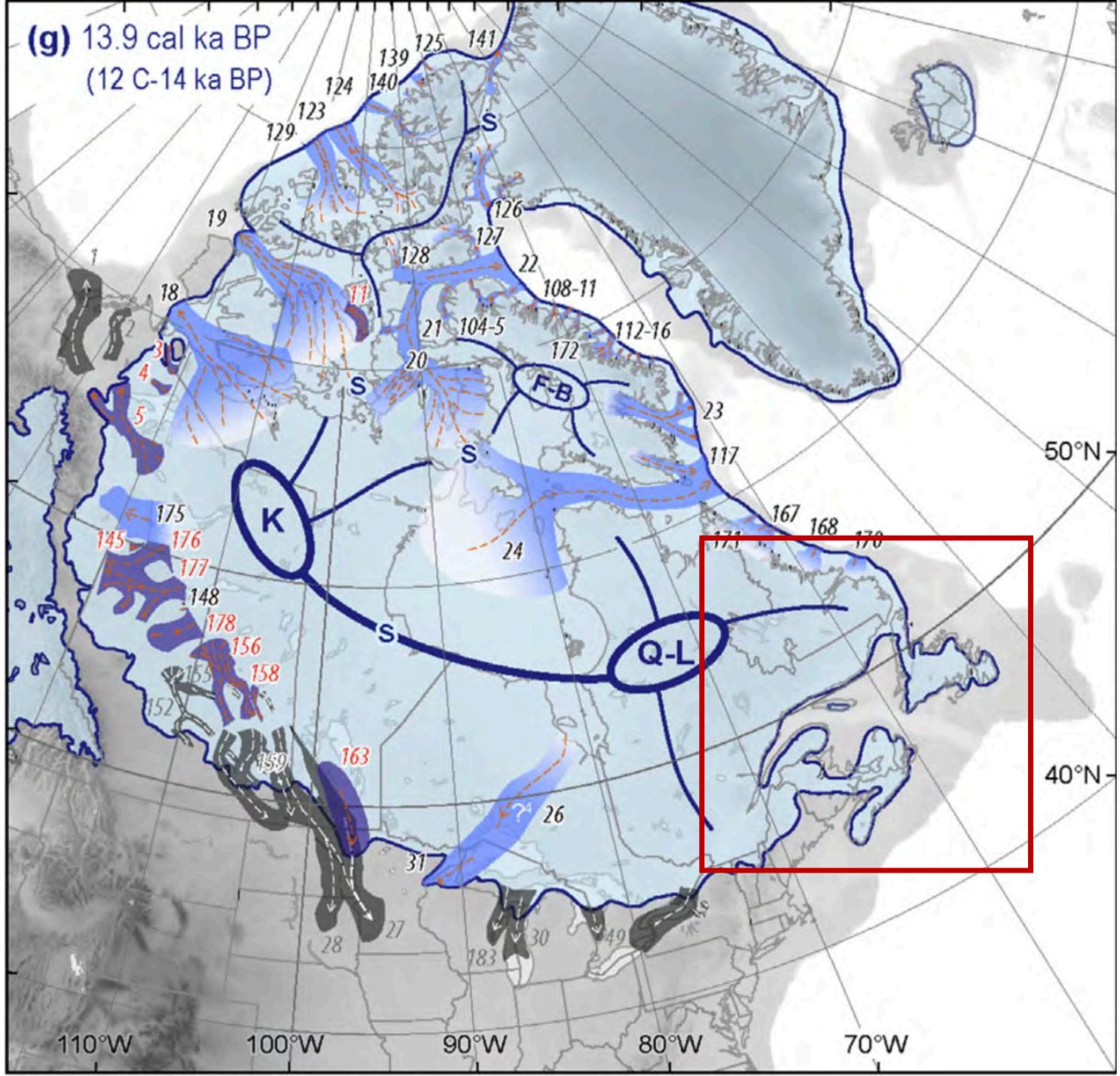


C.F. Larsen et al. / Earth
and Planetary Science
Letters 237 (2005) 548–
560

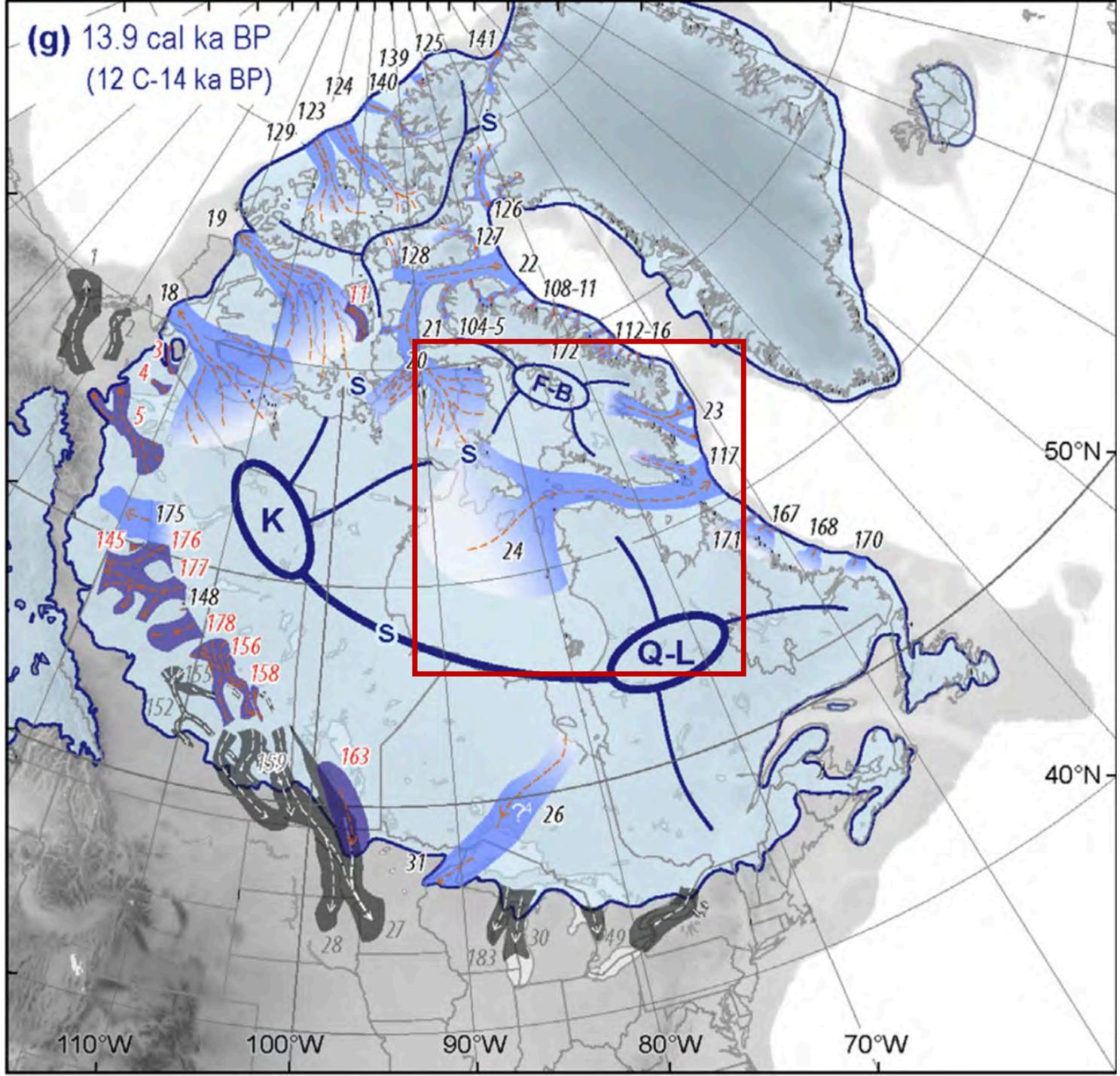
Margold,
Stokes &
Clark,
QSR,
2018



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2018



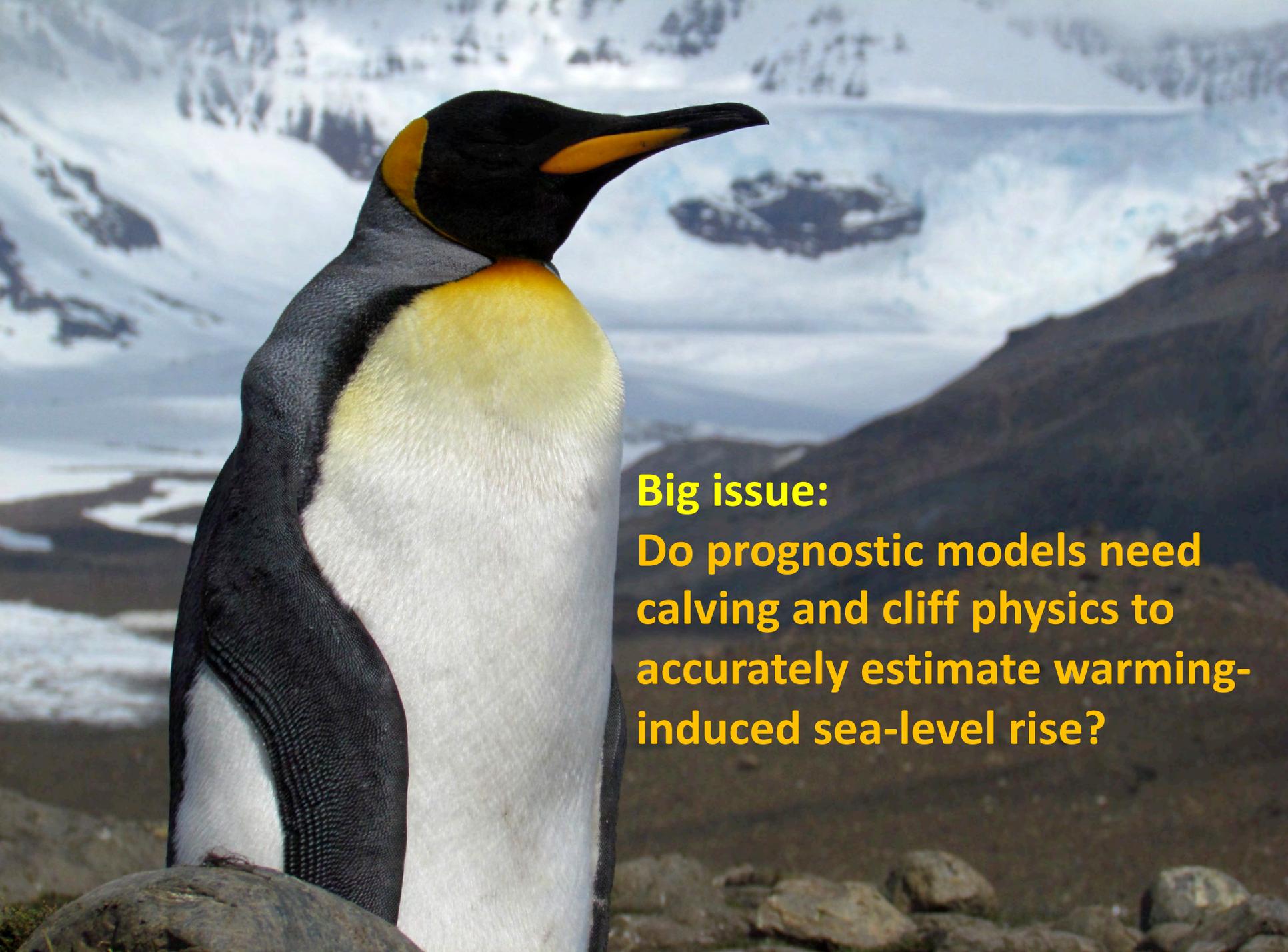
Margold,
Stokes &
Clark,
QSR,
2018



(k) 8.9 cal ka BP
(8 C-14 ka BP)

Margold,
Stokes &
Clark,
QSR,
2018





Big issue:
Do prognostic models need calving and cliff physics to accurately estimate warming-induced sea-level rise?

There are many places on land where you can walk up on the toe of a glacier...



which is what these Greenlandic wild reindeer did, to get away from tundra mosquitoes



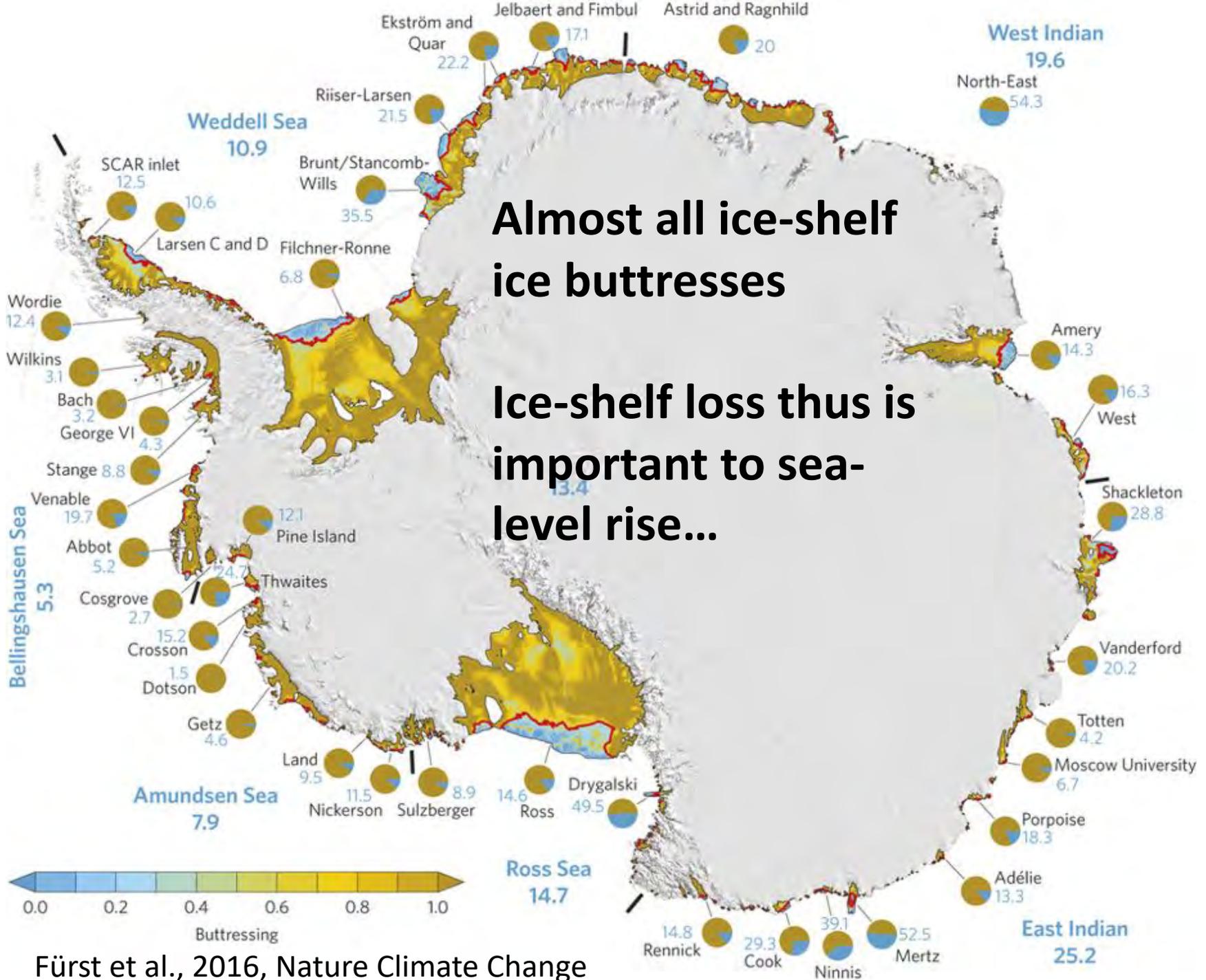
But you can't paddle your sea kayak to the ice front, jump out and walk up—almost all ice flowing into the ocean ends in a cliff, where fracture dominates

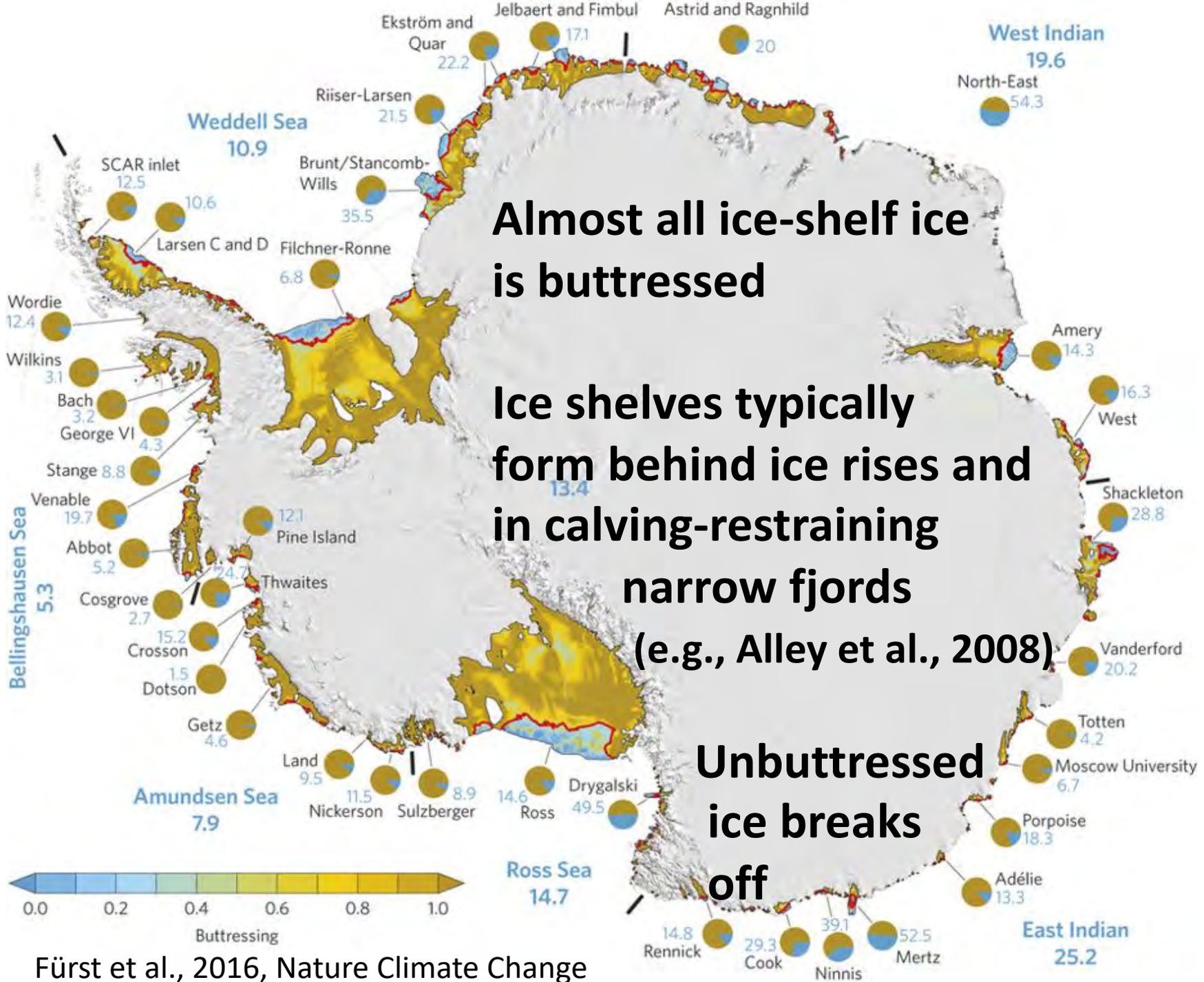
So, a reasonable hypothesis is that calving and cliff failure will continue to be important

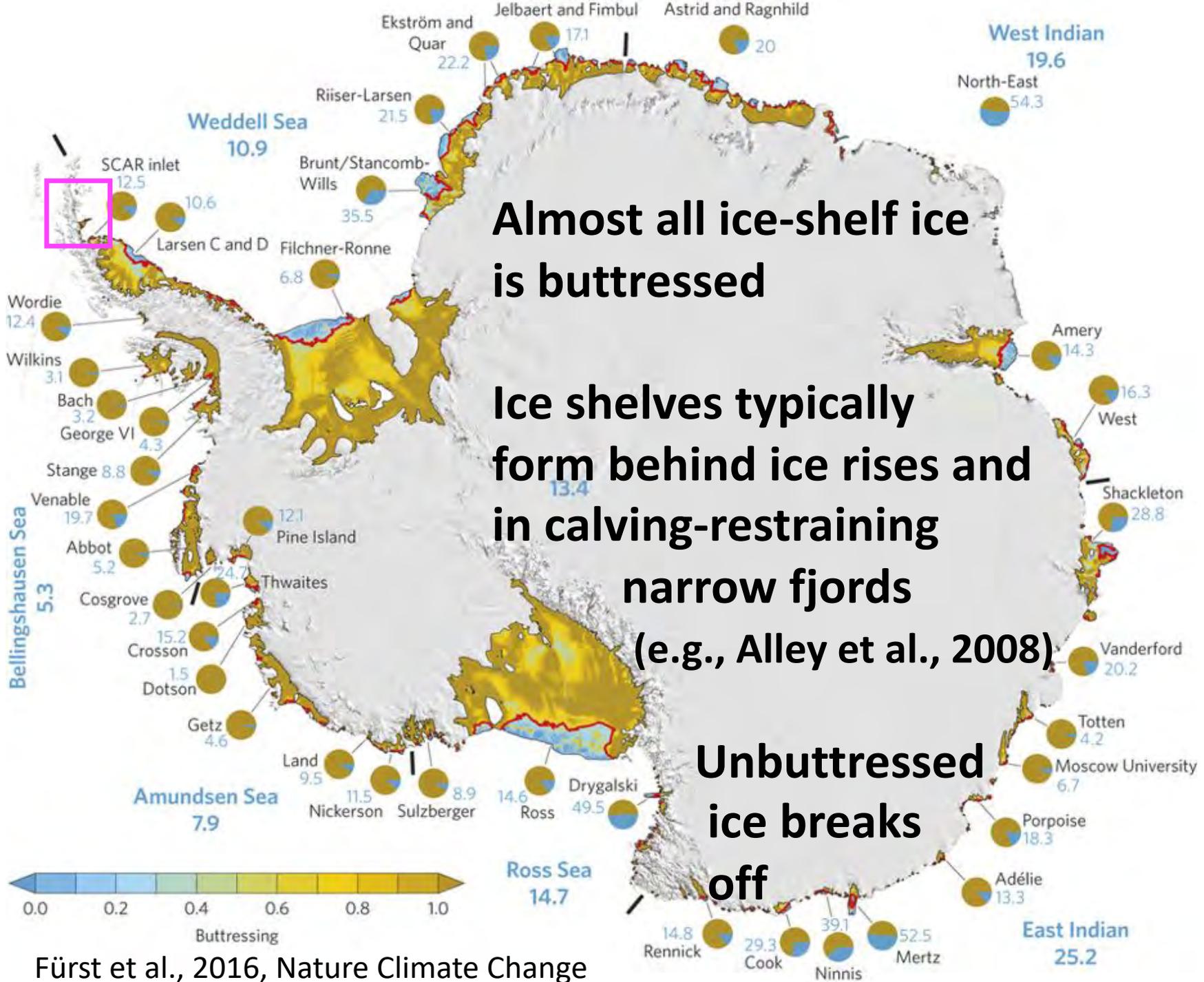


Almost all ice-shelf
ice buttresses

Ice-shelf loss thus is
important to sea-
level rise...



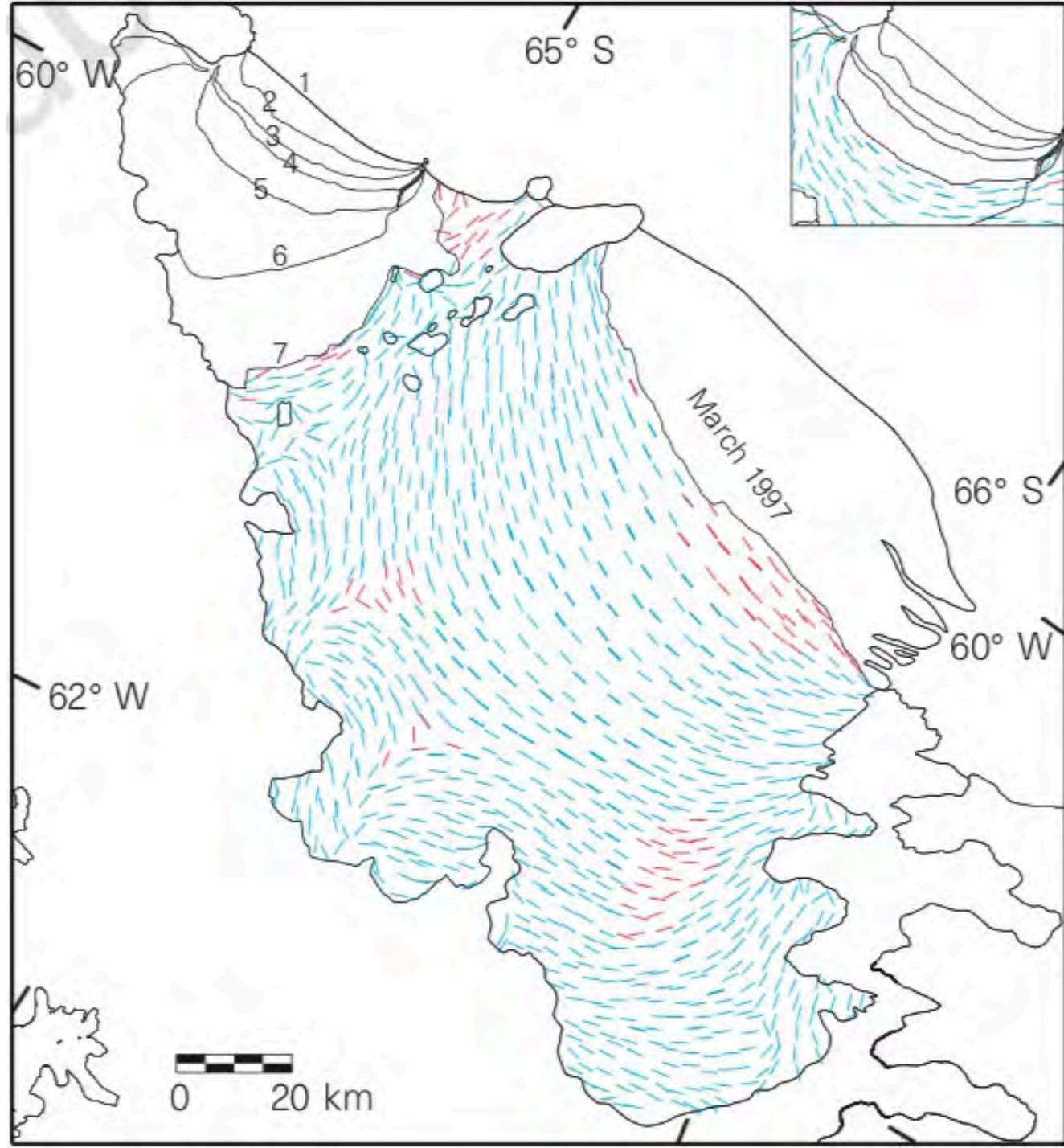




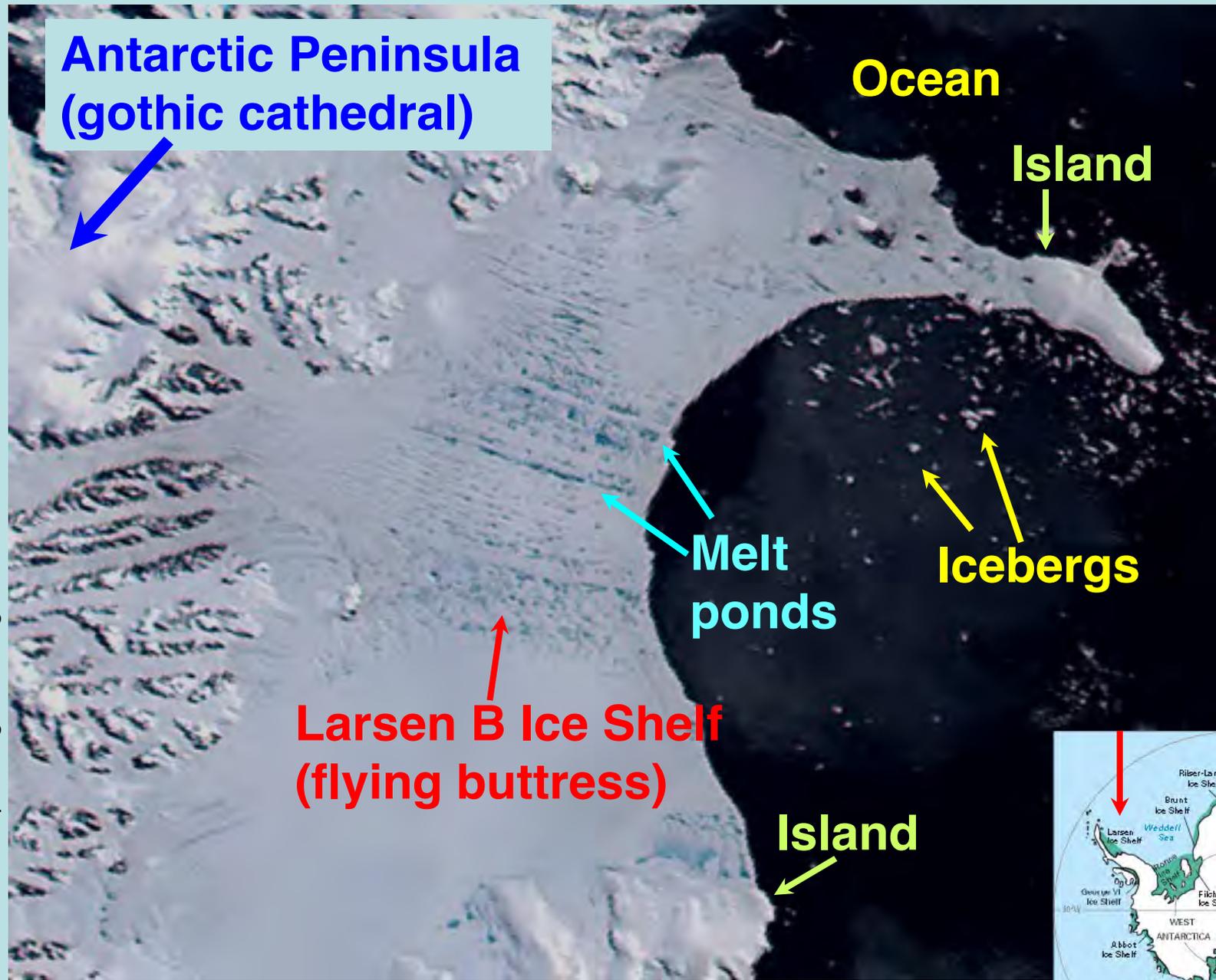
**Doake et al.,
Nature, 1998**

**Successfully
predicted Larsen
B collapse**

**When some
buttressing lost,
ice shelves tend
to break up
completely**



<http://svs.gsfc.nasa.gov/vis/a000000/a002400/a002421/index.html>



**Antarctic Peninsula
(gothic cathedral)**

Ocean

Island

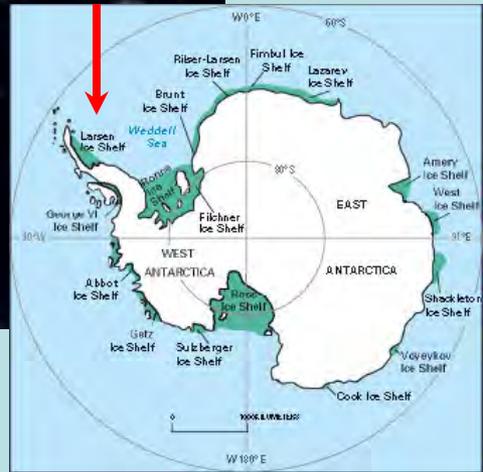
**Melt
ponds**

Icebergs

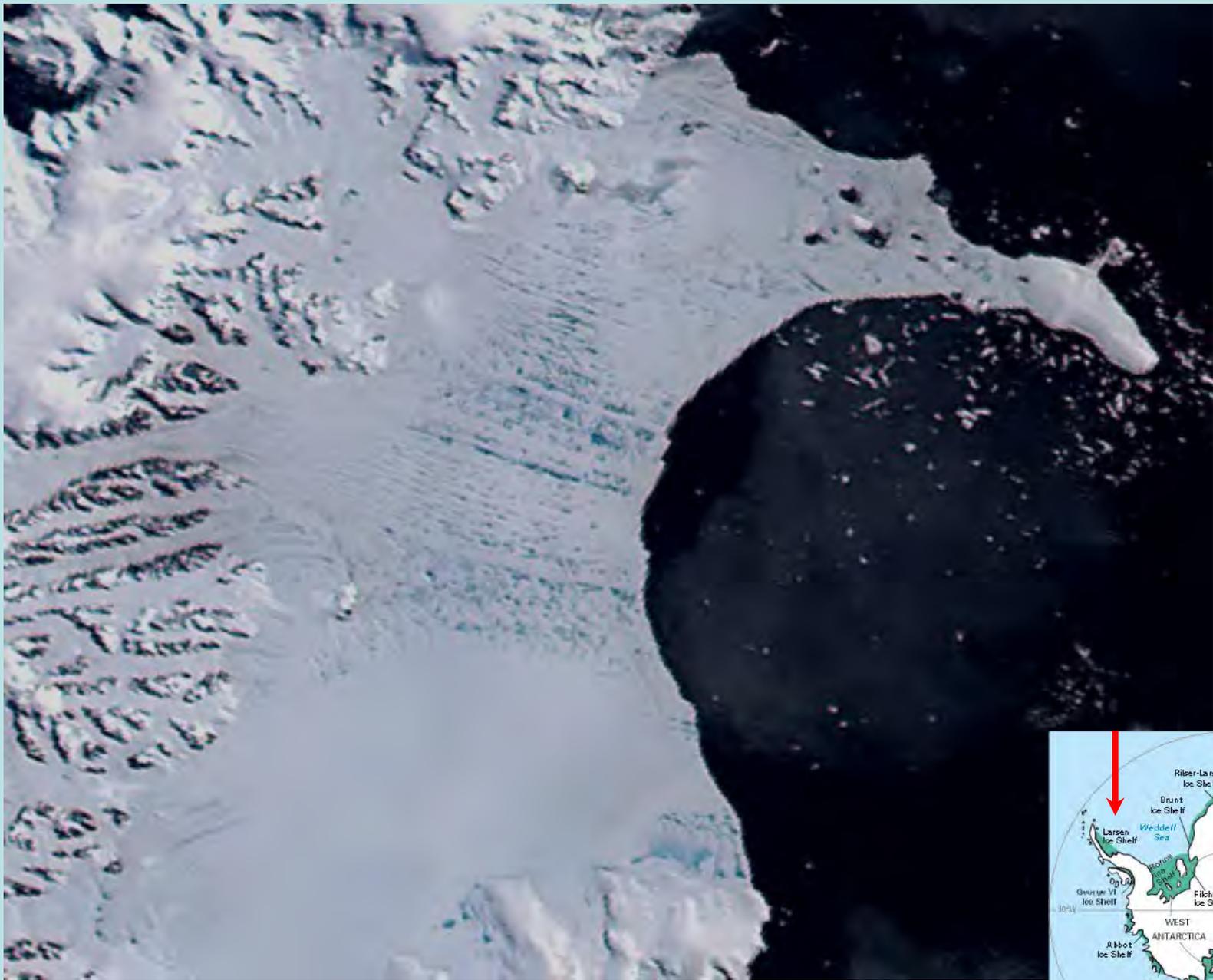
**Larsen B Ice Shelf
(flying buttress)**

Island

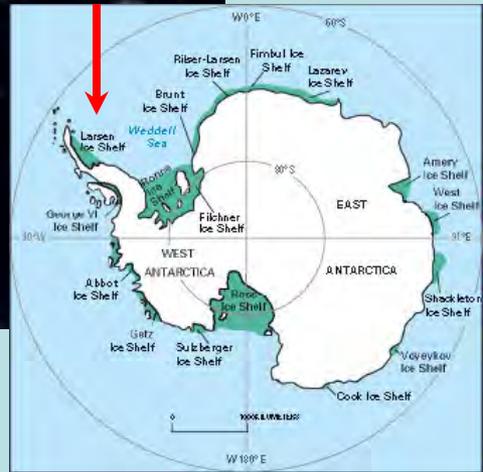
12 mi
20 km



January 31, 2002

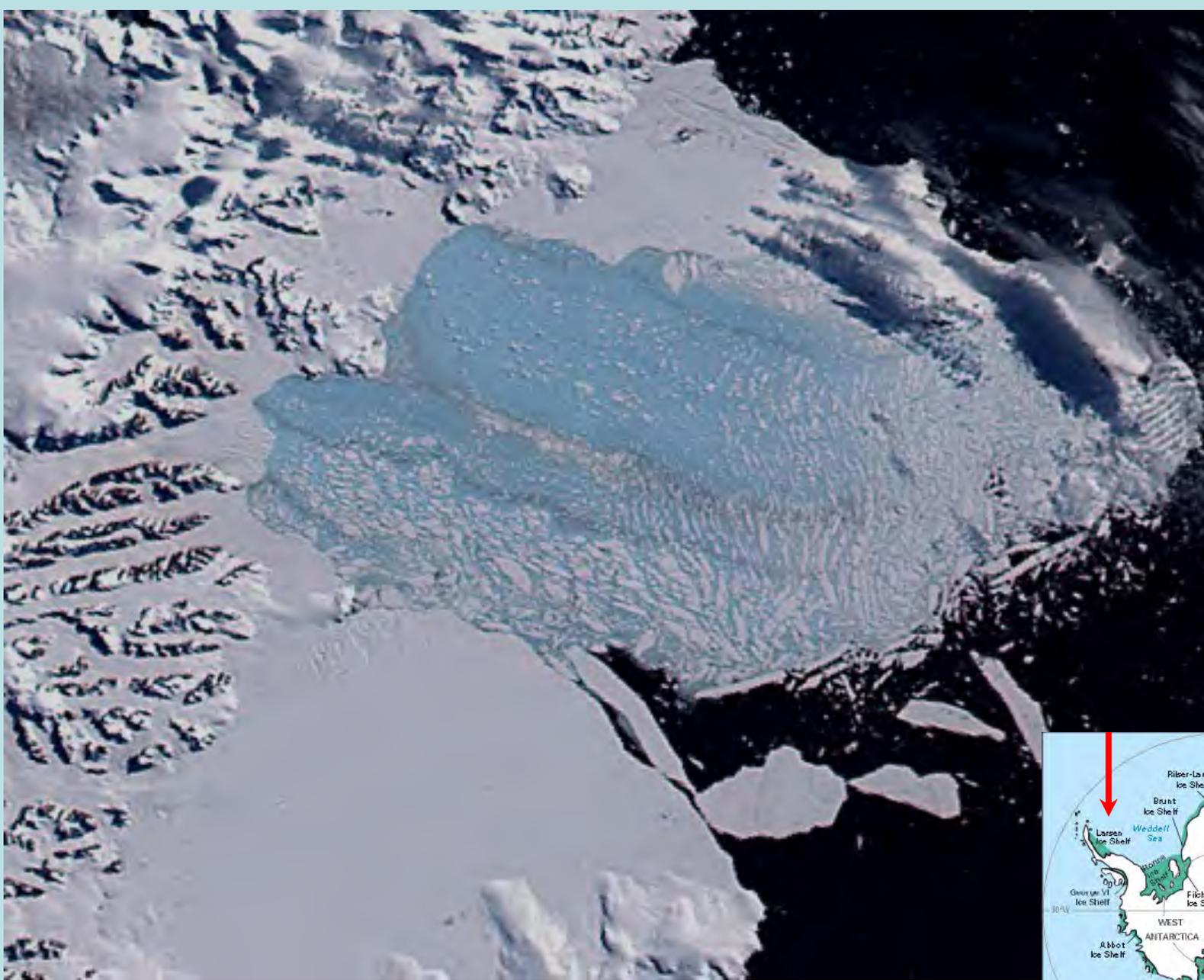


12 mi
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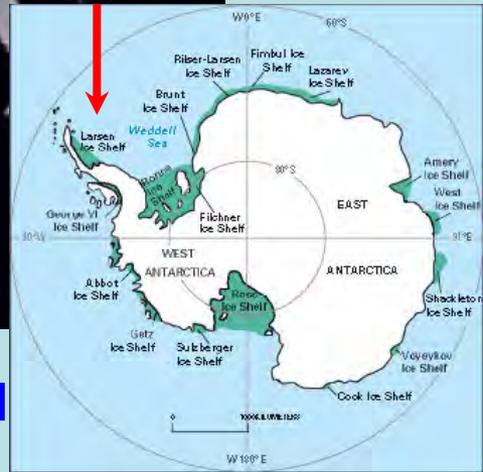


January 31, 2002

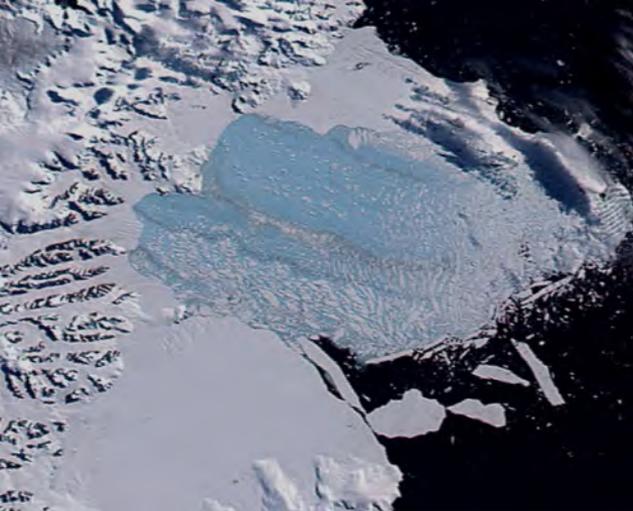
<http://svs.gsfc.nasa.gov/vis/a000000/a002400/a002421/index.html>



12 mi
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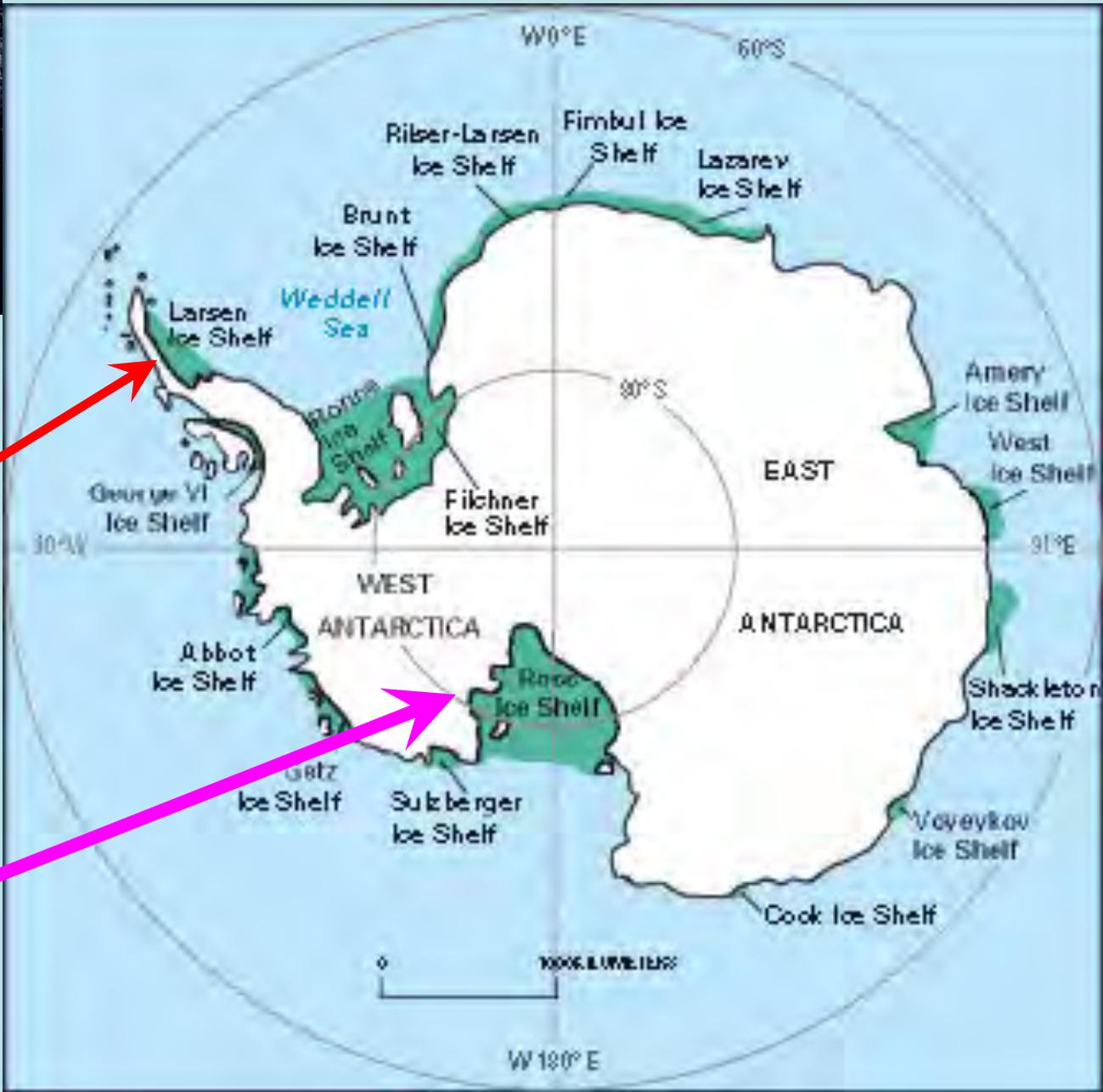


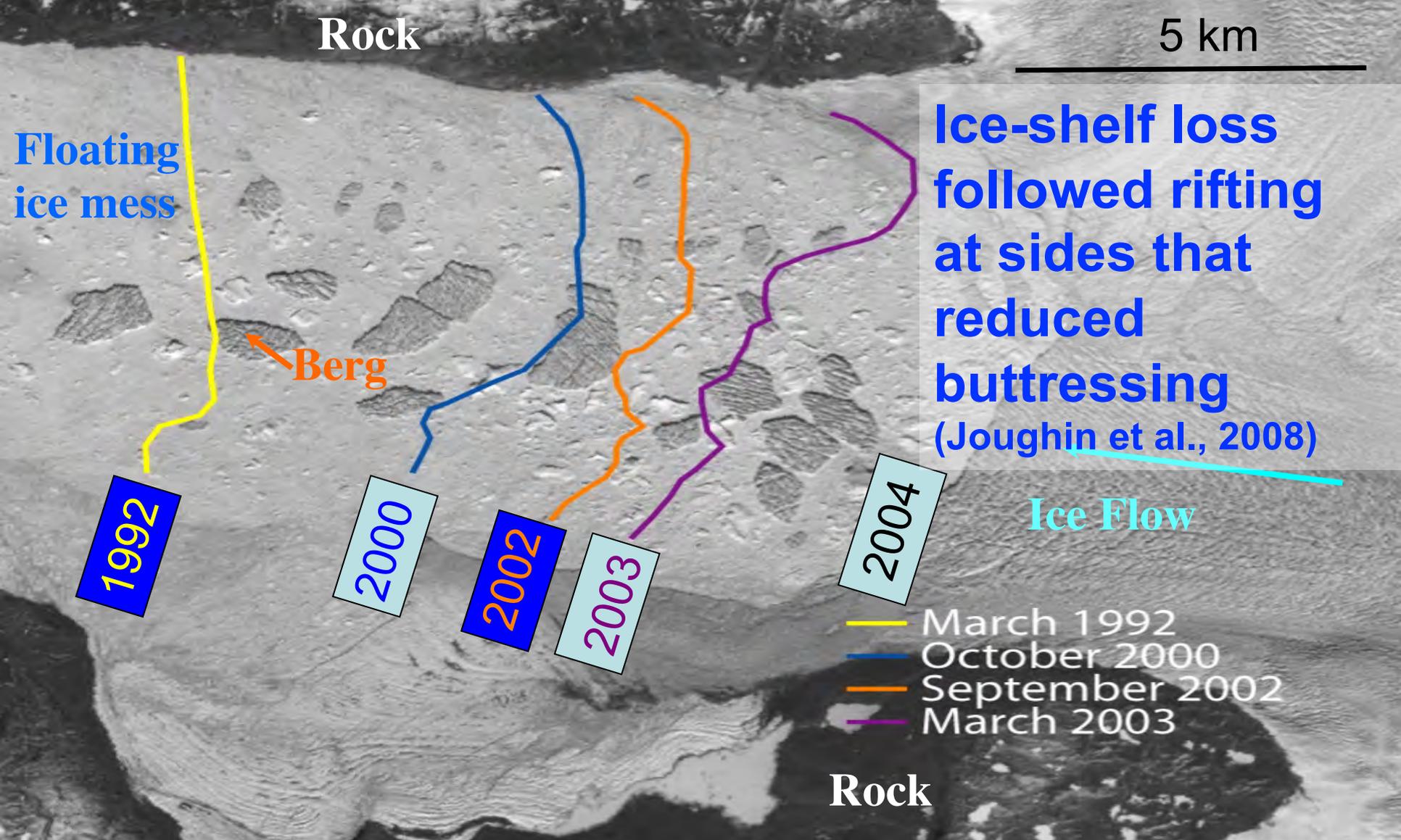
March 7, 2002. **8x tributary flow-speed increase followed**



Not much ice behind Larsen B; loss can't raise sea level much

Many more ice shelves with lots of ice behind them that can raise sea level a lot.





Jakobshavn Isbrae retreat and speed doubling from ocean-warming-induced ice-shelf loss Ian Joughin (Alley et al., 2005).

We've seen ice shelves fall apart, and not grow back

Jakobshavn, Helheim try to re-grow, favored by:

- cold ice, so strong
- much thicker than surface-crevasse depth
- moulin drainage inland keeps water out of crevasses (Parizek et al., 2010)
- narrow fjords give stabilizing side shear, restrict wave forcing, aid mélange formation
- flow so fast that can grow an ice shelf in a winter
- fluctuations in climate give colder years

But shelves still fail to re-grow

Rock

Hysteresis—breaking easier than building

So we see in multiple cases

- Warming (air and/or water) reduces shelf buttressing
- Shelf then breaks off, leaving a grounded calving cliff
- Ice flow accelerates
- And the shelf stays off

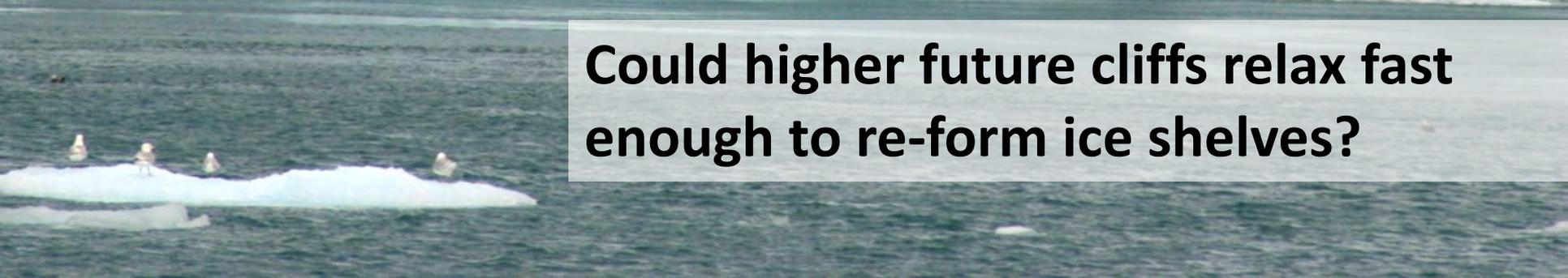




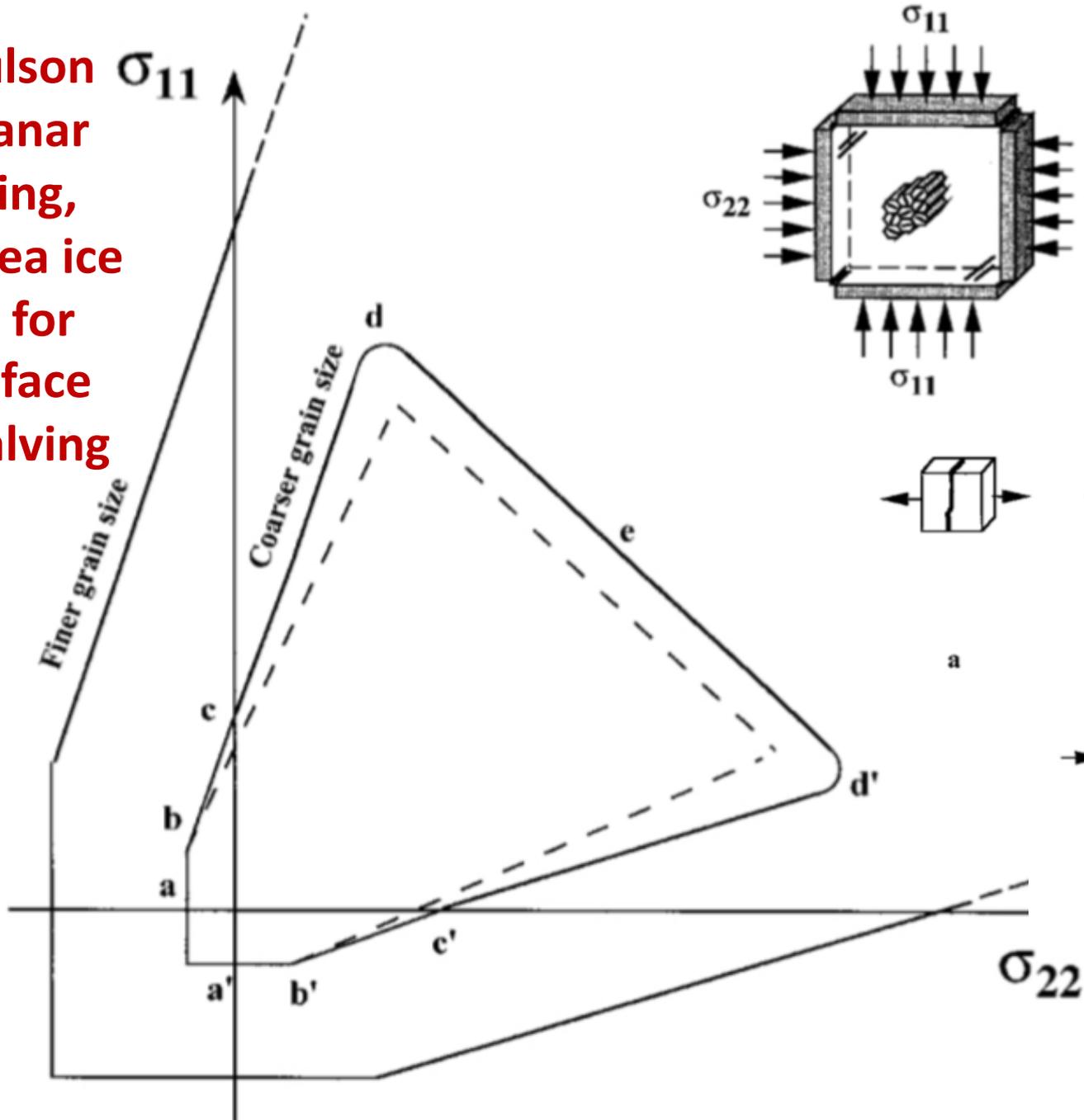
All cliff ice is relaxing by creep, but is breaking faster

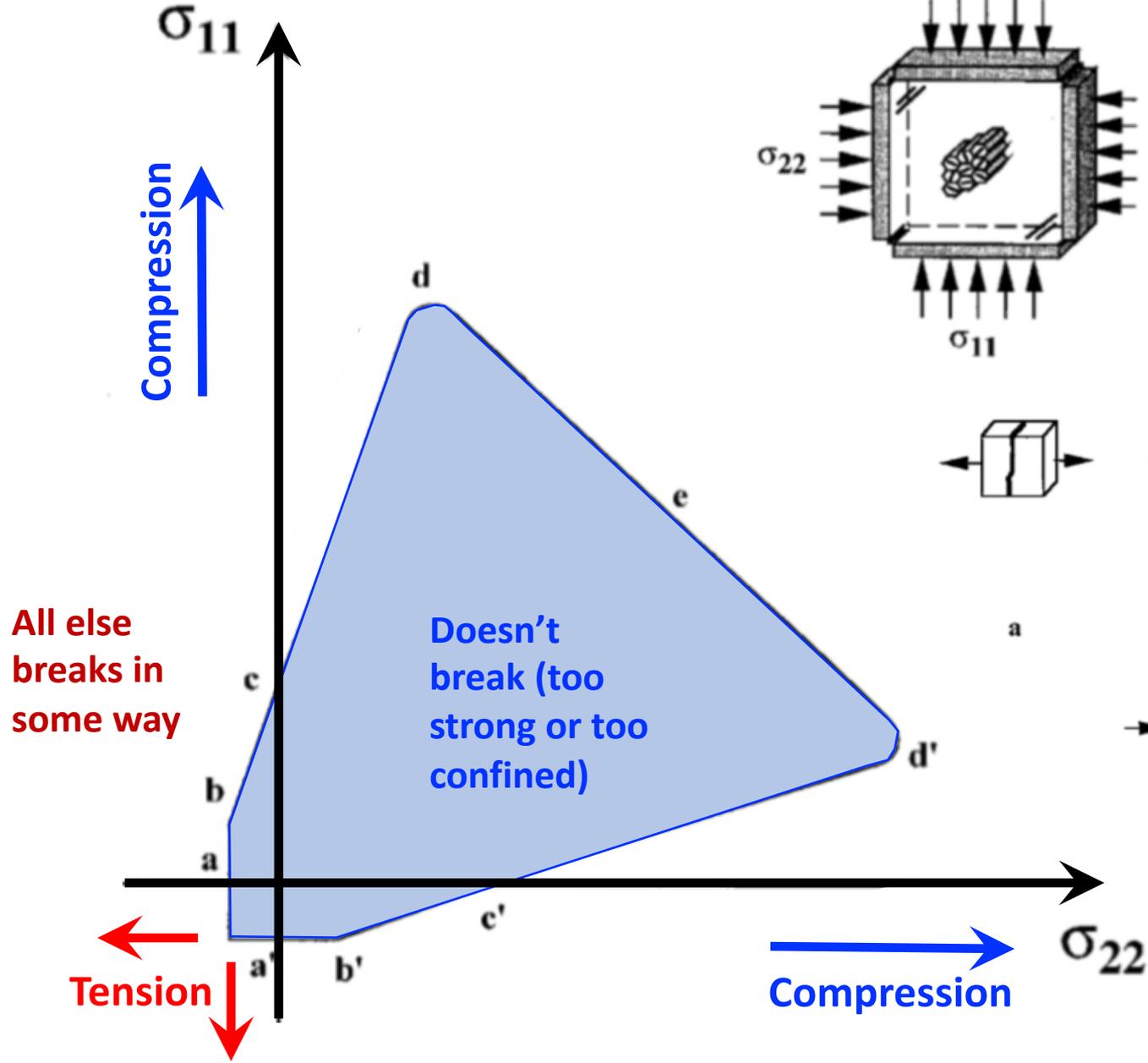


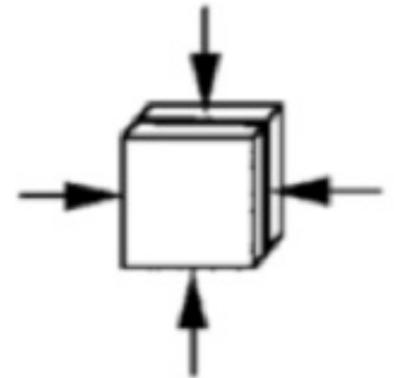
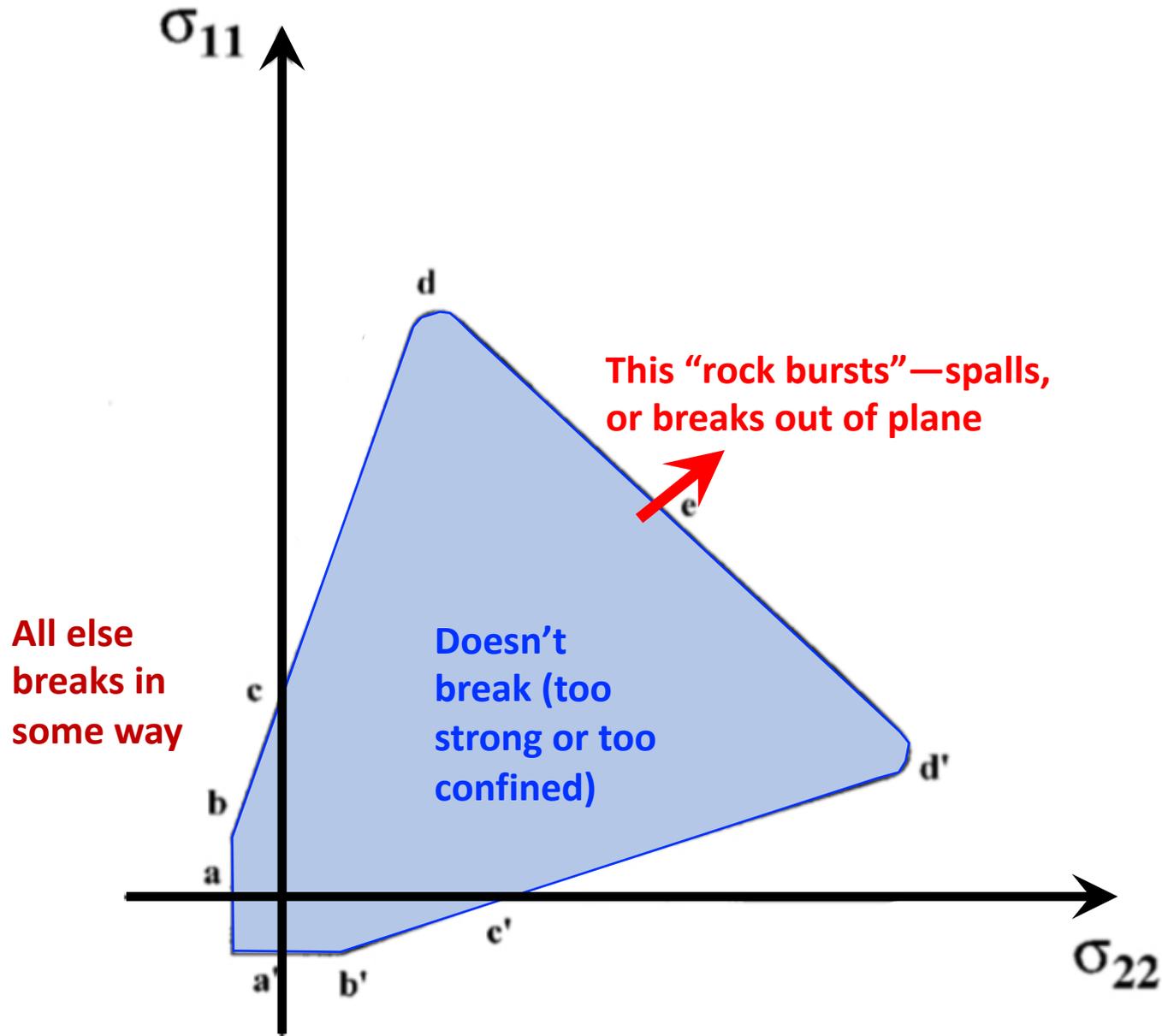
Could higher future cliffs relax fast enough to re-form ice shelves?



Schulson
 —planar
 loading,
 for sea ice
 (and for
 free face
 of calving
 cliff)

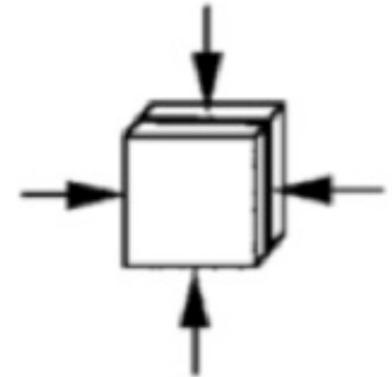
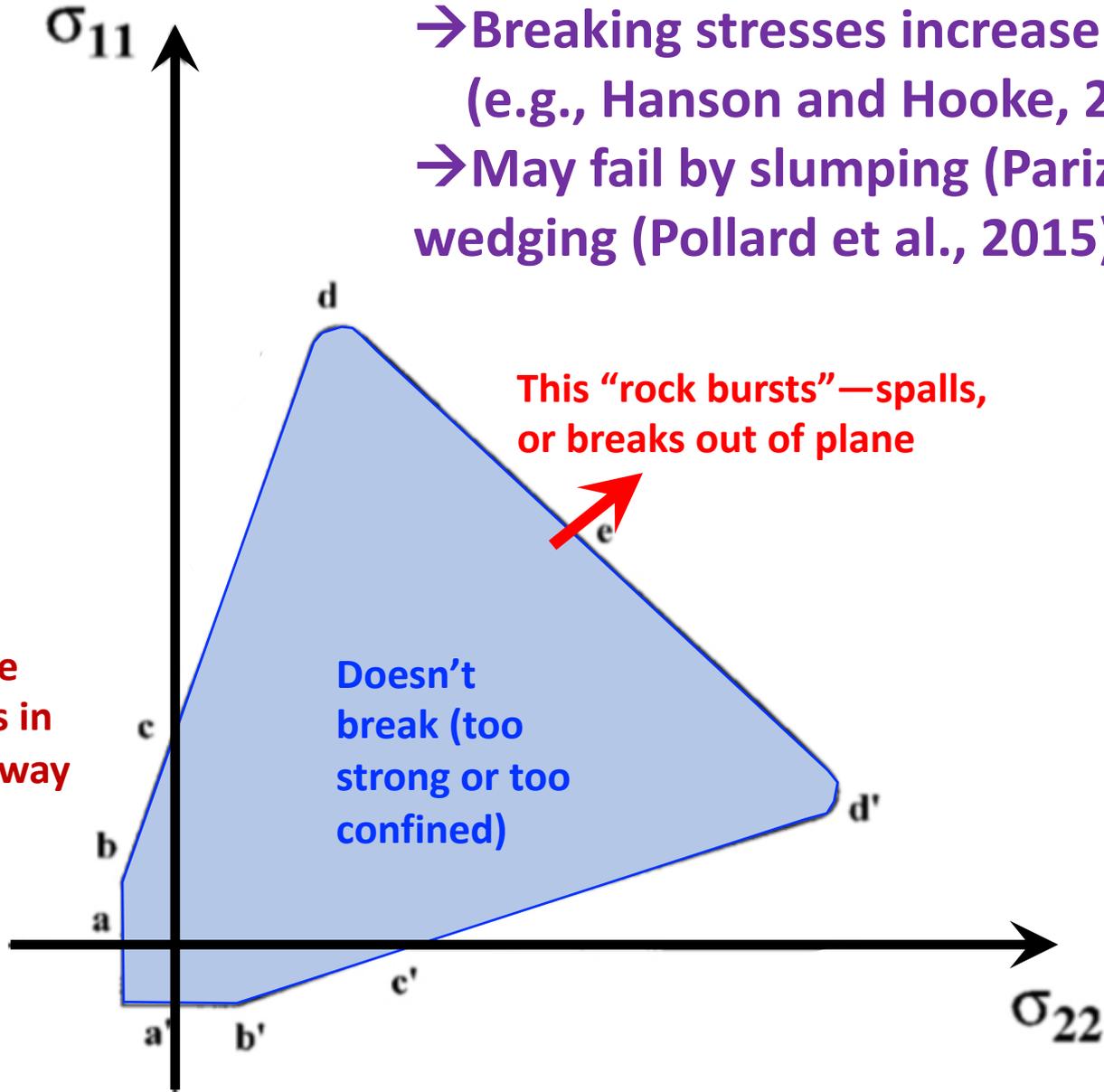






- Free face of a cliff prevents confinement
- Breaking stresses increase with cliff height (e.g., Hanson and Hooke, 2000; 2003)
- May fail by slumping (Parizek et al.), water wedging (Pollard et al., 2015), rock burst...

All else breaks in some way



If creep and subcritical cracking are occurring at a similar rate,
2x cliff height \rightarrow 2x deviatoric stress causing deformation:
creep increases $\sim 10x$; cracking increases $\sim 10^9x$



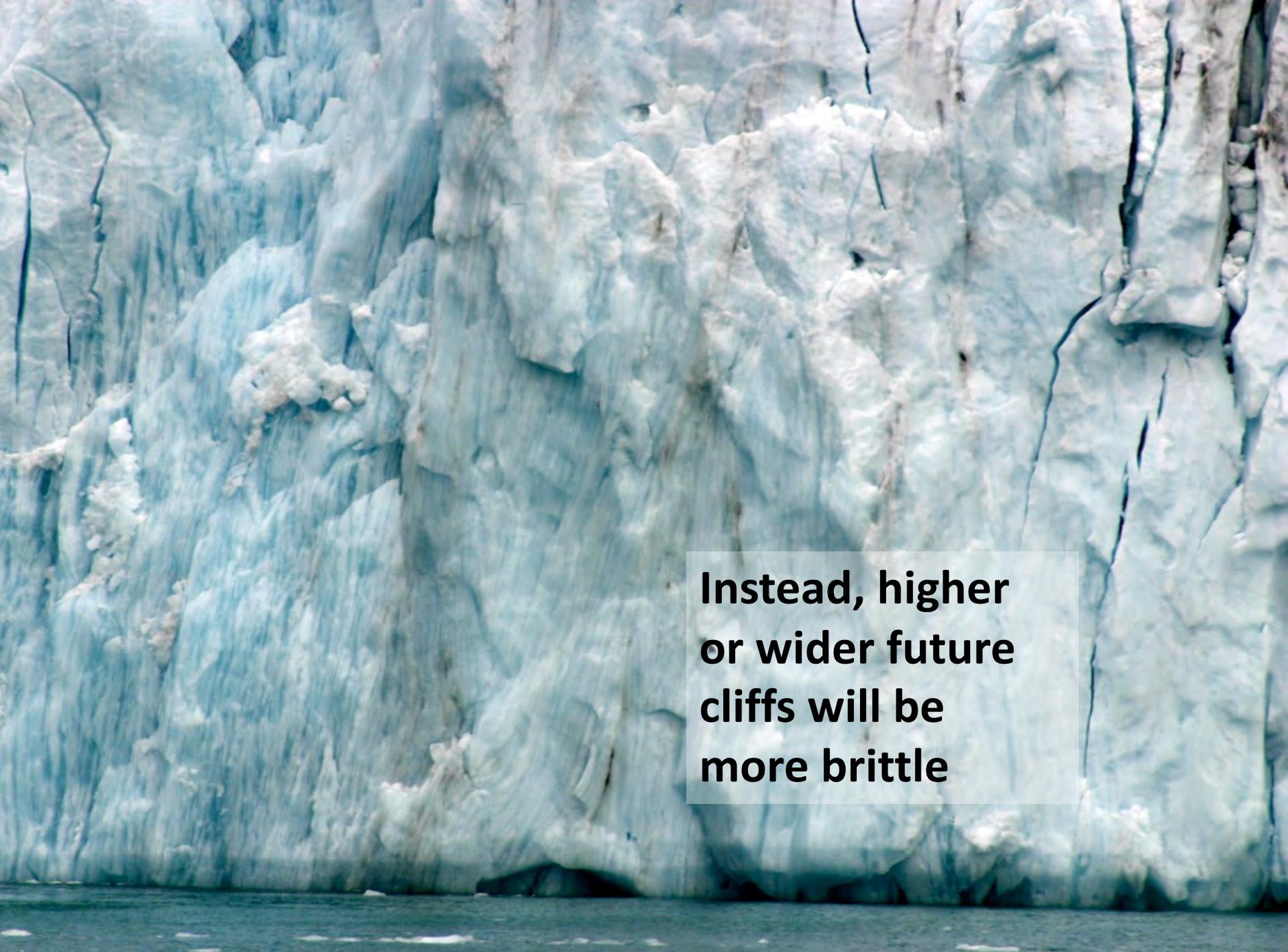
If creep and subcritical cracking are occurring at a similar rate, doubling deviatoric stress causing deformation: creep increases $\sim 10x$; cracking increases $\sim 10^9x$



Brittle-to-ductile or flow-to-fracture transition



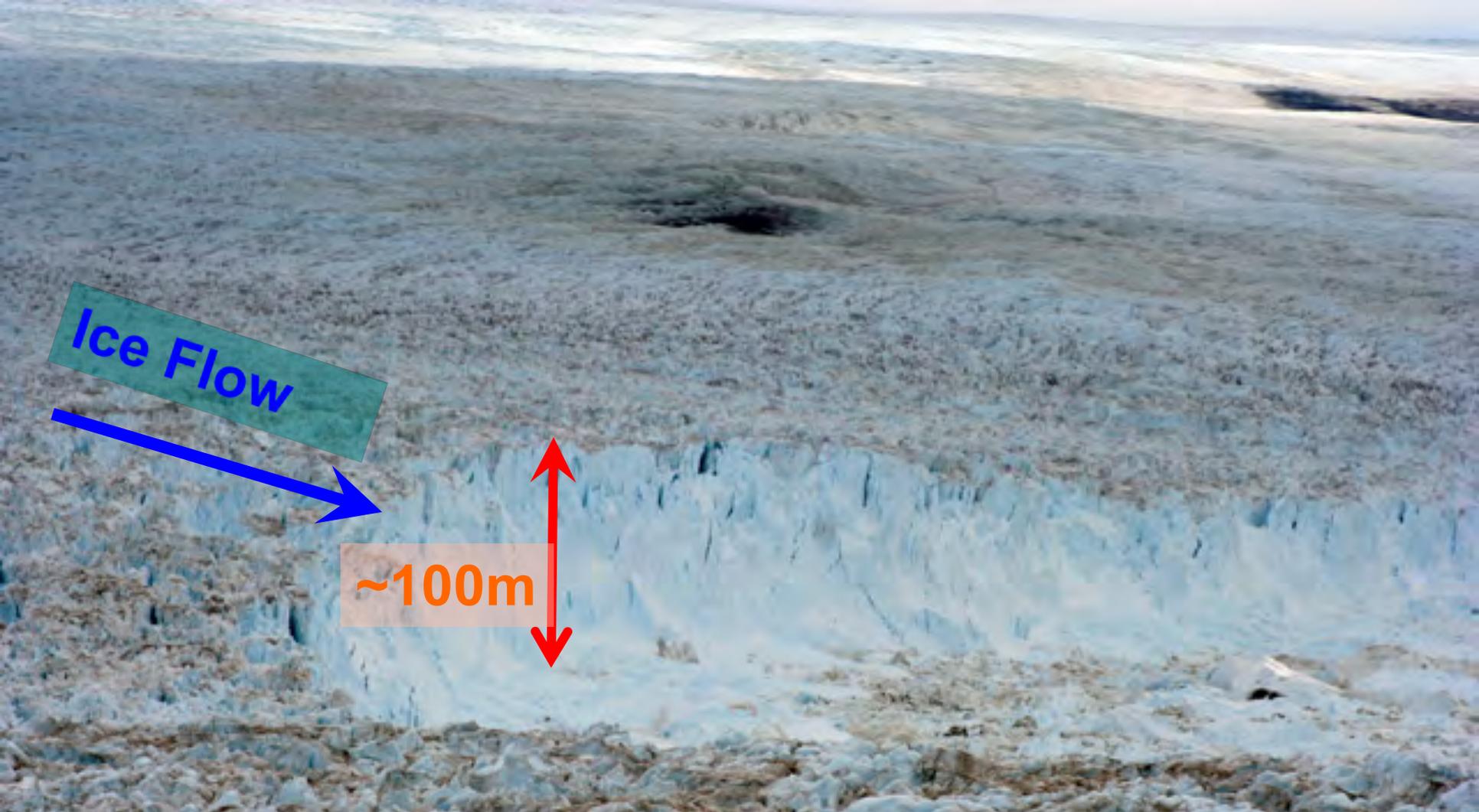
Could higher future cliffs relax fast enough to re-form ice shelves? **No**



**Instead, higher
or wider future
cliffs will be
more brittle**

But does it matter? Yes...





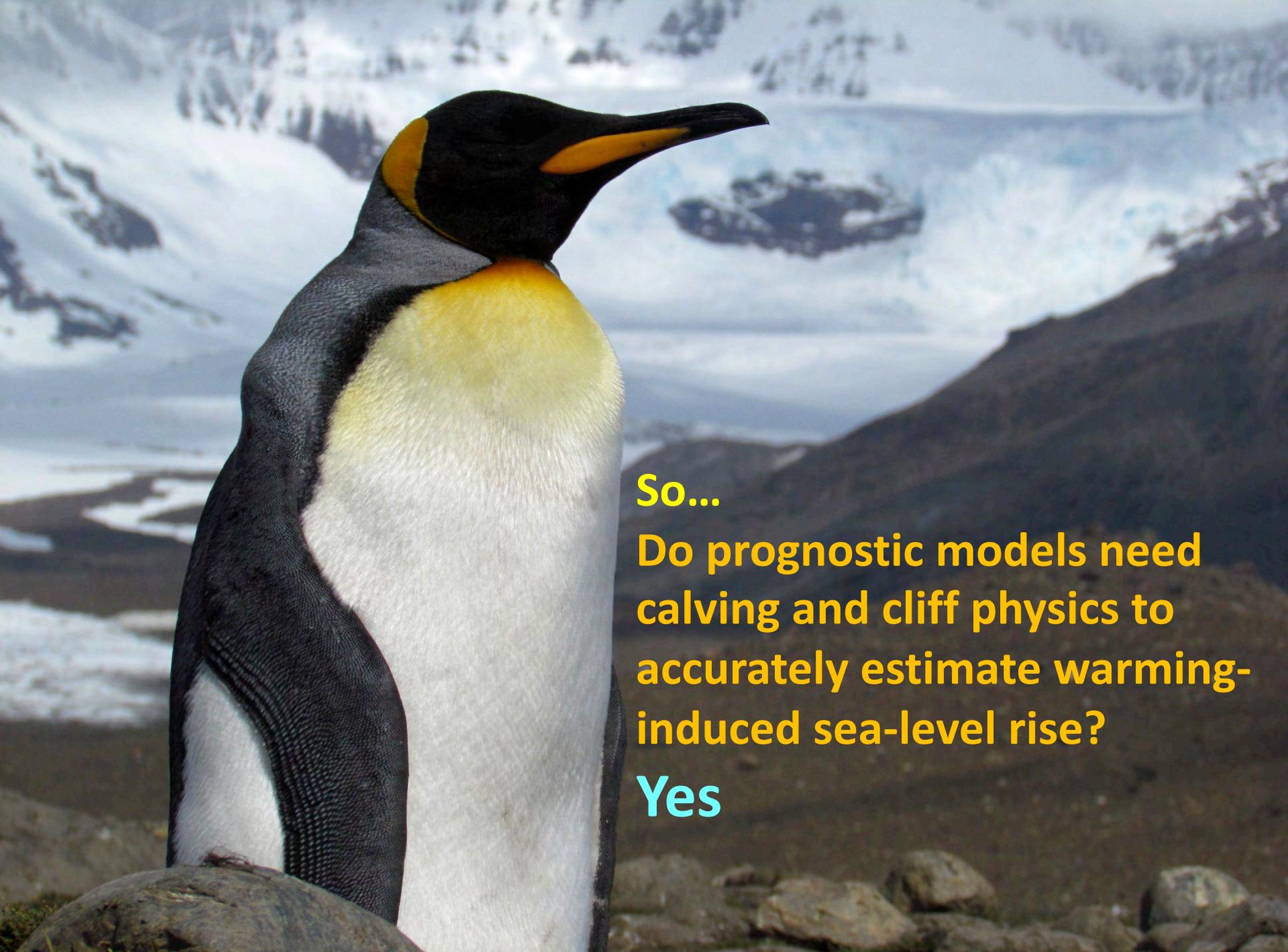
Jakobshavn calves to ~20 m thicker than flotation, reducing drag, speeding flow and thinning (Joughin et al., 2012)



Stresses caused by cliffs enhance thinning through second-invariant effects in flow law (see Parizek et al., AGU; ms. in prep.)



With very rapid failure by water-driven crevassing, slumping, spalling for too-tall cliffs



So...

Do prognostic models need calving and cliff physics to accurately estimate warming-induced sea-level rise?

Yes

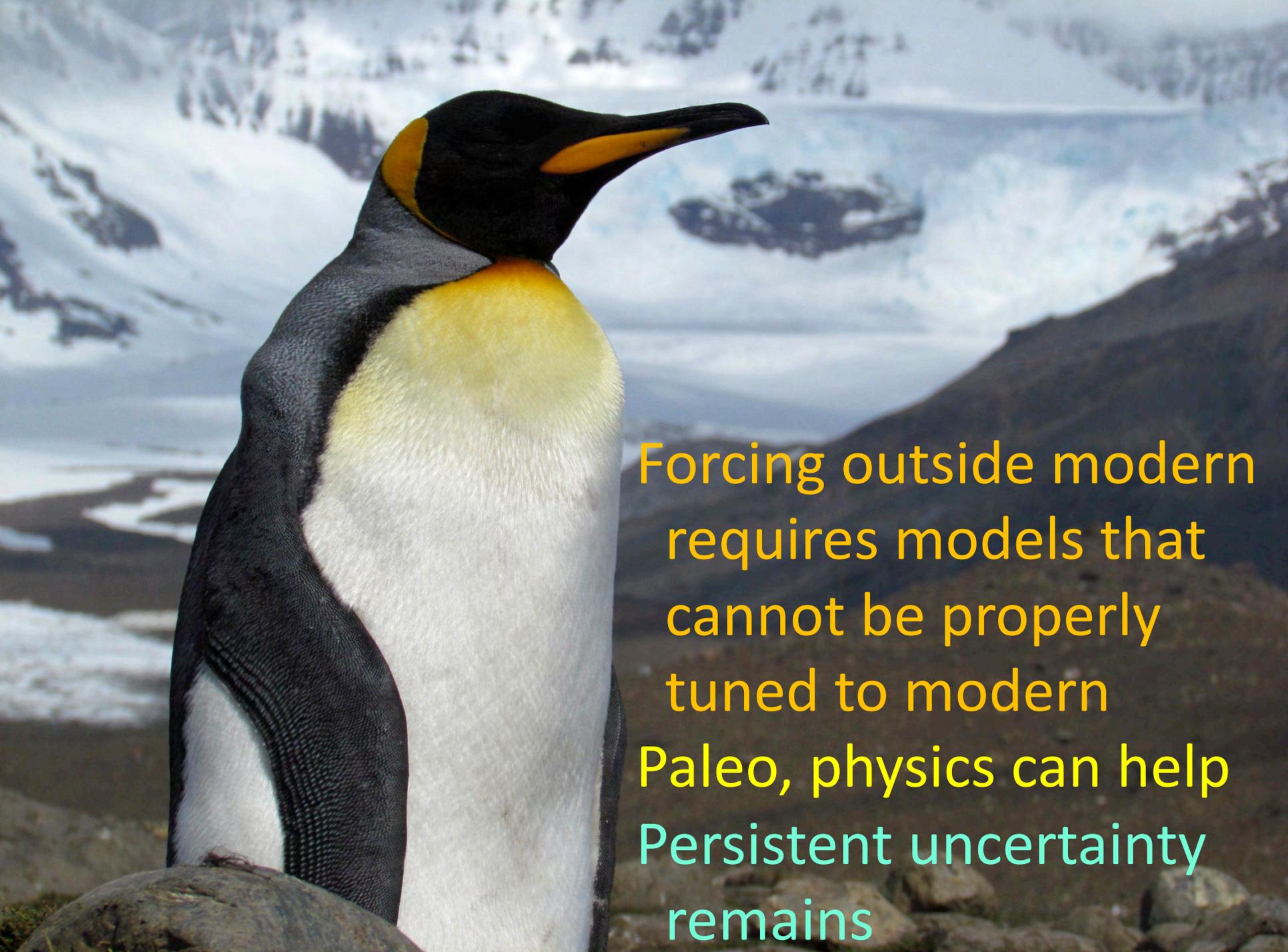
What can we do? Model calving and cliffs...



What can we do? Model calving and cliffs...Or:

- Set ice-shelf buttressing to zero above some temperature**
- Set basal drag to zero a bit inland of flotation (calving to there)**
- Weaken near-grounding-zone ice (second-invariant cliff effects)**
- Set flag—if cliff much taller than 100 m, retreat faster than model**





Forcing outside modern
requires models that
cannot be properly
tuned to modern
Paleo, physics can help
Persistent uncertainty
remains